

1 **Impacts and prognosis of natural resource development on**
2 **water and wetlands in Canada's boreal zone**

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4 Kara L. Webster, Frederick D. Beall, Irena F. Creed and David P. Kreuzweiser
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9 **Kara L. Webster,¹ Frederick D. Beall, and David P. Kreuzweiser.** Natural Resources
10 Canada, Canadian Forest Service, 1219 Queen St. East, Sault Ste. Marie, ON P6A 2E5,
11 Canada.

12 **Irena F. Creed.** University of Western Ontario, Department of Biology, 1151 Richmond
13 St. N., London, ON N6A 5B7, Canada.

14
15 ¹Corresponding author (email: Kara.Webster@nrca.gc.ca).

16 Tel: 705-541-5520

17 Fax: 705-541-5700

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21 **Abstract**

22 Industrial development within Canada's boreal zone has increased in recent decades.
23 Forest management activities, pulp and paper operations, electric power generation,
24 mining, conventional oil and gas extraction, non-conventional oil sand development, and
25 peat mining occur throughout the boreal zone with varying impacts on water resources.
26 We review impacts of these industries on surface water, groundwater and wetlands
27 recognizing that heterogeneity in the dominance of different hydrologic processes (i.e.,
28 precipitation, evapotranspiration, groundwater recharge and runoff generation) across the
29 boreal zone influences the degree of impacts on water resources. Through the application
30 of best management practices, forest certification programs, and science-based
31 guidelines, timber, pulp and paper, and peat industries have reduced their impacts on
32 water resources, although uncertainties remain about long term recovery following
33 disturbance. Hydroelectric power developments have moved toward reducing reservoir
34 size and creating more natural flow regimes, although impacts of aging infrastructure and
35 dam decommissioning is largely unknown. Mineral and metal mining industries have
36 improved regulation and practices, but the legacy of abandoned mines across the boreal
37 zone still presents an ongoing risk to water resources. Oil and gas industries, including
38 non-conventional resources such as oil sands, is one of the largest industrial users of
39 water and, while significant progress has been made in reducing water use, more work is
40 needed to ensure the protection of water resources. All industries contribute to
41 atmospheric deposition of pollutants that may eventually released to downstream waters.
42 Although most industrial sectors strive to improve their environmental performance with
43 regards to water resources, disruptions to natural flow regimes and risks of degraded

44 water quality exist at local to regional scales in the boreal zone. Addressing the emerging
45 challenge of managing the expanding, intensifying and cumulative effects of industries in
46 conjunction with other stressors, such as climate change and atmospheric pollution,
47 across the landscape will aid in preserving Canada's rich endowment of water resources.

48

49 Key words: natural resources, development, hydrology, biogeochemistry, cumulative
50 effects, water quality, water quantity

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53 **1. Introduction**

54 **1.1 Water as a defining feature**

55 Canada is a water rich nation, hosting ~5.4 % (2.9 km³ per year) of the global renewable
56 water supply (5.4 ×10⁴ km³ per year, FAO 2013). It is often reported that Canada
57 contains 20% of the world’s freshwater (Sprague 2007), but this estimate reflects the sum
58 of water stored in lakes, reservoirs, soils and groundwater, and therefore represents the
59 stock of freshwater rather than its renewable supply that is replenished each year from
60 precipitation. Of the water that annually flows into Canada’s oceans, approximately 50%
61 originates in the boreal zone. Seventy-three percent of boreal rivers drain to the Arctic
62 Ocean and Hudson Bay, 20% drain into the Atlantic Ocean, and 7% drain into the Pacific
63 Ocean (Statistics Canada 2010). In addition to the large extent of surface waters,
64 wetlands are a dominant feature of the boreal landscape, occupying 96.5 ×10⁶ ha (17%)
65 of the ecozone and representing close to 20% of the world’s wetlands (NRCan 2013c)
66 (Fig. 1).

67

68 **1.2 Water and its ecosystem services**

69 Canadians value their water resources. A Nanos (2009) research poll revealed 61.6% of
70 Canadians ranked fresh water as the country’s most important resource, ahead of forests,
71 agriculture, oil and fisheries. Water is essential for all living things, and the myriad of water
72 bodies including lakes, ponds, wetlands, rivers, and streams across Canada’s boreal zone
73 support an array of aquatic ecosystem services. These services include groundwater aquifer
74 recharge, contaminant absorption and filtering, river flow regulation through absorbing and
75 releasing excess water, shoreline and erosion protection, clean water supply, and habitat for

76 waterfowl, fish and other biota (Millenium Ecosystem Assessment 2005; Woodward 2009;
77 Wells et al. 2010; Kreutzweiser et al. 2013) (Table 1).

78

79 **1.3. Role and impact of disturbance on boreal water resources**

80 Forest disturbance followed by succession has an important natural role in boreal
81 ecosystems (Venier et al. *in press*). Fire, insect outbreaks, blow downs, and beaver
82 impoundments are all examples of natural disturbances operating on a broad range of
83 spatial and temporal scales that influence boreal water resources. Anthropogenic
84 development creates a different disturbance footprint on the landscape (Table 3). Natural
85 disturbances may be small (blow down) or large (fire, insect) in extent, but typically their
86 impacts are transient in time, such that the disturbance and impacts to water resources are
87 short lived (i.e., <10 years). Anthropogenic disturbances may also be small (e.g., pulp
88 mill) or large (e.g., clearcut, open-pit mine), but typically their geometry is different (i.e.,
89 linear vs. irregular; square or rectangular vs. polygons) and their impact times are longer
90 causing slower or delayed ecosystem recoveries which create legacies for water resources
91 that may be felt for decades or centuries (e.g., acid mine drainage and mine tailings) (Van
92 Geen et al. 1997; Leblanc et al. 2000; National Research Council 2008).

93 Canada's boreal zone is increasingly affected by large-scale industrial activities,
94 with estimates of total anthropogenic disturbance footprint based on interpretation of
95 satellite imagery with approximately 24 million ha (4% of boreal) showing visible forest
96 cutblocks accounting for more than 60% of this disturbance (Pasher et al. 2013) (Fig. 2).
97 Industrial activities increase access for humans to once isolated lakes, rivers, and
98 wetlands through networks of roads, seismic lines, pipelines and transmission lines and

99 can result in removal of water, alteration of hydrologic and biogeochemical flows,
100 increases in erosion and siltation, and increases in pollutants and contaminants in aquatic
101 systems.

102

103 **1.4 Purpose and scope of review**

104 This paper reviews published literature describing the characteristics of boreal water
105 resources, and reviews the status and impacts of natural resource development on the
106 quantity and quality of water resources in the boreal zone and offers prognoses. Industries
107 that use only raw natural resources (i.e., wood, peat, minerals, water) are considered;
108 agriculture (crop or livestock production) and agroforestry are excluded. We review if
109 and how water quantity and quality across the boreal zone are being affected by natural
110 resource development. Consideration is given both to the intensity and extent of resource
111 development impacts and the cumulative effects from combined industrial developments
112 on boreal water resources. We focus on recent (i.e, last 20 years), peer-reviewed literature
113 on studies in the Canadian boreal zone, but include other boreal and non-boreal examples
114 where necessary. In a companion paper, Kreutzweiser et al. (2013) reviewed the impacts
115 and offer prognoses of natural resource development on freshwater aquatic biodiversity
116 in the boreal zone.

117 **2. National perspective of boreal water resources**

118 **2.1 Hydrologic regions**

119 Boreal basins drain into the Atlantic, Pacific, and Arctic oceans. The boreal zone can be
120 classified into hydrologic regions that are defined by the dominance of the different
121 components of the water budget (precipitation, evapotranspiration, discharge) that
122 influence the distribution of wetlands, surface water and groundwater flow pathways, and
123 biotic assemblages (Fig. 3).

124 Devito et al. (2005) proposed a hierarchy of five factors for classifying the
125 dominant controls on water cycling: climate, bedrock geology, surficial geology, soil type
126 and depth, and topography and the properties of the drainage network. Each of these
127 factors operates over different scales (e.g., km² or larger for climate vs. m² for soils) and
128 influences water cycling in different ways. An area as vast as the boreal zone exhibits
129 significant heterogeneity in the factors (and their interactive effects) that control the
130 cycling of water and the sediments and nutrients transported by water through boreal
131 ecosystems.

132 Climate gradients occur both with latitude (primarily a temperature gradient) and
133 longitude (primarily a moisture gradient) (Fig. 4). There is a west to east gradient in
134 annual precipitation, with the western boreal zone receiving considerably less
135 precipitation than the eastern boreal zone (Fig. 4A). When annual precipitation is
136 combined with annual potential evapotranspiration (Fig. 4B), there is an even stronger
137 gradient in the amount of water available for groundwater recharge and surface runoff
138 (Fig. 4C). The seasonality of precipitation must also be considered. For example, in the
139 western boreal zone most precipitation falls in the same months with the highest potential

140 evapotranspiration, whereas, in the eastern boreal zone, precipitation is more evenly
141 distributed throughout the year. Thus, there is a greater potential for drought in the
142 western boreal zone. Broad-scale and re-occurring climate events (e.g., ENSO and PDO)
143 also affect the frequency and intensity of climatic events across Canada (Shabbar and
144 Skinner 2004; Yu and Zwiers 2007).

145 There is also a gradient in geology across the boreal zone. In general, sedimentary
146 rocks dominate in the western boreal zone, whereas intrusive and metamorphic rocks of
147 the Canadian Shield dominate in the north-central and eastern boreal zone. The surficial
148 deposits of the boreal zone are generally dominated by glacial till blanket and veneer,
149 with glaciolacustrine deposits distributed throughout and glaciomarine deposits
150 surrounding the Hudsons Bay Lowlands (Fulton 1995). The soils overlying these deposits
151 vary substantially across the boreal zone (see Maynard et al. 2014).

152 In the western boreal zone, the sub-humid climate, permeable bedrock, deep
153 surficial deposits and relatively flat topography means that hydrology is dominated by
154 vertical fluxes (high evapotranspiration and groundwater recharge), soil water storage
155 and low runoff generation (Fig. 5). In contrast, in the eastern boreal zone, the humid
156 climate, impermeable bedrock, shallow substrates and more complex topography means
157 that the hydrology is dominated by lateral fluxes (high runoff), with limited groundwater
158 recharge and high runoff generation (Fig. 5). The prevalence of different dominant
159 hydrologic processes in different regions suggests that the potential for and severity of
160 alterations to hydrologic regime due to impacts of natural and anthropogenic disturbances
161 will also vary.

162 These same hydrologic regions also explain wetland forms and distribution across
163 the boreal zone. Wetlands are defined as areas with water table at, near or above the land
164 surface for long enough to promote hydric soils, hydrophytic vegetation and biological
165 activities adapted to wet environments (Tarnocai 1980). Wetlands are classified into bog,
166 fen, swamp, marsh or shallow water (National Wetlands Working Group 1997). Wetlands
167 with more than 40 cm depth of peat development are further classified as peatlands
168 (National Wetlands Working Group 1997). Wetlands form as a result of climate and
169 landscape factors that create a positive water balance (precipitation + inflows >
170 evapotranspiration + outflows). Climate factors that promote wetland development
171 include high precipitation and/or low evapotranspiration. Landscape factors associated
172 with wetland development are geomorphic positions that collect water (i.e., depression
173 areas), that drain slowly (i.e., over confined bedrock), or are connected to regional
174 groundwater flows. Minerotrophic wetlands such as shallow open water, fens, marshes
175 and swamps are normally situated at positions in the landscape lower than adjacent
176 hillslopes, such that water and mineral elements are introduced by groundwater or littoral
177 sources in addition to atmospheric sources (Price and Waddington 2000). Thus, in the
178 boreal zone, these minerotrophic wetlands typically form in bedrock depressions (e.g.
179 Canadian Shield), on flat, poorly drained areas (e.g., Western Plains and Hudson Plain),
180 and in discharge zones at the base of long, steep slopes (e.g., Cordillera). Ombrotrophic
181 bogs rely on atmospheric sources of water. They are hydrologically isolated from lateral
182 inflows or upward seepage and thus occur only where precipitation normally exceeds
183 evapotranspiration during the growing season.

184 Taken together, the above-mentioned factors demonstrate the difficulty of making
185 generalizations about hydrologic processes in the boreal zone and emphasize how the
186 “uniqueness of place” (Beven 2000) continues to bedevil attempts at generalization and
187 classification of hydrologic regimes. At present, the science of hydrology does not have a
188 generally accepted classification system, unlike many other disciplines, which would
189 permit the definitions of regions or locales where the dominant hydrologic processes are
190 similar. There have been ongoing calls for the development of a hydrologic classification
191 system (e.g. McDonnell and Woods 2004), and many frameworks have been proposed,
192 built around hydrologic similarity (Wagner et al. 2007), hydrologic response units
193 (Devito et al. 2005) or hydrologic landscapes (Winter 2001). All of the frameworks are
194 similar in that they attempt to capture the co-evolution of climate, bedrock geology,
195 surficial geology, soils, topography, and vegetation and how these factors contribute to
196 the local hydrologic regime. Greater communication among groups within the
197 hydrological community (e.g., scientists and engineers) would help in adopting a
198 common classification. A robust hydrologic classification system would be required to
199 extrapolate results from the spatially sparse research sites to the broader landscape.

200

201 **2.2 Properties and attributes of wetlands in the boreal zone**

202 The boreal zone encompasses both the boreal and subarctic peatland regions (Tarnocai
203 and Stolbovoy 2006), with 67% of the boreal peatlands occurring as bogs, 32% as fens,
204 and the remainder (1%) as swamps and marshes. Two notable extensive boreal peatland
205 areas occur in the Mackenzie River Valley in the northwest and the Hudson Plain in the
206 east, the latter being the single largest peatland complex in North America and the second

207 largest in the northern hemisphere (Gorham 1991; Glooschenko et al. 1994). Wetland
208 type is largely determined by the strength of connection to the groundwater system and
209 the physical and chemical nature of the underlying geology (National Wetlands Working
210 Group 1997). The natural succession of boreal wetlands is from open water to fen and
211 eventually to bog as peat accumulation becomes sufficient to disconnect the peatland
212 from groundwater (Zoltai et al. 1988; Tarnocai and Stolbovoy 2006). In northern parts of
213 the boreal zone, peatlands may contain permafrost (i.e., ground that remains frozen for
214 greater than two years) (Brown 1967). Frozen peat can develop in more southerly
215 latitudes within peatlands where insulating moss (*Sphagnum* spp.) and black spruce
216 (*Picea mariana*) allow frost to persist throughout the year (McLaughlin and Webster
217 2013).

218 Peat is classified as soil having 30% or more organic matter (>17% organic
219 carbon) by weight (Soil Classification Working Group 1998) and has a relatively low
220 bulk density (0.07 to 0.25 Mg per m³) compared to mineral soils (Turchenek et al. 1998).
221 Low bulk density creates high (average of 92%) porosity (volume of pore space as a
222 proportion of total volume). This allows for greater water retention than mineral soils.
223 These properties also affect the movement of water, with high hydraulic conductivity in
224 the top layer of peat (~80% of water is released) and low hydraulic conductivity in deeper
225 peat (~15% of water is released) (Turchenek et al. 1998).

226 Water table levels within wetlands fluctuate seasonally and annually as a result of
227 precipitation patterns and wetland size and connectivity to surface, shallow subsurface
228 and deeper groundwater sources (Pelster et al. 2008). During dry periods, considerable
229 water can be stored in wetlands that provide small, sustained baseflows to streams and

230 evapotranspiration to the atmosphere (Mitsch and Gosselink 2007). In contrast, during
231 wet periods, the capacity of wetlands to store water is exceeded and the saturated areas
232 act as conduits of overland flow from the uplands to the streams. In areas of permafrost
233 (e.g., Hudson Plain, Mackenzie River Valley), groundwater-surface water interactions are
234 restricted to shallow, localized flows in the surface active layer above permafrost during
235 periods of thaw (Woo and Young 2003).

236

237 **2.3 Land-water linkages and their influence on water resources**

238 The interactions of the mass and energy cycles and the physical, chemical and biological
239 processes that drive them must be considered to understand the potential impacts of
240 natural resource development on boreal water resources.

241 Water enters the ecosystem as precipitation. It is first intercepted by surface
242 vegetation, where a portion is returned to the atmosphere by evaporation, and the
243 remainder falls to the ground as throughfall. Water evaporates from plant and soil
244 surfaces or infiltrates and is transpired by vascular vegetation, collectively known as
245 evapotranspiration, and any water remaining is stored in the soil. Upon soil saturation,
246 water flows *via* surface, shallow subsurface or deeper groundwater pathways, creating
247 areas of discharge (upward movement of water to the surface) to surface bodies of water
248 or recharge (downward movement of water) into groundwater storage. Storm and
249 snowmelt intensity (e.g., gentle rain or intense downpour) and the properties of the upper
250 surface of the ground such as soil texture (percentage of sand, silt and clay), amount of
251 organic material (e.g., porous peat materials), and depth of confining layers (e.g.,
252 impermeable bedrock or frozen permafrost layer) will determine the dominant flow paths

253 or route that water will take. In general, surface flows predominate where there is steep
254 relief, shallow soils and/or deeper soils with saturated conditions, and deep flows
255 predominate where there is gentle relief, deep soils and unsaturated conditions (Dunne
256 and Leopold 1978).

257 Forest ecosystems are sources of nutrients to surface water, with important
258 downstream water quality implications (Waring and Schlesinger 1985; Chapin et al.
259 2011). Within forests, water flow pathways influence water chemistry by determining the
260 materials with which water interacts. Water travelling along shallow flow pathways will
261 interact with organic-rich compounds such as freshly fallen leaves and/or roots and forest
262 floor materials, whereas deeper flow pathways interact with mineral-rich parent
263 materials. The residence time of water moving along the flow path will also influence
264 water quality (McGuire et al. 2005). Residence times are governed by presence or
265 absence of storage structures (e.g., wetlands), the gravitational gradient (e.g., steep vs.
266 gentle slope), and the physical structure of the soils (e.g., fine vs. coarse materials). Short
267 residence times create oxic conditions and result in less interaction with soils that may
268 leach into the water than longer residence times that have anoxic conditions and more
269 interaction with soils (McGuire et al. 2005). The residence time of water in turn alters the
270 soil chemical environment, which alters the form and thus also the solubility and mobility
271 of nutrients (Chapin et al. 2011).

272 Water quantity and quality changes throughout the year reflecting changes in
273 precipitation inputs (e.g., snowmelt, summer drought, fall storms), dominant flow
274 pathways, and the dynamic biological and chemical environments along the flow
275 pathways. For example, water chemistry is typically more dilute in spring due to low

276 biological processing under the snowpack and high quantities of available water during
277 snowmelt. In contrast, water chemistry is typically more concentrated in the summer due
278 to high biological processing and low quantities of available water (Band et al. 2001).

279 In the eastern boreal zone, with shallow upland soils overlying the high relief of
280 the Canadian Shield, water quantity and quality is predominantly determined by variable
281 source area (VSA) regulated near-surface flows (Fig. 5). VSA refers to the concept that
282 runoff-generating areas in the landscape vary in size and configuration over time. In
283 VSA-controlled landscapes, as the groundwater table rises to intersect the surface soils,
284 nutrients that have previously accumulated in the surface soils are mobilized and flushed
285 to the stream (Creed and Sass 2011). Topography, *via* VSA control, influences the
286 hydrologic flushing of nutrients in various ways. It affects (a) the generation of nutrient
287 supply (e.g., nutrient-poor areas develop if soil conditions are too dry or too wet); (b)
288 potential expansion *vs.* contraction rates of the VSAs (e.g., catchments with a greater
289 potential for lateral expansion of source areas will have longer flushing times and higher
290 rates of nutrient export, while catchments with less potential for lateral expansion of
291 source areas will have shorter flushing times and lower rates of nutrient export); and (c)
292 the transport of flushable nutrients to surface waters, which is a function of both the size
293 and configuration of the VSA (e.g., catchments with larger, hydrologically-connected
294 VSAs will have larger nutrient export, whereas catchments with smaller or
295 hydrologically-disconnected VSAs will have lower nutrient export or more leaching to
296 groundwater) (Creed and Sass 2011). This mechanism, or variations of it, has been used
297 to explain the export of carbon (Hornberger et al. 1994; Dillon and Molot 1997; Creed et

298 al. 2003, 2008; Richardson et al. 2009; Mengistu et al. 2013), nitrogen (Creed et al. 1996;
299 Creed and Beall 2009; Mengistu et al. 2013), and phosphorus (Mengistu et al. 2013).

300 The western boreal zone is characterized by deeper upland soils and lower relief,
301 and water quantity and quality is typically not regulated by VSA dynamics (Fig. 5). In
302 non-VSA-dominated landscapes, the predominance of primarily vertical flows and deep
303 subsurface pathways regulate water and nutrient transfer from land to aquatic systems
304 (Creed and Sass 2011). Thus, consideration has to be given to local, intermediate and
305 regional groundwater flows to understand the chemistry of surface waters. In boreal
306 forests with these complex hydrogeological systems, predicting the chemical responses of
307 surface waters to disturbance is extremely difficult (Devito et al. 2000). However, the
308 size and configuration of surface saturated areas where water flowpaths created by these
309 complex hydrogeological systems intercept with the land surface have been used to
310 explain total phosphorus loading to boreal surface waters (e.g., Devito et al. 2000; Evans
311 et al. 2000) and trophic status of lakes (Sass et al. 2007, 2008a, 2008b) in the western
312 boreal zone.

313 Wetlands can have a large impact on water quality, even if they occupy only a
314 small percentage of basin area. Wetlands are unique in that they act as reservoirs of water
315 and bioreactors of nutrients during drier periods and conduits of water and nutrients
316 during wet periods. In particular, there is a strong connection between the presence and
317 extent of wetlands and the concentration of dissolved organic carbon in surface waters
318 (Dillon and Molot 1997; Creed et al. 2003; 2008). Similarly, wetlands have been shown
319 to have an important role in nitrogen and phosphorus loading to surface waters (e.g.,

320 Creed and Beall 2009; Mengistu et al. 2013) as well as contributors of mercury to aquatic
321 ecosystems (Branfireun et al. 2005).

322

323 **2.4 Climate regulation and change**

324 Canada's water-rich boreal zone, including its glaciers, permafrost, snowpack, streams,
325 wetlands, lakes, and groundwater, not only regulates climate, but also is sensitive to
326 predicted changes in climate. Boreal ecosystems regulate weather and climate through
327 both direct mechanisms such as transpiration cooling and albedo effects (Oke 1978;
328 Barnett et al. 2005) and indirect mechanisms such as freshwater inputs to the Arctic
329 Ocean (Bates et al. 2008) and carbon sequestration (Webster and McLaughlin 2013; Kurz
330 et al. 2013).

331 Changes to global climate and hydrologic cycles have the potential to strongly
332 impact boreal water resources (Price et al. 2013). The key impacts from climate change
333 on the water cycle will come in the form of extreme events such as droughts and floods,
334 seasonal shifts in flow regimes, and reduced winter ice coverage (NRTEE 2010). These
335 hydrologic changes to the water cycle are likely to have positive feedbacks (i.e., negative
336 outcomes). Climate models show that changes in temperature and precipitation are likely
337 to continue affecting the partitioning of water between evapotranspiration and runoff as
338 well as the amount of water stored in glaciers, snowpack, lakes, wetlands, soils and
339 groundwater. Recent modelling studies suggest that the interior of North America,
340 including the western boreal zone, will become much drier and may no longer be able to
341 support forests (Hogg 1994; Dai 2011, 2013; Price et al. 2013). These large scale

342 vegetation changes are likely to have large impacts on water resources (Baldocchi et al.
343 2000).

344

345 **3. Risks from natural resource development to water**

346 **resources**

347 **3.1. Introduction**

348 Water is critical to the operation and future development of the energy, forest, and mining
349 sectors, all of which contribute to Canada’s economic wealth and social well-being
350 (NRTEE 2010). The natural resource sectors make a significant contribution to the
351 Canadian economy representing 12.5% of the country’s gross domestic product (GDP) in
352 2009 (Table 2). In 2005, the natural resource sectors, which include thermal power
353 generation, oil and gas, mining, forestry and agriculture, accounted for approximately
354 84% of both Canada’s gross water use (the total volume withdrawn from water bodies)
355 and total water consumed (withdrawn and not replaced) (Statistics Canada 2010). The
356 thermal power-generating sector was responsible for the greatest gross water use, while
357 agriculture accounted for the greatest water consumed (Statistics Canada 2010; NRTEE
358 2010). The current extent of the anthropogenic footprint related to natural resource
359 development in Canada’s boreal zone (Fig. 2) is expected to increase within these sectors
360 by 50 to 60% by 2030, placing even more pressure on Canadian water resources (NRTEE
361 2010).

362 **3.2. Road development**

363 ***3.2.1 Introduction***

364 Transportation infrastructure is an artefact of humans interacting with the landscape
365 (Coffin 2007). Roads are necessary for servicing all natural resource industries and the
366 network of roads that cut across the boreal landscape renders vast areas of land as “road-

367 affected” (Coffin 2007). Pasher et al. (2013) estimated that linear disturbance features
368 across the boreal totalled approximately 600 000 km, with roads and seismic exploration
369 lines contributing to more than 80% of the linear disturbance. There is little scientific
370 literature on impacts of roads in the boreal zone even though roads are a pervasive linear
371 disturbance that has no natural corollary. They create gaps, fragment habitats, alter access
372 to humans and predators, provide vectors for introduction of invasive species, change
373 local climate, and influence local hydrology (Trombulak and Frissell 2000). The effects
374 of roads often last well beyond their construction (or even their use), with older roads
375 tending to have greater effects on water resources because of the places in the landscape
376 where they were constructed and because construction practices were more intrusive than
377 more recently constructed roads (Wemple et al. 1996, 2001; Lugo and Gucinski 2000;
378 Wemple and Jones 2003; Dutton et al. 2005). These effects persist as long as the road
379 remains a physical feature, continuing to alter flow routing even after abandonment and
380 revegetation (Trombulak and Frissell 2000).

381

382 **3.1.2 Water quantity.** Road construction associated with natural resource development
383 influences water quantity through the creation of road surfaces, cutbanks, ditches, and
384 culverts (Jones and Grant 1996) that contribute to changes in the timing and routing of
385 runoff (Trombulak and Frissell 2000). Road surfaces create increased overland flow due
386 to an impervious surface layer that creates reduced infiltration capacity (La Marche and
387 Lettenmaier 2001; Tague and Band 2001). Cutbanks intercept upslope subsurface flow
388 (La Marche and Lettenmaier 2001). The depth of the road cut, which varies with local
389 slope, and width of the road will determine the amount of subsurface runoff that is

390 intercepted (Tague and Band 2001). Ditches and culverts provide a conduit of flows
391 directly to the channel, increasing the flow routing efficiency by extending the drainage
392 network (Wemple et al. 1996). Road location on the landscape, specifically position on
393 hillslope profile, influences the effects of roads on hydrology (Jones et al. 2000).
394 However, there are instances where, instead of concentrating flow, ditches and culverts
395 diffuse flow if water is routed to relatively dry areas. Thus, the effect of this
396 concentration of flow will depend upon the characteristics of the receiving area (Tague
397 and Band 2001). Wetland habitats can be destroyed or created with alterations to surface
398 or subsurface flow (Trombulak and Frissell 2000). For example, wetland road crossings
399 may block drainage passages and groundwater flows, effectively raising the upslope
400 water table and lowering the downslope water table (Forman and Alexander 1998;
401 Partington and Gillies 2010).

402 Roads, in combination with forest harvesting, can have dramatic impacts on both
403 the timing and volume of water during storms (Jones and Grant 1996). However, despite
404 some generalizations that roads have significant effects on peak flow, the impacts will
405 vary with underlying geology, soil texture, vegetation, type of road construction, and both
406 seasonal precipitation and local storm events (Lugo and Gucinski 2000; Tague and Band
407 2001). For example, in regions with high precipitation and high relief landscapes (e.g.,
408 much of the eastern boreal zone, with the exception of the Clay Belt), the net effect of
409 roads is typically a more rapid delivery of water to stream channels during storms,
410 resulting in a higher and earlier peak in storm flow (Jones and Grant 1996; LaMarche and
411 Lettenmaier 2001). Hydrologic effects are likely to persist for as long as the road remains

412 a physical feature altering flow routing (Trombulak and Frissell 2000). Little effort has
413 been expended in understanding the impacts of roads in low-relief landscapes.

414

415 **3.1.3 Water quality.** Roads affect water quality by increasing inputs of dust, generating
416 higher erosion rates and therefore sediment loads, and increasing the turbidity of the
417 waters (Reid and Dunne 1984; Partington and Gillies 2010) thereby impacting aquatic
418 biota (Kreutzweiser et al. 2013). Travel intensity, road surface type, vegetation cover,
419 climate, geologic substrate, road maintenance, and road-stream connectivity are primary
420 factors in regulating sediment production in road systems (Lugo and Gucinski 2000).
421 High rates of sediment production from road surfaces occur in the year immediately
422 following road construction but diminish rapidly over time (Wemple et al. 2001).
423 Trombulak and Frissell (2000) identified five classes of chemicals that roads contribute to
424 the environment: heavy metals, organic molecules, ozone, nutrients, and chemicals for
425 de-icing and dust control. This study indicates that most contamination declines within 20
426 m of the road but that elevated levels of contaminants often occur 200 m or more from
427 the road (Trombulak and Frissell 2000).

428 **3.1.4 Prognosis**

429 Roads will continue to be a part of industrial landscapes of the boreal zone. It is likely
430 that road networks will become denser and expand farther north in the boreal zone as
431 natural resource exploration, development, and the associated infrastructure increase.
432 Consequently, the influence on water resources from road construction is also likely to
433 increase. Road construction has a lasting legacy. Even if access to roads is restricted, the
434 road footprint never completely disappears, and hydrologic flows continue to be affected.

435 There is potential to offset the impacts of roads to some degree by mitigation measures.
436 Recent technological advances enable land planners and managers to improve the
437 location of roads, the location and installation of culverts, and the selection of appropriate
438 road surface materials on roads near water. For example, the Government of Alberta has
439 used a technique for mapping wet areas to generate databases available to public to
440 optimize road design (White et al. 2012). In collaboration with Ontario Ministry of
441 Natural Resources, FP Innovations has summarized the tools and resources available for
442 best management practices for roads (Partington and Gillies 2010) and state-of-practice
443 review (Gillies 2011) focussing on the unique challenges for road construction through
444 wetlands.

445 **3.2 Forest Management**

446 ***3.2.1 Introduction***

447 The forest sector contributed \$23.6B (1.9%) to Canada's GDP, \$17.2B to
448 Canada's trade balance, and 233 900 direct jobs in 2011 (NRCan 2013a). Forest
449 management activities occur in all provinces, and to a much more limited degree in the
450 Yukon and Northwest Territories, with the boreal zone host to a significant fraction of
451 forest management activities in Canada (Fig. 2C). Within the spectrum of forest
452 management activities, including timber harvesting, site preparation, planting,
453 competition management, and stand tending, only timber harvesting activities have
454 received significant research attention with regards to their impact on water resources.
455 Timber harvesting in the boreal zone typically employs clearcut harvesting (e.g., usually
456 clearcuts of less than 260 ha in Ontario's boreal; OMNR undated) systems and post-
457 harvest treatments to emulate the natural pattern of fire-origin, even-age stands of the

458 dominant species (generally spruce, jack pine, lodgepole pine, and aspen) (Long 2009;
459 Sibley et al. 2012). Forest management strategies in the boreal zone have evolved and
460 large, regularly shaped cuts typical of the past have been replaced with smaller,
461 irregularly shaped cuts containing residual material to address biodiversity concerns
462 (Potvin et al. 1999; Venier et al. *in press*). An emerging issue in boreal forestry is
463 biomass harvesting for bioenergy. Until relatively recently, slash (non-merchantable
464 wood comprised of small diameter trees, tops and branches) from the harvest was either
465 left on the ground on-site or de-limbed at the road. Now there are opportunities for use of
466 this biomass in alternative wood products (e.g., wood pellets and chips) or in energy
467 production (Paré et al. 2011). The removal of this material has implications for site
468 productivity and water resources (Thiffault et al. 2010).

469 Our understanding of how timber harvesting affects water resources is based
470 largely on a century of comparison studies of harvested vs. unharvested paired-basin
471 studies (c.f. Brown et al. 2005 for a recent meta-analysis of these studies). The primary
472 assumption underlying paired-basin studies, that changes in streamflow between a
473 harvested and unharvested basins with similar climate, geology and vegetation isolate and
474 quantify the impacts of harvesting, has proven to be a powerful tool in understanding the
475 effects of timber harvesting. Canada has a long history of paired-basin studies (Buttle et
476 al. 2000; Krezek et al. 2008; Mallik and Teichert 2009). However, not all of these studies
477 have forest management activities as a primary focus and only a small proportion are
478 located in the boreal zone (Fig. 3, Table 4). These paired-basin studies, combined with
479 stand- or site-level studies, have contributed greatly to our understanding of the effects of
480 timber harvesting of water quantity and quality.

481

482 ***3.2.2 Water quantity***

483 In general, reduction of forest cover by timber harvesting (and other natural and
484 anthropogenic disturbances) reduces canopy interception and evapotranspiration, leading
485 to increases in soil moisture (Buttle and Metcalfe 2000; Guillemette et al. 2005; Mallik
486 and Teichert 2009). The characteristics of the topography, soil, and surficial geology will
487 determine how the increase in soil moisture is expressed in terms of increases in
488 groundwater recharge *vs.* runoff. The processes connecting the changes in coverage,
489 composition, and structure of tree canopies due to harvesting and the increase in
490 groundwater recharge or runoff are numerous, primarily driven by the water balance and
491 radiation balance (National Research Council 2008). There are, of course, many localized
492 processes and the exceptions to these general principles demonstrate the importance of
493 the particular combination of climate, geology, topography, and soils (Devito et al. 2005)
494 on the response of a particular basin to forest management activities.

495 Forest canopy intercepts precipitation where a portion is subsequently returned to
496 the atmosphere by evaporation or delivered to the ground surface as throughfall and
497 stemflow. The amount of precipitation intercepted varies with tree species. Elliott et al.
498 (1998) reported that interception accounted for 58%, 46%, and 16% of precipitation in
499 mature white spruce, jack pine, and aspen stands, respectively, in north central
500 Saskatchewan. Price et al. (1997) reported that black spruce stands in northern Manitoba
501 intercepted between 15 and 60% (seasonal average of 23%) of precipitation, with higher
502 rates of interception occurring during smaller precipitation events (Price et al. 1997).

503 In the boreal zone, snowfall is a significant contributor to the water balance of
504 most basins. Pomeroy et al. (1998) found that snow interception also differs among forest
505 types, with largest interception in black spruce stands, intermediate in jack pine stands,
506 and smallest in mixed wood (aspen and white spruce) stands. Sublimation of the
507 intercepted snow was approximately 38 to 45%, 30 to 32%, and 10 to 15% for black
508 spruce, jack pine, and mixed wood stands, respectively (Pomeroy et al. 1998). Comparing
509 mature, clearcut, and regenerating stands in northern Saskatchewan, Pomeroy and
510 Granger (1997) found that clearcuts accumulated the most snow, followed by mature
511 mixed wood stands, mature jack pine stands, and regenerating stands (15 years old) had
512 the least accumulation. Similar results were found in other regions, including conifer
513 stands in British Columbia (Winkler et al. 2005) and Alberta (Golding and Swanson
514 1986), where the effect of forest removal was to increase snow accumulation, resulting in
515 greater delivery of water to the soil substrates.

516 The size of canopy openings in relation to the surrounding forest influences the
517 radiation balance within forest stands. Increases in solar radiation have pronounced
518 effects on evapotranspiration and snowmelt (Winkler 1999). In a study of a variety of
519 forest types in northern Manitoba, Metcalfe and Buttle (1999) showed that increasing
520 canopy density, measured as gap fractions, reduced incoming short-wave radiation and
521 wind speeds, reducing latent and sensible heat fluxes and subsequent melt rates. Pomeroy
522 and Granger (1997) found that clearcuts underwent an earlier and more rapid melt. Faria
523 et al. (2000), in sites in central Saskatchewan, found that the rate of snow melt was
524 inversely correlated with snow water equivalent (higher melt rates with lower snow water
525 equivalents), and covariance between melt rate and snow water equivalent led to the

526 largest acceleration of snow cover depletion in medium density stands and the smallest
527 acceleration in high density stands.

528 The impacts of timber harvesting on runoff related to the water and radiation
529 balances, together with site-specific changes to water flow pathways associated with land
530 surface erosion, compaction, rutting, and loss of organic matter, have been the subject of
531 many general reviews (e.g., Bosch and Hewlett 1982; Stednick 1996; Brown et al. 2005),
532 including several examining the Canadian experience (Buttle et al. 2000, 2005, 2009;
533 Moore and Wondzell 2005; Mallik and Teichert 2009). The majority of these studies
534 focused on how timber harvesting affects peak flows because of concerns for flooding
535 and associated increases in stream scouring and bank undercutting, which in turn can
536 affect water quality and aquatic habitats downstream through transport and subsequent
537 deposition of sediment (Alila and Beckers 2001). Studies examining the effects of timber
538 harvesting on base flows and long-term (> five years) recovery are largely absent.

539 In the eastern boreal zone, dominated by a hydrologic regime characterized by
540 high precipitation and relatively high relief, runoff responses to forest management are
541 extremely variable, ranging from negligible to moderate depending on the amount of
542 basin disturbed and the prevalence of skid trails and road ditches. The “Ruisseau des
543 Eaux-Volées” Experimental Watershed (REVEW) in boreal Quebec, forested primarily
544 by balsam fir but with some white spruce, white birch, and black spruce, has been the site
545 of a number of studies on timber harvesting impacts of streamflow. Plamondon and
546 Ouellet (1980) reported on the effects of harvesting 31% of a 394 ha basin and found no
547 changes in annual, seasonal, and monthly runoff, and no changes in the timing of peak
548 and base flows. A subsequent study demonstrated that harvesting 85% of a 122 ha basin

549 increased peakflow by 63% and reduced lag times in the storm hydrograph (Guillemette
550 et al. 2005). The relatively large peak flow response, when compared to 50 other similar
551 studies globally, was attributed to increased channelized flow to the basin outlet due to
552 skid trails and road ditches (Guillemette et al. 2005). Another REVIEW study compared
553 harvesting 50% of four small basins (< 50 ha) with harvested areas distributed at different
554 distances from the stream and found that the configuration of the harvests had little effect
555 on peak flows (Tremblay et al. 2008). A study of larger basins (400 to 11 900 ha) in the
556 boreal forest of northeastern Ontario, with uplands dominated by black and white spruce
557 with minor occurrence of balsam fir, jack pine, white birch, and trembling aspen, and
558 lowland areas dominated by tamarack and black spruce, where the proportion of the
559 basins harvested varied from 5 to 25%, reported no definitive effects on annual runoff or
560 peak flow timing or magnitude (Buttle and Metcalfe 2000). They did, however, report
561 that there was some augmentation of base flows in summer months and suggested the
562 muted responses reflected the ability of large basins to buffer the response to timber
563 harvesting. Metcalfe and Buttle (1999) noted that water storage and evaporation in small
564 wetlands and ephemeral surface depressions is a fundamental component of the basin
565 water balance that may influence runoff response.

566 In the western boreal zone, dominated by a hydrologic regime characterized by
567 low precipitation and relief, runoff responses were more difficult to discern. In the aspen
568 dominated mixedwood forest of northeastern Alberta, Devito et al. (2005) found that
569 harvesting effects were largely obscured by interannual variation in precipitation and soil
570 water storage potential. In these areas, where deep sedimentary deposits are interspersed
571 with clay lenses, greater infiltration and increased groundwater recharge lead to raised

572 water tables following harvest (Smerdon et al. 2009) rather than creating runoff. Runoff
573 generation only occurred under exceptional circumstances (one in 20 years) and within
574 specific landscape units (ephemeral draws and wetlands) that were only rarely
575 hydrologically-connected to surface waters.

576 Swanson and Hillman (1977) examined the effects of timber harvesting in nine
577 control and nine harvested basins in the foothills of the Rocky Mountains in west-central
578 Alberta, where precipitation and relief are much greater than the boreal plains to the east.
579 In this lodgepole pine – white spruce forest, they reported that snowmelt runoff increased
580 59%, growing season runoff increased 27%, and peak flow response to summer storms
581 increased 1.5 to 2 times. In another foothills basin (Streeter Creek), a 50% harvest
582 distributed as small (<1 ha) cutblocks increased growing season runoff by 175%
583 (Swanson et al. 1986).

584

585

586 **3.2.3 Water quality**

587 Timber harvesting causes many physical, geochemical, and biological changes to
588 a forest ecosystem. Alterations to the soil's temperature, structure and nutrient cycling
589 (Table 5) can affect the character of soil pore water, which is subsequently reflected in
590 surface water chemistry as water is transported through the landscape by hydrologic
591 flows, impacting downstream energy and nutrient cycling (Buttle et al. 2000) and aquatic
592 biodiversity (Kreutzweiser et al. 2013).

593

594 **3.2.3.1 Water temperature.** The impacts of timber harvesting on stream water
595 temperatures are variable. Timber harvesting results in higher incident radiation on the
596 soil surface that results in an increase in soil temperatures. Timber harvesting also results
597 in changes in wind patterns that influences soil temperatures. Changes in soil temperature
598 influence temperatures of saturated and unsaturated water flow and, thus, the
599 temperatures in downstream aquatic ecosystems (Johnson and Jones 2000). Temperatures
600 in small headwater streams in sub-boreal Engelmann spruce and subalpine fir forest
601 ecosystems of British Columbia remained four to six degrees warmer and diurnal
602 temperature variation remained higher than in the control streams regardless of riparian
603 buffer retentions five years after the completion of timber harvesting treatments
604 (MacDonald et al. 2003a). Although initially the high-retention treatment acted to
605 mitigate changes in temperature, successive years of wind throw reduced forest canopy
606 density creating temperature impacts equivalent to a clearcut (MacDonald et al. 2003a).
607 In contrast, in more eastern parts of the boreal zone, Tremblay et al. 2009 found that
608 summer daily maximum and minimum stream temperatures remained within $\pm 1^{\circ}\text{C}$ in

609 balsam fir dominated forests that were clearcut but had a 20 m riparian buffer. Similarly,
610 Kreutzweiser et al. (2009b) found that only at the most intensively harvested riparian
611 buffers were there significant ($\sim 4^{\circ}\text{C}$) increases in daily maximum stream temperatures
612 for about six weeks in the first summer in boreal mixedwood stands after logging than in
613 pre-logging years or reference sites, after which temperatures returned to normal.
614 Steedman et al. (2001) found that it was only early summer littoral water temperatures of
615 small lakes in clearcut shorelines of northwestern Ontario that were associated with
616 increases of 1 to 2°C in maximum temperatures and increases of 0.3 to 0.6°C in average
617 diurnal temperature range, compared with undisturbed shorelines or shorelines with 30 m
618 riparian buffer strips three years after harvesting. Timber harvesting has been shown to
619 have variable effects on water temperatures, with the largest impacts usually in early
620 season of the year following canopy openings when the riparian canopy is drastically
621 reduced either by harvesting or post-harvest wind throw.

622

623 **3.2.3.2 Suspended sediments.** Erosion occurs when there has been physical disruption to
624 the soil surface (e.g., from rutting or mechanical site preparation) and where local
625 topography results in increased soil moisture or channelized flows that cause slope failure
626 and creates delivery channels to receiving waters (Hutchinson and Moore 2000; Lewis et
627 al. 2001; but see MacDonald et al. 2003b). Erosion results in the transport of sediment
628 into adjacent streams, wetlands and lakes. This material takes with it nutrients that are
629 embedded within or sorbed (*i.e.*, adhered) onto it (Prepas et al. 2006). MacDonald et al.
630 (2003b) found modest increases in suspended sediments during snowmelt in headwater
631 streams in two harvested basins immediately post-harvest, but declined rapidly within

632 two or three years. Tremblay et al. (2009) also observed increased suspended sediment
633 concentrations post-harvest, although there was not enough pre-harvest data to
634 statistically confirm this result. Kreutzweiser et al. (2009a) measured fine sediments in
635 three harvested boreal catchments that included partially-harvest riparian buffers, and
636 detected significant post-harvest increases in fine sediments at only one of three sites, and
637 only in the first year after harvest. In addition to the sediments from erosion, Steedman
638 and France (2000) found that wind-blown sediment from black spruce - jack pine
639 clearcuts, roads, and skid trails, where soil disturbance is large and able to reach the lake,
640 may have led to elevated levels of littoral sedimentation, although it was thought this
641 mechanism would not cause important changes in water quality. Although some studies
642 do not examine sedimentation directly, many have found that water clarity is reduced by
643 timber harvesting over a variable range in time (one to four years) following harvest
644 (Carignan et al. 2000; France et al. 2000; Steedman 2000; Steedman and Kushneriuk
645 2000; Knapp et al. 2003; Garcia et al. 2007; Bertolo and Magnan 2007; Winkler et al.
646 2009) although one study found no effect on water clarity (Prepas et al. 2001a).

647

648 **3.2.3.3 Dissolved organic carbon and nutrients.** Changes occur to soil nutrient cycles
649 post-harvest (see review by Kreutzweiser et al. 2008). Nutrient cycles such as carbon,
650 nitrogen and phosphorus are driven primarily by microbial activities. Timber harvesting
651 affects both the environmental constraints on and the sources of organic material for
652 decomposition (Kreutzweiser et al. 2008). The warm, mesic and aerated conditions
653 created from timber harvesting operations are ideal for microbial metabolism and their
654 mineralization of organic matter. The rate and timing of decomposition is also

655 constrained by the amount of organic material left following harvest and its quality (i.e.,
656 more easily degradable leaf or needle litter and fine roots vs. less easily degradable
657 branches and stems) (Hazlett et al. 2007). The amount and type of residue (slash) retained
658 on the site will differ depending on timber harvest approach and whether whole tree (de-
659 limbing at road, less residue in plot) or tree-length (de-limbed where cut, more residue on
660 plot) methods were used. Following timber harvest, the demand for nutrients from
661 vegetation is low because of the removal of plants and the lag time in regeneration
662 (Kreutzweiser et al. 2008). Thus, high concentrations of nutrients can accumulate in the
663 soil and soil pore water, creating the potential for leaching into groundwater or loading to
664 downstream ecosystems (Kreutzweiser et al. 2008).

665 ***Dissolved Organic Carbon (DOC):*** Disruption of the forest floor containing
666 easily decomposed residues (leaves and bark) or damage to wetlands where there is
667 deeper organic layers in combination with changes to soil environmental conditions that
668 enhance microbial activity following harvesting can create hydrologically mobile sources
669 of dissolved organic carbon (DOC). A rise in groundwater level following timber
670 harvesting can mobilize DOC both during peak and base flow conditions (Laudon et al.
671 2009). Most studies reviewed found an increase in surface water DOC concentrations
672 following harvest that were maintained for a few years followed by a decline towards
673 pre-harvest conditions in lakes (Carignan et al. 2000; France et al. 2000; Garcia et al.
674 2007; Lamontagne et al. 2000; Steedman 2000; Bertolo and Magnan 2007; Hausmann
675 and Pienitz 2009; Winkler et al. 2009). A few studies reviewed showed no effect in
676 streams (Hillman et al. 1997) or lakes (Knapp et al. 2003; Desrosiers et al. 2006), and one
677 study showed a negative effect in lakes (France et al. 1996) of harvesting on surface

678 water DOC concentrations. In a survey of lakes in the eastern boreal zone, the average
679 increase in concentration of DOC attributable to harvesting was 2 mg/L, with a range of -
680 2 to +5 mg/L (France et al. 2000).-The specific silvicultural practices applied have an
681 impact on how long surface water DOC concentrations remained elevated. For example,
682 Schelker et al. (2012) determined that while clearcutting increased DOC concentrations,
683 site preparation (intentional ground disturbance in preparation for planting) had an even
684 more profound effect on DOC concentration. However, since DOC concentrations in
685 undisturbed basins are closely linked to presence and extent of permanently or transiently
686 saturated soils (Dillon and Molot 1997; Creed et al. 2003, 2008), the presence of these
687 features may confound harvesting effects (Hillman et al. 1997).

688 **Nitrogen:** Timeber harvesting effects on nitrogen mobility and export are
689 mediated by microbial processes affecting mineralization and nitrification (Mallik and
690 Teichert 2009). Studies examining the impacts of timber harvesting on nitrogen in
691 surface waters have observed a variety of responses in boreal lakes. These include
692 increases in total nitrogen (Lamontagne et al. 2000; Steedman 2000; Garcia et al. 2007;
693 Hausmann and Pienitz 2009), organic nitrogen (Carignan et al. 2000) and inorganic
694 nitrate-nitrogen (Tremblay et al. 2009). Other studies have shown no effect on total
695 nitrogen (Garcia and Carignan 2000; Prepas et al. 2001a; Bertolo and Magnan 2007;
696 Winkler et al. 2009) or nitrate-nitrogen (Carignan et al. 2000; Lamontagne et al. 2000).
697 Variation in nitrogen response to harvesting is due to local site factors such as vegetation
698 type controlling carbon to nitrogen ratio, local environmental conditions such as
699 temperature and moisture controlling rates of mineralization, and local topography
700 affecting its aquatic vs. atmospheric fate (Holmes and Zak 1999; Lamontagne 2000).

701 **Phosphorus:** Timber harvesting effects on phosphorus export is related to soil
702 erosion and soil properties that influence biotically controlled mineralization processes or
703 abiotically controlled adsorption and desorption processes. Landscape position affects
704 these processes (Mallik and Teichert 2009). Studies examining the impacts of forestry on
705 phosphorus loading to surface waters reveal a variable response. Several studies observed
706 increased concentrations of total phosphorus in lakes (Carignan et al. 2000; Lamontagne
707 et al. 2000; Prepas et al. 2001a, Garcia et al. 2007; Hausmann and Pienitz 2009; Winkler
708 et al. 2009). Other studies observed no change in concentrations of total phosphorus
709 concentrations (Evans et al. 2000; Steedman 2000; Bertolo and Magnan 2007) or a
710 decrease in total phosphorus concentrations (France et al. 1996) or phosphate (Tremblay
711 et al. 2009) in lakes.

712 **Base cations and anions:** Base cation (calcium, magnesium, sodium, potassium)
713 and anion (chloride, sulphate) exports are controlled primarily by soils, geology, pH, and
714 organic matter associations rather than decomposition processes. Base cations in water
715 were generally observed to increase following timber harvesting. Increases were
716 observed for calcium (Carignan et al. 2000; Lamontagne et al. 2000), magnesium
717 (Tremblay et al. 2009), sodium (Lamontagne et al. 2000), and potassium (Carignan et al.
718 2000; Lamontagne et al. 2000; Tremblay et al. 2009). However, Steedman (2000)
719 observed decreases in calcium and magnesium. Trends in acid anions are less well
720 studied, and include sulphate increases (Carignan et al. 2000) or no sulphate change
721 (Garcia and Carignan 2000; Prepas et al. 2001a; Hausmann and Pienitz 2009), and
722 chloride increases (Carignan et al. 2000; Lamontagne et al. 2000; Steedman 2000) or no
723 chloride change (Prepas et al. 2001a; Hausmann and Pienitz 2009). Water pH is primarily

724 driven by balances in cations and anions and, in general, timber harvesting appeared to
725 have little effect on pH (Hillman et al. 1997; Prepas et al. 2001a; Steedman 2000; Garcia
726 et al. 2007; Hausmann and Pienitz 2009), although Tremblay et al. (2009) found a slight
727 decline in pH following timber harvesting.

728 Studies investigating the impacts of forestry on water quality parameters within
729 the boreal zone have tended to focus more on lakes than streams, with more studies from
730 the eastern boreal zone compared to the western boreal zone. Although the majority of
731 studies have shown short-term increases in many of the nutrients, it is not well known
732 what the longer-term changes are following timber harvesting and stand regrowth.
733 Nutrients may differ in their responses, with mobile nutrients (potassium, chloride,
734 sulphate, nitrate) that are released following a harvest being rapidly flushed out of the
735 basin (e.g., 50% decrease in three years), while other nutrients show little change or
736 continued increases after three years (Carignan et al. 2000). The rate of decline in nutrient
737 export will be determined by microbial demand for nutrients during decomposition,
738 vegetation demand for nutrient uptake, and water availability for solubilising and
739 mobilizing nutrients.

740

741 **3.2.3.5 Mercury.** Forestry operations have been shown to increase mercury and methyl
742 mercury output from boreal catchments (Porvari et al. 2003). Clearcutting and/or site
743 preparation significantly increases the mobility of total mercury and methyl mercury
744 accumulated in forest soil and may be an important factor for the total input of mercury to
745 boreal freshwater ecosystems (Porvari et al. 2003). Garcia et al.'s (2007) study in the
746 Haute-Mauricie region of northwest Quebec found that the concentration of methyl

747 mercury in lakes where forests were harvested were significantly higher than in reference
748 lakes, and these changes paralleled those in DOC concentrations. Similar to what was
749 observed with DOC, Sørensen et al. (2009) suggested that the majority of mercury
750 outputs are primarily from site preparation, and not from timber harvesting, and therefore
751 mitigative measures can be taken to reduce contamination of surface waters by mercury
752 caused by site preparation.

753

754 ***3.2.4 Prognosis***

755 More than a century of research into interactions between forestry and water
756 resources has led to significant improvements in forest management practices to protect
757 water resources (Ice et al. 2010). Significant improvements have been made in forest
758 management planning, beneficial management practices, and certification standards.
759 Improving road placement (White et al. 2012), restricting machine activity in wet areas
760 (e.g., OSIC 2011), and optimizing the design of riparian buffers (Creed et al. 2008) are
761 examples of changes that have led (or could lead) to improvements in water quality
762 particularly in reducing sediment and nutrient runoff.

763 The role that riparian buffers play in mitigating forest management impacts on
764 water resources has been evaluated (Buttle 2002). Fixed-width riparian buffers have
765 become the norm for streamside protection from forest management activities, due to
766 their relatively easy implementation in forest planning, yet few experiments have been
767 done to test the efficacy of riparian buffers of a particular width or explore site or
768 landscape-specific modifications (Richardson et al. 2012). A couple of examples from the
769 boreal zone suggest that the uncut “donuts” are not essential for maintaining water

770 quality. Tremblay et al. (2009) found that the proximity of the cutblock to the stream
771 network and logging within riparian buffers did not appear to affect water quality,
772 however the harvest was only of moderate intensity (50% cut). Prepas et al.'s (2001b)
773 study within the western boreal forest determined that where climatic and physiographic
774 variability produces complex hydrologic pathways, standard width riparian buffers are
775 not adequate for protecting aquatic systems. In the same region, Creed et al. (2008)
776 developed a remote sensing technique where a probability map of wet area formation was
777 calculated from a time series of remotely sensed images and related to the return period
778 of discharges from the basin. The relationship between wet area and return period was
779 proposed as an approach for estimating risk associated with harvesting in these critical
780 areas, enabling land managers to design variable width riparian buffers based on the level
781 of risk they are willing to accept with respect to increasing nutrient loading to surface
782 waters following harvesting (Creed et al. 2008). A move away from a "one size fits all"
783 approach to the design of riparian buffers towards a customization of buffer widths that
784 considers relevant landscape-specific factors based on easily derived terrain information
785 could further strengthen forest management planning.

786 Modern forest management strategies aim to emulate natural disturbance as much
787 as possible in its methods (Hunter 1993). This has implications for the protection of water
788 resources, because natural disturbances often occur in riparian buffers, and emulating
789 those disturbances will require timber harvesting closer to water than previously allowed
790 under conventional riparian buffers (Kreutzweiser et al. 2012). This emphasis on
791 emulation of natural disturbance is providing impetus for new directions in basin and
792 riparian forest management (Naylor et al. 2012), but implementation across the boreal

793 zone will need to consider and adjust for landscape contexts including regional
794 hydrologic regimes and disturbance patterns (Sibley et al. 2012). Implementation will
795 also need to consider conditions under which riparian buffer harvesting does not mimic
796 natural disturbances. Although achieving harvest patterns that are similar to fire patterns
797 near water can be successful, creating similar response in processes is more difficult. For
798 example, fires in riparian buffers result in standing dead trees that can provide shade
799 (temperature control) and pulsed inputs of large wood (organic matter input and
800 retention) to streams, whereas harvesting in riparian buffers generally does not (Moore
801 and Richardson 2012). A similar problem is encountered with emulating the impacts of
802 fire on water quality. In a series of syntheses, Buttle et al. (2000), Carignan and Steedman
803 (2000), Pinel-Alloul et al. (2002) and Nitschke (2005) showed that there were substantial
804 differences between fire and harvesting. For example, DOC, mercury, sodium, and
805 potassium responded more strongly to harvesting whereas nitrate-nitrogen, phosphorus,
806 calcium, and magnesium responded more strongly to fire, with other parameters
807 responding similarly.

808 Given the importance of forestry operations in the boreal zone, a large knowledge
809 gap remains in understanding the persistence of impacts of timber harvesting activities.
810 Most studies do not continue beyond two or three years post-harvest. While studies of
811 this duration may capture recovery of rapidly responding effects, it may not capture some
812 of the longer-term impacts or the time scales of recovery in runoff and water quality that
813 have been observed in some forest ecosystems (Carignan and Steedman 2000; Nitschke
814 2005). Furthermore, studies on the impacts of forestry on water resources have been
815 predominantly conducted outside of the boreal zone or within the eastern boreal zone and

816 the applicability of these results to the entire boreal zone is unknown. Also, studies on the
817 impacts of forestry on water quality within the boreal zone have focused primarily on
818 lakes, not streams. Although some general trends in water quality parameters were
819 observed, there was considerable variability and inconsistency in water quality
820 observations, reflecting the substantial spatial heterogeneity and temporal variability of
821 drivers of hydrologic and biogeochemical processes within the boreal zone (*sensu* Devito
822 et al. 2005), preventing a predictive understanding of the impacts of forest management
823 on water quality (Carignan and Steedman 2000; Pinel-Alloul et al. 2002; Kreutzweiser et
824 al. 2008).

825 Several studies have considered the impacts of harvesting intensity and typically
826 the impacts became more pronounced with increasing intensity of harvest, when
827 normalized by the lake surface area or volume (Carignan et al. 2000; Pinel-Alloul et al.
828 2002). However, forest management studies have typically focused on low order,
829 headwater basins, and the spatially cumulative impacts of timber harvesting at larger
830 scales has rarely been evaluated (Buttle and Metcalfe 2000), or in combination with other
831 landscape disturbances (Buttle et al. 2005). Furthermore, scaling impacts from the stand
832 level to the basin or landscape scale still remains a challenge.

833 .

834 ***3.2.5 Unique considerations for timber harvesting within peatlands***

835 ***3.2.5.1 Introduction.*** Timber harvesting in treed peatlands presents unique challenges for
836 forestry. Timber harvesting within peatlands is not a widespread practise in the Canadian
837 boreal zone as it is in the boreal zone of northern Europe (Poulin and Pellerin 2001),
838 although it occurs in certain areas, such as the Clay Belt in Ontario and Quebec and in

839 north central Alberta. Tree productivity in peatlands is generally low because of a high
840 water table, poor aeration (Campbell 1980), low substrate temperature (Liefvers and
841 Rothwell 1987), and inadequate nutrient availability (Tilton 1978). However, treed
842 peatlands have marketable black spruce (*Picea mariana*) and tamarak (*Larix laricina*)
843 that are harvested primarily as high-quality pulpwood (Locky and Bayley 2007).

844

845 **3.2.5.2 Water quantity.** In treed peatlands, water table increases after harvesting is
846 referred to as “watering-up” (Dubé et al. 1995). Watering-up is mainly due to a reduction
847 in interception and evaporation of precipitation and reduction in transpiration following
848 forest cover removal (Dubé et al. 1995) as well as flow path disturbances during forestry
849 operations (Aust et al. 1997). Immediately following timber harvest, water table levels
850 can increase by 4 to 22 cm (Berry and Jeglum 1991; Dubé et al. 1995; Roy et al. 1997,
851 2000). These elevated water tables can persist for many years. Pothier et al. (2003) did
852 not detect water table recovery five years after harvesting, and Marcotte et al. (2008)
853 found that water table levels in undrained plots were still 5 to 7 cm higher than the pre-
854 cut levels 10 years after clearcutting, which was attributed to precipitation interception by
855 vegetation being only 50% of the pre-harvest value. Watering-up creates conditions
856 conducive for paludification (i.e., peat accumulation) and is considered a threat to future
857 forest productivity (Lavoie et al. 2005). Silvicultural practices within peatlands
858 commonly include draining to improve forest productivity prior to harvest (Poulin et al.
859 2004) and after harvesting (Roy et al. 2000) to improve soil conditions for future tree
860 regrowth by increasing the depth of aeration and allowing previously shallow rooting
861 systems to expand deeper into the soil (Liefvers and Rothwell 1987; Prévost et al. 1997;

862 Silins and Rothwell 1999; Roy et al. 2000), although this practice would have negative
863 implications for atmospheric carbon exchange (Kurz et al. 2013). Watering up can alter
864 local downstream flows, although there has been no recent work to quantify the potential
865 magnitude of impact. Drainage also impacts downstream flows, with one study showing
866 drainage increased and sustained summer low flows by 25% (Prévost et al. 1999).

867

868 **3.2.5.3 Water quality.** The increase in water table following timber harvest in treed
869 peatlands creates an anoxic environment (Aust et al. 1993), root growth (Kozlowski
870 1984; Lieffers and Rothwell 1986), and nitrate assimilation (Morris 1997). Locky and
871 Bayley (2007) and Prévost et al. (1999) found that surface waters in young clearcuts had
872 significantly larger nutrient concentrations compared with controls, likely due to soil
873 warming. Peatland draining prior to timber harvest causes changes in water chemistry,
874 opposite to that of watering up, as a result of increased oxygen transport into the peat
875 (Silins and Rothwell 1998). Numerous negative effects of peatland draining have been
876 observed, including increases in suspended sediments, nutrients, specific conductivity,
877 and pH of peat soil water (Lieffers and Rothwell 1987; Lieffers 1988; Rothwell et al.
878 1993; Sheehy 1993; Paavilainen and Paivänen 1995; Prévost et al. 1997, 1999; Lavoie et
879 al. 2005). These water chemistry changes have the potential to influence downstream
880 water quality during periods when the peatlands are hydrologically connected in the
881 regional flow network (e.g., during ditching and in high flows in following weeks, to
882 months or even years [Prévost et al. 1999]).

883 **3.2.5.4 Prognosis.** Changes in demand for different wood products will determine if
884 forest management activities within treed peatlands are likely to expand or decline.

885 Extensive research that occurred during eras when demand for forest products from
886 peatlands was high and from the European experience has identified many beneficial
887 management practices to minimize impacts on water resources. Although peatland
888 draining following clearcutting or pre-commercial harvesting helps in the recovery from
889 “watering up”, it is considered a costly and remedial approach to the problem (Marcotte
890 et al. 2008). For example, drainage should be limited to the first cohort stand following
891 harvesting (Lavoie et al. 2005) because the improvement in tree growth may be relatively
892 small without additional fertilization (Jutras et al. 2002). Preventive measures, including
893 silvicultural treatments that promote regeneration and evapotranspiration, along with
894 protecting understory vegetation, should be employed to limit water table rise (Lavoie et
895 al. 2005). Partial and shelterwood cutting methods are among the best options, since
896 watering-up was found to be roughly proportional to cutting intensity (Päivänen 1980;
897 Pothier et al. 2003). Partial cutting scenarios also conserve the vertical complexity of the
898 stand that is essential to the recovery process (Marcotte et al. 2008). This requires the use
899 of appropriate equipment to minimize site disturbances while the ground is frozen (Locky
900 and Bayley 2007). Applying these practices will help in minimizing negative impacts to
901 water resources within commercially viable forested peatlands.

902

903

904 **3.3 Pulp and paper operations**

905 ***3.3.1 Introduction***

906 Pulp and paper is an important component of Canada's forest product industry
907 and in 2010 contributed to 0.7% of the GDP (NRCan 2013a) and \$9.8 B per year in
908 exports from newsprint and wood pulp (NRCan 2011). During the same year, Canada
909 ranked first in world production of newsprint (13.8%) and second in world production of
910 wood pulp (11%). Boreal forests are the dominant contributor to the pulp and paper
911 industry. As of 2011, there were 17 active and 7 closed or decommissioned pulp and
912 paper mills in the boreal portion of the commercial forest (Fig 3. in Brandt et al. 2013).

913 The pulp and paper industry is capital-intensive, characterized by complex system
914 processes to convert cellulose fiber from trees into a wide variety of traditional products
915 such as pulps, papers, and paperboards (Environment Canada 2012a) and emerging
916 products such as fibre-bioplasic composites (NRCan 2013b). Wood is reduced to fiber
917 either by cooking in chemicals or by mechanical means (Environment Canada 2012a).
918 The fibers are then mixed with water, adhering to one another as the water is removed by
919 pressure and heat. Chemical pulping (kraft and sulfite process) uses sulphur to extract
920 fiber that is exceptionally strong and is used for magazines, printing and graphics papers,
921 grocery bags, and corrugated packaging (Biermann 1993). Mechanical pulping mills
922 physically shred wood into pulp with grindstones and/or heat to produce pulp that has
923 weaker fibers and is commonly used for newspapers (Biermann 1993).

924

925

926 **3.3.2 Water quantity**

927 Large volumes of water are required for processing and cooling during pulp and
928 paper production, however most of the water is not consumed. The majority of water is
929 taken from surface waters (98%), primarily rivers (76%) and lakes (19%) (NCASI 2010).
930 About 92% of the water used by pulp and paper processes is returned directly to the
931 surface waters (river 74%, lake 13%) following treatment. 8% is evaporated during
932 manufacturing and wastewater treatment, and about 0.4% is imparted to products or solid
933 residuals (NCASI 2010). The large amounts of water used in processing and cooling
934 reappear as effluent (Pokhrel and Viraraghavan 2004). It is estimated that mills release
935 between 50 000 to 150 000 m³ of effluent per day (Chambers et al. 2001; Hewitt et al.
936 2006) and in 2007 total national effluents flows were estimated at $1\,560 \times 10^6$ m³ with an
937 additional 229×10^6 m³ in non-contact cooling water flow (NCASI 2010). Over a period
938 of 17 years (1992 and 2009), water use intensity within the industry has declined from
939 ~ 85 m³ per tonne to ~ 55 m³ per tonne (FPAC 2011). There is no evidence to suggest that
940 boreal water quantity is significantly affected by water use in the pulp and paper industry.

941

942 **3.3.3 Water quality**

943 Pulp mill effluent is a known pollutant of surface waters (Walden 1976). Pollutants are
944 generated at various stages of pulping and paper making process (Pokhrel and
945 Viraraghavan 2004). Effluents are a complex combination of waste streams produced in
946 debarking, pulp washing, bleaching, and regeneration of cooking chemicals (Pokhrel and
947 Viraraghavan 2004; Hewitt et al. 2006). Chemicals in effluents may be from the wood
948 itself or from chemicals added during the pulping and bleaching process and include

949 suspended solids, lignins, resins, fatty acids, volatile organic carbon (e.g., terpenes,
950 alcohols, phenols, methanol, acetone, chloroform, etc.), mercury and inorganic chlorine
951 and organochlorine compounds.

952 During the 1980s, pulp and paper production worldwide became an area of
953 increased environmental scrutiny by the public as dioxins and furans in effluents and
954 paper products were found to have toxic effects on aquatic organisms (McMaster et al.
955 2006*b*; Kreutzweiser et al. 2013). As a result of studies conducted internationally and in
956 Canada, new regulations came into force in 1992 to set revised limits for biochemical
957 oxygen demand (BOD; the amount of oxygen needed to decompose organic matter), total
958 suspended solids (TSS), and dioxins and furans (Environment Canada 2005*b*). To meet
959 the new regulatory limits, the industry was required to make process changes (e.g.,
960 switching from elemental chlorine to chlorine dioxide in the bleaching process) and
961 treatment changes to reduce pollution and environmental impact (McMaster et al. 2006*b*).
962 Effluents are now treated with both primary treatment to remove solids in settling basins,
963 and secondary treatment to adsorb, settle, and promote microbial breakdown of
964 biodegradable material in order to reduce BOD and levels of toxic organic compounds
965 (Pokhrel and Viraraghavan 2004; Hewitt et al. 2006).

966 Since the Pulp and Paper Effluent Regulations under the Fisheries Act came into
967 force in 1992, there have been declining trends in discharge loads of total suspended
968 solids, adsorbable organic halides, dioxins, and furans and declines in BOD (Chambers et
969 al. 2000; NCASI 2010). Dioxins have been virtually eliminated through process changes
970 in mills and decreased toxicity of effluents through biological effluent treatment,
971 reducing BOD substances by 90% and suspended solids by 70% (FPAC 2011).

972 Changes to effluent treatment to reduce contaminants have not reduced nutrient
973 load. Eutrophication that results from elevated nutrient concentrations in water remains a
974 predominant environmental issue for pulp mill effluents (Bothwell 1992; Biermann 1996,
975 Chambers et al. 2000, 2006). This is particularly a concern in northern boreal rivers
976 where they are naturally nutrient deplete (Chambers et al. 2000). The increased nutrient
977 concentrations result in increased production of phytoplankton and aquatic plants, cause
978 changes in the abundance and composition of consumers, and contribute to declines in
979 dissolved oxygen (Smith et al. 1999). Research conducted as part of the Northern River
980 Basins Study (NRBS) in the western boreal zone (see Gummer et al. 2000) found
981 elevated levels of total nitrogen and total phosphorus in the Athabasca and Wapiti rivers
982 (Chambers et al. 2006). Pulp mills contributed 22% of the P and 20% of the N load
983 discharged from the Wapiti to the Smoky river, and 6 to 16% of the P load and 4 to 10%
984 of the N load in the Athabasca River (Chambers et al. 2000). These nutrient
985 concentrations were elevated downstream of the pulp mills, particularly during the low
986 discharge periods of fall and winter (September to April; Chambers et al. 2000). These
987 elevated nutrient concentrations extended 4 to 120 km downstream of the point of
988 effluent discharge (Chambers et al. 2000).

989

990 **3.3.4 Prognosis**

991 The GDP attributed to the Pulp, Paper and Paperboard Mills industry has
992 decreased from \$8.1 billion in 2002 to \$5.7 billion in 2011. Global market trends,
993 including increased demand for paper for small printers but declines in newsprint coupled
994 with changing markets in which there is greater use of recycled fibre and fibre from

995 plantations (Whitman 2005), has affected the viability of the Canadian pulp and paper
996 industry (FPAC 2011). Changes or recovery in traditional markets and acceptance and
997 application of new technologies will impact how this sector develops into the future.

998 Declining newsprint demand and improved effluent treatments indicate that
999 effects of toxic compounds in pulp mill effluents on water quality are likely to decline in
1000 the boreal zone. Furthermore, compliance with toxic substance control regulations over
1001 the past couple of decades has resulted in improved water quality of pulp mill effluents
1002 (Environment Canada 2005*b*). However, studies from the western boreal forest indicate
1003 that nutrient additions from mill effluents continue to be a persistent challenge and will
1004 pose risks of eutrophication effects in downstream water bodies. Schindler and Lee
1005 (2010) point out that eutrophication in riverine boreal lakes is an increasing problem, and
1006 they suggest that the contribution of pulp mill effluents has to be considered along with
1007 cumulative effects of municipal sewage, agricultural, and shoreline developments. As a
1008 result of research from the Northern River Basin Study, nutrient loading to the Athabasca
1009 and Wapiti rivers has improved and pulp mills are required to develop nutrient
1010 minimization plans and investigate new methods to reduce nutrient discharge (Chambers
1011 et al. 2000). These examples from the western boreal suggest that continued
1012 improvements in effluent processing technologies are warranted to further reduce nutrient
1013 and suspended sediment loads.

1014

1015

1016 **3.4 Electric power development**

1017 ***3.4.1 Introduction***

1018 Canada is the third largest producer of hydroelectricity in the world (behind China
1019 and Brazil) accounting for approximately 10.5% of the world's total production
1020 (Observ'ER 2011). In 2010, electricity generation in Canada amounted to 588.9 terawatt
1021 hours (NRCan 2011), with hydroelectric generation representing 59.1% of total
1022 generation, and thermal sources (e.g., coal, nuclear, natural gas and petroleum products
1023 and waste) contributing 23.4% (NRCan 2011). In the same year, Canada exported \$2.2B
1024 in electricity, exclusively to the United States (NRCan 2011).

1025 In 2000, there were 713 large dams greater than 5 m in crest height across Canada
1026 in the boreal zone (Brandt et al. 2013). These dams produce a large portion of Canada's
1027 hydroelectric power, with 39% percent of Canada's hydroelectric capacity generated
1028 from rivers arising in or flowing through the Boreal Shield ecozone alone (as defined by
1029 ESWG 1996) (Urquizo et al. 2000) (Fig. 9 in Brandt et al. 2013). Two hydroelectric
1030 power projects within the boreal zone, Le Grande complex (Quebec) and Churchill Falls
1031 (Labrador), are ranked among the largest projects in the world in terms of capacity and
1032 reservoir size (International Hydropower Association 2007; Lewis 2012). Other large
1033 dams in the boreal zone can be found on the Moose River in Ontario, Nelson River in
1034 Manitoba, Saskatchewan River in Saskatchewan, and Peace River in British Columbia.

1035 Canada has diverted more water by damming rivers than any other country
1036 (Dynesius and Nilsson 1994; Ghassemi and White 2007). Almost all of the diverted water
1037 volume in Canada (97%) is used for hydroelectric power generation (Lasserre 2007).
1038 Although the distribution of dams is considerable (Fig. 7), the boreal zone remains an

1039 area with the greatest number of large, undammed, free-flowing river systems in North
1040 America (Fig. 8, Dynesius and Nilsson 1994).

1041 The impacts of hydroelectric power generation on water resources differ
1042 depending on the mode of generation, whether it is a conventional hydroelectric, thermal,
1043 or run-of-the-river installation (Baxter 1977; Renöfält et al. 2010). Conventional
1044 hydroelectric power requires construction of dams to impound water to generate
1045 electricity and is the most common type of power generation in the boreal zone. Thermal
1046 power generation uses heat energy from fossil fuels or uranium to produce steam to drive
1047 turbines. Run-of-the-river installations are small, low capacity hydroelectric power
1048 installations where natural flow of the river itself is used, requiring little to no water
1049 retention within reservoirs. In addition to generating stations and impoundments,
1050 hydroelectric power developments also bring with them thousands of kilometres of
1051 transmission lines along linear corridors that are cleared of trees and maintained by
1052 herbicides and cutting (Urquizo et al. 2000; Wells et al. 2010), creating impacts, at least
1053 initially, that are similar to forest management activities (see Section 3.2).

1054

1055 **3.4.2 Water quantity.** Of all the modes of hydroelectric power generation, conventional
1056 hydroelectric dams have the largest impact on boreal water resources, with impacts both
1057 upstream and downstream of the dam (Baxter 1977; Rosenberg et al. 1995, 1997;
1058 Environment Canada 2004; Woo and Thorne 2009; Renöfält et al. 2010). Direct and
1059 often obvious *upstream* impacts include flooding of riparian buffers, wetland and upland
1060 habitats, conversion of lotic (flowing) environments to lentic (standing water) systems,
1061 shoreline erosion with water level fluctuations (including thermokarst slumping where

1062 permafrost exists), and altered groundwater recharge patterns. The amount of water
1063 impounded upstream of a dam differs greatly from one site to another depending on the
1064 local topographic relief (Baxter 1977). Upstream flooding associated with conventional
1065 hydroelectric dams is also the main cause of peatland loss in the country (Rubec 1991).
1066 Although recent data were not available, in the early 1990s, approximately 900 000 ha of
1067 peatlands were flooded, mostly in Québec, Manitoba and Alberta (Rubec 1991; Urquizo
1068 et al. 2000).

1069 Downstream impacts of hydroelectric dams include decreased groundwater
1070 recharge, streambank erosion, and associated change in stream morphology (Baxter 1977;
1071 Rosenberg et al. 1995, 1997; Environment Canada 2004; Woo and Thorne 2009).

1072 Conventional hydroelectric dams do not diminish the magnitude of downstream water
1073 flow, unless it involves large reservoirs that lead to increased evaporation losses.
1074 However, the timing of water flow is significantly altered to meet electricity demands
1075 (Woo and Thorne 2009). Changes to downstream flow occur immediately after the dam
1076 construction is complete; flows are dramatically reduced during the time required to fill
1077 the reservoir, sometimes for up to several years (Rosenberg et al. 1995; Déry and Wood
1078 2005). Once the dam is in operation, the natural river flow regime is altered, including
1079 changes to the magnitude, frequency, duration, timing, and rate of change of flows
1080 (Magilligan and Nislow 2005). Alteration to the natural river flow regime has
1081 consequences not only for water resources, but also for the entire ecological integrity of
1082 the river system (Poff et al. 1997).

1083 Hydroelectric power developments characteristically trap high spring flows for
1084 storage in reservoirs and release higher-than-normal flows in winter when the power is

1085 needed (Woo and Thorne 2009). This results in an attenuation of the normal hydrograph
1086 in spring and enhancement in winter (Rosenberg et al. 1997). This leads to a decrease in
1087 the amplitude of annual variations in water levels, although a higher frequency of small-
1088 amplitude, short-period variations may result as the discharge is varied with electricity
1089 demand (Baxter 1977). For example, Peters and Prowse (2001) found in the Peace River
1090 that even at 1 100 km downstream of the reservoir average winter flows were 250%
1091 higher and annual peaks (1, 15 and 30 day peaks) were 35 to 39% lower.

1092 The regulation of flows that result in little to no annual spring flooding has been
1093 identified as an important impact on delta ecosystems that depend upon flooding for
1094 replenishment of water and nutrients (Woo and Thorne 2009). The Peace-Athabasca
1095 Delta, the largest boreal freshwater ecosystem in the world that has been recognized as
1096 both a Ramsar and UNESCO World Heritage Site, has been drying since impoundment
1097 of the Peace River at the Bennett Dam. Weirs were built to recreate the hydraulic
1098 damming effect of the pre-impoundment Peace River to restore flooding to the delta
1099 (Rosenberg et al. 1995), but data suggest the delta continues to dry, a condition
1100 exacerbated by climatic variability and climate change (Culp et al. 2000; Prowse and
1101 Conly 2000; Beltaos et al. 2006; Peters et al. 2006, Timoney 2006; Peters and Buttle
1102 2010) and other socioeconomic changes (Timoney and Lee 2009). However,
1103 paleolimnological evidence indicates that recently observed dryness is part of a longer
1104 trend, which began some 20-40 years prior to Peace River regulation and not out of the
1105 bounds of historical variability (Wolfe et al. 2005).

1106 Many large hydroelectric power projects divert water from the natural basin into a
1107 neighbouring basin (Ghasseni and White 2007). For example, water from the Churchill

1108 River is diverted to the Nelson River in Manitoba, and the Eastmain and Caniapiscau
1109 rivers are diverted into the LaGrande River in Quebec. Such diversions further exacerbate
1110 downstream flow patterns, by decreasing flows in the contributing river and increasing
1111 flows in the receiving river (Newbury et al. 1984; Woo and Thorne 2009). Disruptions of
1112 freshwater flows, particularly for the many boreal rivers flowing northward to the Arctic
1113 Ocean, affect the timing of the spring breakup of ice in the lower reaches of the river,
1114 hastened by the increased hydrostatic pressure of water flowing from the south to the
1115 north (Baxter 1977). Any or all of these changes in upstream and downstream flow
1116 regimes have implications for the aquatic organisms that live in those habitats (Rosenberg
1117 et al. 1997; Kreuzweiser et al. 2013).

1118 Other forms of hydroelectric power generation have less impact on water
1119 quantity. Power generation by run-of-the-river installations is very small (<10 Megawatt)
1120 in comparison to conventional hydroelectric dams (> 2 000 Megawatt) (Paish 2002; Egré
1121 and Milewski 2002). Run-of-the-river hydroelectric installations create minimal changes
1122 in flow, since they have little to no reservoir retention. The purpose of the dam in a run-
1123 of-the-river installation is to direct and control the flow of the stream and little water is
1124 impounded (Baxter 1977). However, if there is an increase in the popularity of run-of-
1125 the-river power generation, the cumulative impacts of multiple installations along a river
1126 section will need to be considered (Douglas 2007). Thermoelectric power is one of the
1127 largest water users (but not consumers) among natural resource sectors in Canada
1128 (Statistics Canada 2010). Oil and coal-fired thermal power generating plants generally
1129 use and discharge more water than natural gas-fired plants, and gas-fired plants are
1130 becoming more prevalent. The main impact on water quantity with thermal power

1131 generating stations is the loss of water through evaporation during steam generation or
1132 during cooling, if cooling towers are parts of the operation (Smith and Tirpak 1989;
1133 Statistics Canada 2010).

1134

1135 **3.4.3 Water quality.** Conventional hydroelectric power generation, through changes in
1136 upstream and downstream flow regimes, has large impacts on water quality. Upstream
1137 changes to water quality include sedimentation from bank erosion, nutrient enrichment
1138 from leaching of flooded material, and changes in thermal regime due to shorter water
1139 residence times and less mixing (Baxter 1977). Perhaps one of the most significant
1140 effects of forming upstream reservoirs is the release of methyl mercury, a strong
1141 vertebrate neurotoxin, into water (Tremblay et al. 2004; Rosenberg et al. 1997). Methyl
1142 mercury is of particular concern in the boreal zone because of the high density of organic
1143 matter deposits found in wetlands, and riparian areas, which when combined with
1144 naturally occurring or anthropogenically-deposited mercury and flooding induced anoxic
1145 conditions result in ideal conditions for mercury methylation (Kelly et al. 1997; Heyes et
1146 al. 2000). The decomposition of organic matter in flooded lands, particularly wetland
1147 soils in reservoirs, fuels the microbial methylation of inorganic mercury to methyl
1148 mercury (Hall et al. 2005). Impoundments increase the surface area of potential mercury
1149 methylation by imposing anoxia over the entire flooded area and by facilitating the
1150 exchange of nutrients and methyl mercury between the ground surface and the surface
1151 water (Heyes et al. 2000). Ecosystem-scale experiments from the Experimental Lakes
1152 Area have improved our understanding of mercury dynamics in reservoirs (e.g., Bodaly
1153 et al. 2004; St. Louis et al. 2004; Hall et al. 2005). Flooding of newly created reservoirs

1154 creates more methyl mercury than flooding of existing lakes (Rosenberg et al. 1995).
1155 Flooding of wetland areas is more detrimental than flooding of upland areas since the
1156 higher concentration of organic matter produces methyl mercury for longer periods than
1157 flooding of upland areas (St. Louis et al. 2004; Hall et al. 2005). In a flooding experiment
1158 in boreal upland forests, Hall et al. (2005) found that within five weeks of flooding,
1159 methyl mercury concentrations in the reservoir outflows exceeded those in reservoir
1160 inflows. The initial pulse of methyl mercury production in all experimental reservoirs
1161 lasted for two years, after which net demethylation began to reduce the pools of methyl
1162 mercury in the reservoirs, but not back to levels found prior to flooding (Hall et al. 2005).

1163 Many of the impacts caused by conventional hydroelectric dams on the streams
1164 below them are the reverse of those produced on the reservoirs above them (Baxter
1165 1977). Attenuated spring flows reduce organic matter and nutrient loading to downstream
1166 waters during biologically active periods, ultimately influencing coastal productivity
1167 (Rosenberg et al. 1997). Other downstream impacts include altered water temperature
1168 regimes (cooler in the summer and warmer in the winter) and increased turbidity due to
1169 sediment inputs from erosion (Baxter 1977).

1170 As for other types of hydroelectric power generation, run-of-the-river has few
1171 impacts on water quality, allowing sediment to pass through relatively unimpeded
1172 (Kondolf 1997). Thermal power generating stations have impacts on water quality, with
1173 the primary concerns related to elevated temperatures of discharge waters and potential
1174 pollution from corrosion-control products used in cooling waters (Donahue et al. 2006;
1175 NRTEE 2010). Madden et al. (2013) observed discharge water temperatures could be 9.5
1176 to 10°C higher than intake water during summer. Although no boreal specific values

1177 could be found, this effect could be larger in northern ecosystems where intake
1178 temperatures would be lower, particularly in the winter. These thermal effects may be
1179 quite localized, but may impact lake stratification and other chemical and biological
1180 processes (Baxter 1977; Environment Canada 2004). In the United States, thermal
1181 regulations have caused a shift in the design of new cooling systems, from once-through
1182 cooling systems that discharge heat back into water sources, to cooling systems that
1183 evaporate water with towers and ponds that alleviate the thermal pollution but result in
1184 increased water consumption due to evaporation (Smith and Tirpak 1989).

1185 Thermal power generation, particularly with coal-fired generating stations, also
1186 result in atmospheric emissions of pollutants that may be deposited on local water bodies,
1187 or enter the airshed to be subsequently deposited on downwind water bodies. The impact
1188 of the pollutant emission and deposition is similar to that described for other resource
1189 industries (see section 3.7). Donahue et al. (2006) found using paleolimnological
1190 reconstructions within a 35 km radius of coal-fired power plants at Wabamun Lake in
1191 central Alberta that sediment concentrations of mercury, copper, lead, arsenic and
1192 selenium had increased by 1.2 to 4 times since the four power plants had been built in
1193 1950. Pollution abatement technologies will become increasingly important, particularly
1194 if increased thermal power generating capacity in Alberta is met by expansion of the
1195 coal-burning industry (Donahue et al. 2006).

1196

1197 **3.4.4 Prognosis.** Electric power generation has differential impacts across the boreal
1198 zone. With high topographic relief and wetter climates, the eastern and cordilleran parts
1199 of the boreal zone are ideal for hydroelectric power generation, while flatter and drier

1200 areas of west-central part of the boreal zone are dependent primarily on thermal power
1201 generation. Thus, the specific impacts of these different technologies will be manifested
1202 in the areas in which they occur.

1203 There are currently approximately 55 thermal power generation stations across the
1204 boreal zone, mostly in Alberta and the Northwest Territories. An increasing demand for
1205 cleaner power will likely lead to a decline in fossil fuel based thermal generating stations
1206 creating an energy gap that will likely to be filled by hydroelectric power generation
1207 based on global trends (Kumar et al. 2011). Northern rivers in the boreal zone hold the
1208 most remaining potential for large-scale hydroelectric power development in Canada
1209 (Fig. 11; Environment Canada 2004), and technology exists to more than double the
1210 existing hydroelectric power capacity in Canada (Canadian Hydropower Association
1211 2008). From 2011 to 2030, there is a potential for 158 hydropower projects totalling 29
1212 000 Megawatts of new capacity in Canada (HEC Montréal 2011). A large number of new
1213 conventional hydroelectric power developments have already been identified across the
1214 boreal zone (Environment Canada 2004; Fortin and Collu 2008). Although run-of-the-
1215 river installations are more environmentally friendly, their low capacity output is likely to
1216 limit their use to meeting specialized local power demands. However, efforts are being
1217 made to improve technologies related to small hydropower (1-50 Megawatt) capacity
1218 (e.g., NRCan 2009). Increased use of other alternative energies for electricity generation,
1219 particularly biomass and geothermal, may bring new impacts on surface and ground
1220 water resources.

1221 Many of the impacts of conventional hydroelectric power generation can be
1222 mitigated through avoidance of diversions, careful site planning, and use of power

1223 storage technologies. For example, upstream impacts can be minimized by reducing the
1224 size of reservoirs. This can be achieved by careful site selection to ensure minimal land is
1225 flooded (Baxter 1977; Mailman et al. 2006). Downstream impacts can be minimized by
1226 allowing for natural flow regimes. Small dams can be converted to run-of-the-river
1227 installations and larger dams can use pump storage or other means to store potential or
1228 electrical energy (Baxter 1977; Egré and Milewski 2002; Mailman et al. 2006; Renöfält
1229 et al. 2010). Most of these environmental mitigation measures require flow modifications
1230 that could reduce power production. An important challenge for river management will
1231 be to identify situations where measures involving relatively small production losses can
1232 have major ecological advantages (Renöfält et al. 2010).

1233 Mitigation strategies have been proposed that may reduce the impacts of methyl
1234 mercury although not all strategies have been adequately developed or tested. Some of
1235 these mitigation strategies are theoretical and not feasible at large scales and often shift
1236 the mercury to a different fate, whether it be volatilized back to the atmosphere *via*
1237 burning or transported downstream *via* regulated water flows (Mailman et al. 2006). The
1238 best mitigation strategy is to minimize the area flooded (i.e., deepening channels or
1239 avoiding flat topography) and ensure that wetland areas with high organic matter are
1240 avoided (Rosenberg et al. 1995, 1997; Bodaly et al. 2004). Other mitigation strategies
1241 include pre-flooding activities to reduce methyl mercury inputs, including removal of
1242 standing trees or burning before flooding (e.g., Rosenberg et al. 1995; Mailman and
1243 Bodaly 2005, 2006), and post -flooding activities that involve flow regulation to release
1244 methyl mercury during periods of peak flow production, intensive fishing to remove
1245 methyl mercury bioaccumulated in fish (e.g., Verta 1990; Surette et al. 2003), adding

1246 selenium to reduce methyl mercury bioaccumulation (e.g., Turner and Rudd 1983),
1247 adding lime to acidic systems to reduce methyl mercury bioaccumulation (e.g., Wiener et
1248 al. 1990), adding phosphorus to cause growth dilution by stimulating productivity (e.g.,
1249 Kidd et al. 1999), demethylation of methyl mercury by ultraviolet light (e.g., Sellers et al.
1250 1996), capping and dredging bottom sediments (e.g., Hecky et al. 1987), and aerating
1251 anoxic bottom sediments and waters (e.g., Matilainen et al. 1991). The optimal mitigation
1252 strategy will depend on the system and situation. Mitigation strategies reviewed by
1253 Kreutzweiser et al. (2013) for reducing impacts of hydroelectric power generation on
1254 aquatic biodiversity are equally relevant to reducing effects on water quantity and quality.

1255 With many hydroelectric dams constructed 40 to 50 years ago, issues related to
1256 maintaining, upgrading, and decommissioning aging hydroelectric installations are
1257 increasing (Environment Canada 2004). Long-term planning is necessary to identify the
1258 needs for these facilities. Facility improvements can be viewed as opportunities to
1259 maximize hydroelectric power generation through implementation of technological
1260 advances and to minimize environmental impacts. Facility decommissioning can be
1261 viewed as opportunities to avoid further adverse effects on water resources both in
1262 reservoirs and downstream areas (Hart et al. 2002; Stanley and Doyle 2003; Environment
1263 Canada 2004). Bednarek (2001) reviews the ecological impacts of dam removal,
1264 demonstrating that restoration of unregulated flow regimes increases biotic diversity by
1265 returning riverine conditions and sediment transport to formerly impounded areas;
1266 however, there are increased sediment loads (and the contaminants they contain) over the
1267 short-term. There is a scarcity of empirical knowledge on environmental responses to

1268 dam removal (Hart et al. 2002), particularly how dam characteristics may influence their
1269 decommissioned effects (Poff and Hart 2002).

1270 Comprehensive cumulative effects assessments of hydroelectric mega-projects is
1271 challenging because downstream areas often are out of the jurisdiction of the agency
1272 responsible for the upstream water development project and for studying its potential
1273 impacts (i.e., the problem is created in one province, but the impacts are felt in another)
1274 (Rosenberg et al. 1995). These assessments are highly complex, expensive, and require
1275 long-term databases from before and after the project, which are seldom available.
1276 Confounding these assessments is a warming climate. Winter flows will increase, spring
1277 freshet dates will advance, but peak flows will decline, as will summer flows due to
1278 enhanced evaporation (Woo et al. 2008; Price et al. 2013). It is predicted that such
1279 changes could reduce hydroelectric capacity by 15% by 2050 (NRTEE 2009).

1280

1281 **3.5 Mining**

1282 ***3.5.1 Introduction***

1283 Canada is one of the leading mining nations in the world, producing more than 60
1284 minerals and metals. In 2010, total mineral products had domestic exports valued at
1285 \$81.4B per year (21.8% of total exports) and mining and mineral-processing industries
1286 contributed another \$34.7B per year to the Canadian economy, representing 2.8% of the
1287 national GDP. Total direct employment in mining and mineral-processing industries was
1288 308 000 or 2.1% of Canada's total employment (NRCan 2011).

1289 Eighty percent of the mining within Canada occurs in the boreal zone, where there
1290 are about 100 active mines, 6 smelters, and 9 coal mines (Fig. 2E, and Fig 6. in Brandt et

1291 al. 2013). In addition, there are over 10 000 mining and exploration sites across Canada
1292 in varying stages of decommissioning and remediation (Tremblay and Hogan 2006),
1293 including more than 1 300 abandoned mines in the boreal zone (Brandt et al. 2013). The
1294 types of materials mined and how they are mined differ across the boreal zone, including
1295 ferrous metals, precious metals, base metals, precious gems, and coal that are extracted
1296 from deep below the ground or at the surface (e.g., open pit or placer mining from rivers).
1297 Mining is not environmentally benign; impacts to water resources can occur throughout
1298 the mine cycle, from exploration, extraction, and smelting to waste management,
1299 including containment and remediation (Ptacek et al. 2004) and even in the creation of
1300 ice roads for winter access in remote mines (Cott et al. 2008). Most impacts of mining on
1301 water resources arise from discharge or seepage of mine effluents and acid mine drainage
1302 into receiving waters, and from atmospheric deposition of smelter-produced acidic and
1303 metal pollution (Gunn 1995; Urquizo et al. 2000). The resulting acidification and
1304 contamination of receiving waters have implications for the aquatic biodiversity in those
1305 systems (Kreutzweiser et al. 2013).

1306

1307 ***3.5.2 Water quantity***

1308 Various steps in the mine lifecycle involve water, potentially leading to
1309 disruptions in surface and groundwater quantity. The impacts are variable depending on
1310 type of mining (e.g., surface vs. subsurface), surface overburden (upland vs. wetland
1311 soils), underlying geology, mining methods, and the material being mined. During the
1312 exploration phase, water quantity is primarily affected by seismic lines, roads,
1313 excavations, field camps, and other exploration activities, all or any of which can disrupt

1314 surface water flow pathways (see section 3.1). Sampling and drilling activities may
1315 disrupt groundwater flow pathways if they come in contact with subsurface aquifers, and
1316 can release underground water sources to the surface if done improperly. These potential
1317 impacts of mining on water quantity in the boreal zone are described in at least two non-
1318 published reports (BCMOE 2008; NMC 2008); our literature search did not find recent
1319 peer-reviewed publications or reviews of potential mining impacts on water quantity in
1320 boreal basins. Younger et al. (2002) provides an extensive review of the impacts of
1321 mining on water resources, based largely on studies in non-boreal regions.

1322 During mine operation, water quantity impacts vary with the mining method.
1323 Surface mining affects surface water flow pathways. Rivers and lakes may need to be
1324 diverted or drained to access the underlying materials. In the instances where extensive,
1325 shallow mineral deposits are located just below the ground surface, large-scale removal
1326 of overburden is required to access the underlying resource via open pit mining (Ptacek et
1327 al. 2004). When these surface mines are in wetland-rich areas, a large amount of wetland
1328 area will be drained and peat will be removed. Diamond mining at Victor Mine in
1329 Ontario’s Hudson Bay lowlands is a current example of near-surface mining in wetland-
1330 rich areas that can impact the surface water and groundwater levels (CEAA 2004).
1331 Diamond mines are usually focused on kimberlite pipes that are often surrounded or
1332 overlain by water that must be diverted or displaced for mining operations (Foote and
1333 Krogman 2006).

1334 The removal of surface mined substrate and overburden (i.e., soil or peat)
1335 decreases the hydraulic gradient. This causes a rise in the water table, and wells must be
1336 drilled to pump excess water. This “dewatering” increases vertical recharge of the

1337 surrounding land, causing groundwater levels to decline in the vicinity of the excavation,
1338 and potentially reducing surface water levels of local watercourses and wetlands to levels
1339 that could be environmentally damaging (Atkinson et al. 2010; CEAA 2011). The extent
1340 of water table decline depends on hydraulic properties of subsurface materials and
1341 pumping rates. For example, in wet areas, water losses in wetlands from mine dewatering
1342 can be significant and irreplaceable for the duration of the mine (Whittington and Price
1343 2012).

1344 Underground mining can also divert water sources and pathways. Underground
1345 openings at a gold mine in the Northwest Territories have caused changes to groundwater
1346 flow and to the circulation of surface flow to depth (Douglas et al. 2000). Placer mining
1347 removes deposits associated with stream beds, so this type of mining activity involves
1348 water damming and diversion that can influence natural drainage patterns and increase
1349 erosion and sedimentation rates (Pentz and Kostaschuk 1999). Placer mining accounts for
1350 about 5% of Canada's annual gold production, and Canadian placer gold mining is almost
1351 exclusively conducted in mountain regions of the Yukon Territory and British Columbia
1352 (McCracken et al. 2007).

1353 Water quantity may also be affected by mining at the processing stage. During
1354 metal and mineral processing, water is extracted or diverted from local watercourses for
1355 processing, cooling, and diluting or for treating mined materials. During the remediation
1356 phases, replacement of overburden and creation of new features (e.g., diversion ditches)
1357 usually alters surface and subsurface drainage patterns (Ptacek et al. 2004; CEAA 2011).

1358

1359 **3.5.3 Water quality**

1360 Most published literature on the impacts of mining in the boreal zone focused on
1361 water quality. The degree of impact of mining on water quality depends on the nature of
1362 the ore, its host rock, and the way that it is mined and processed (Wireman 2001). Mine
1363 related effluents, seepages, and emissions during mining and smelting cause the majority
1364 of water quality impacts, but impacts can also arise from leaching of waste rock and
1365 tailings well after mine closures (Leblanc et al. 2000; Urquizo et al. 2000; Ptacek et al.
1366 2004). These waste materials and other exposed rock surfaces from mining often contain
1367 sulphide minerals that, when oxidized, produce acid drainage that elevates concentrations
1368 of metals in receiving waters (Blowes et al. 2003). Depending on the ore and its
1369 byproducts, these metal concentrations in water can include heavy metals as well as
1370 cyanide, arsenic, or radioactive compounds. In the absence of effective remedial actions,
1371 these waste impoundments have the potential to oxidize and produce acid mine drainage
1372 for centuries (Blowes and Jambor 1990; Moncur et al. 2005; Kalin et al. 2006).

1373 Studies over the last two decades in or on the edge of Canada's boreal zone
1374 showed that surface waters near historic and active mining and smelting operations often
1375 continue to contain elevated levels of various metals. Some of these elevated
1376 concentrations are from continuing discharges or leaching of effluents and tailings, and
1377 some reflect continuing mobilization of metals from saturated soils near smelters. Using
1378 an assessment approach based on region-specific objectives, de Rosemond et al. (2009)
1379 analyzed water quality from sampling sites downstream of mining effluent outputs and
1380 found that average water quality indices were significantly or measurably lower than at
1381 corresponding reference sites. Many (about 80%) of these mining sites (n = 71) were in

1382 or on the edge of the boreal zone, and the water quality indices were largely reduced by
1383 elevated concentrations of metals and nutrients. Some approached or exceeded values
1384 considered harmful to aquatic life.

1385 Wang and Mulligan (2006) reviewed the environmental occurrence of arsenic in
1386 Canada, and reported that tailings from historic and recent gold mine operations in
1387 several regions continue to leach arsenic to surface waters resulting in concentrations
1388 more than 100 times acceptable levels, posing risk of harm to aquatic organisms. Historic
1389 silver mining in northern Ontario has also left large volumes of arsenic-bearing mine
1390 wastes that leach to nearby watercourses resulting in aqueous arsenic concentrations at
1391 least 10 times higher than Canadian drinking water standards, although the spatial extent
1392 of this exceedance appeared to be limited to within 1 km of the tailings outflow (Kwong
1393 et al. 2007).

1394 Selenium concentrations in lake water at 5 to 10 times higher than reference
1395 levels were found up to 10 km downstream from uranium mining operations in northern
1396 Saskatchewan (Muscatello et al. 2008). In a separate study from the same region,
1397 Muscatello and Janz (2009) reported selenium concentrations in water of a lake at the
1398 outflow of an effluent management system of approximately 20 times above reference
1399 levels. Wiramanaden et al. (2010) found that although selenium concentrations in lakes
1400 downstream of uranium mining in their study area were fairly low some distance from the
1401 discharge point, they were above reference levels and accumulated in fish resulting in
1402 adverse effects.

1403 Cyanide-containing tailings from an open pit gold mine near the boreal zone in
1404 eastern Canada that closed in 1992 continue to leach mercury and other metals (e.g., Cu,

1405 Zn, Pb) from a contaminated groundwater plume into a headwater stream, resulting in
1406 highly elevated mercury concentrations (Maprani et al. 2005; Al et al. 2006). An earlier
1407 study at the same site near the boreal zone also reported elevated aqueous gold
1408 concentrations in stream water originating from the weathering of tailings piles
1409 (Leybourne et al. 2000). Singh and Hendry (2013) concluded that there continues to be
1410 long-term risk of nickel and uranium leaching to surface and groundwater from waste-
1411 rock piles at an active uranium mine in northern Saskatchewan.

1412 Moncur et al. (2006) found that 50 years after a zinc-copper mine in Manitoba
1413 closed, metal and sulphate concentrations in a lake below the outflow from a tailings
1414 impoundment were still elevated, with peak concentrations up to 8 500 mg L⁻¹ of iron,
1415 100 mg L⁻¹ of aluminum, and 20 000 mg L⁻¹ of sulphate, although concentrations varied
1416 spatially throughout the lake. A small lake in northwestern Ontario received tailings
1417 water with high concentrations of metals since the 1960s, but showed significant declines
1418 in metal loading, especially copper, zinc, and nickel, to the lake after improved water
1419 treatment and management practices by the mid 1990s. However, arsenic concentrations
1420 increased through that period to the end of the study, apparently the result of
1421 remobilization of arsenic stored in sediments (Martin and Pedersen 2002). Similarly, a
1422 lake near Flin Flon, Manitoba that received metal loadings for more than 50 years from
1423 metal mining and processing in the region showed increasing levels of zinc in recent
1424 years despite improved water treatment and reduced discharges of zinc (Bhavsar et al.
1425 2004). They attributed the increased zinc concentrations to mobilization of historically
1426 deposited zinc in the lake drainage basin. Further evidence of the remobilization of stored
1427 metals in basins was shown by Outridge et al. (2011). They measured metal

1428 concentrations in cores from peat and lake sediments (reflecting trends in water
1429 chemistry) in the Flin Flon region and found that although mercury and other metal
1430 concentrations in peat declined over time in accordance with reduced smelter emissions,
1431 mercury and zinc increased in lake sediments. They attributed these increases to the
1432 mobilization of historically deposited metals in surrounding basins.

1433 Nickel mining and smelting in the Sudbury region on the edge of the boreal zone
1434 in Ontario have left a legacy of metal-contaminated soils in regional basins that continue
1435 to export metals to receiving waters. Although it is clear from earlier studies (e.g., Keller
1436 and Pitblado 1986) that emission controls and other remedial actions have reduced metal
1437 loading to lakes within the region, recent studies continue to find elevated metal
1438 concentrations in surface waters. Rajotte and Couture (2002) and Eastwood and Couture
1439 (2002) reported that several lakes in the Sudbury region contained elevated levels of
1440 aluminum, cadmium, copper, nickel, and zinc, often exceeding provincial water quality
1441 objectives by five times or more and with negative implications for fish condition.
1442 Hruska and Dubé (2004) found that effluents being discharged to stream water contained
1443 several metals or their byproducts at concentrations 10 to 22 times higher than reference
1444 levels, with selenium 64 times higher than reference. Ponton and Hare (2009) reported
1445 that nickel concentrations in lakes near mining and smelting operations were up to 200
1446 times higher than in reference lakes. Several metals were found at elevated
1447 concentrations in stream water receiving metal mining effluents, and these concentrations
1448 led to elevated metal body burden in fish (Weber et al. 2008). Further information on
1449 trends in metal contamination of lakes in the Sudbury region is found in Keller et al.
1450 (2007) and Keller (2009).

1451 Coal mining also produces waste rock that can oxidize and become sources of
1452 acidity, metals, and salinity that leach to receiving waters (Zielinski et al. 2001). Open pit
1453 coal mines, by virtue of their associated overburden removal that can extend for tens of
1454 square kilometers (e.g., Strong 2000), can disrupt hydrologic flowpaths and affect
1455 downstream water quality. Although coal mining is active in the boreal zone, particularly
1456 in western regions, and those mining installations are under national regulatory
1457 assessment (e.g., EUB-CEAA 2000), our search of the published literature did not find
1458 any studies reporting the impacts of coal mining on water quantity or quality in the boreal
1459 zone. Studies from non-boreal regions show that coal mining can adversely affect surface
1460 water quality (e.g., Zielinski et al. 2001; MacCausland and McTammany 2007) and
1461 similar risks would be expected from coal mining activities in the boreal zone. The lack
1462 of published literature to assess these risks in the boreal zone remains a critical
1463 information gap for coal mining assessment.

1464

1465 ***3.5.4 Prognosis***

1466 Several aspects of mineral and metal mining activities have the potential to affect
1467 water quantity (i.e., water levels of natural water bodies) through water diversion,
1468 overburden removal and replacement, and intentional dewatering. However, it appears
1469 that compliance with current mining regulations and implementation of mitigation
1470 measures can reduce these impacts on water quantity. In at least one example, mitigation
1471 measures proposed for a major surface gold mine were considered effective at reducing
1472 the risks of adverse and long-lasting changes in water levels of local watercourses
1473 (CEAA 2011). Mitigation measures for impacts on water quantity include the

1474 establishment of water flow monitoring stations, water flow supplementation by pumping
1475 when necessary to maintain base flows within 15% of seasonal norms, active open pit
1476 flooding to increase aquifer recovery after mine closure, proper drainage channel
1477 installation to manage and minimize water diversion, and managing site infrastructure to
1478 minimize the overall footprint on water flow paths and wetlands (BCMOE 2008; CEAA
1479 2011). The mining sector in general is not a significant consumer of water, and threats to
1480 water quantity from mining of metals and minerals across the boreal zone do not appear
1481 to be a constraint on future mining development (Ptacek et al. 2004; NRTEE 2010).
1482 Potential impacts to water quantity from current mining activities and the effectiveness of
1483 mitigation measures to minimize these impacts could not be assessed from the published
1484 literature. Peer-reviewed scientific studies on these issues around water quantity are
1485 lacking and represent an information gap.

1486 Water quality is often impacted near mining and smelting operations in the boreal
1487 zone. Historic mining areas, tailings ponds, and older and abandoned mines and sites are
1488 more likely to contribute to elevated metal concentrations in receiving waters than newer
1489 installations (Urquizo et al. 2000; Ptacek et al. 2004). Stakeholder working groups that
1490 have been formed to assess and report on the effects of metal mining on aquatic
1491 ecosystems have concluded that recent metal mining effluent regulations and best
1492 practices have improved water quality at many sites, but that elevated concentrations of
1493 metals and other dissolved solids continue to be detected and are linked to adverse effects
1494 on aquatic organisms in many areas (AQUAMIN 1996; Ptacek et al. 2004; Environment
1495 Canada 2012b; Kreutzweiser et al. 2013). The related issues of atmospheric deposition of
1496 acidic pollutants from processing emissions and their effects on water quality are ongoing

1497 (see section 3.8). Emissions of acidifying pollutants have declined by about 50%
1498 nationally, and by much more in some areas (CCME 2011), but atmospheric deposition
1499 of acidic pollutants still exceeds critical loads, and acidifying emissions are still
1500 increasing in other areas (Environment Canada 2005a; FPTC 2010).

1501 The impacts on water quality can be mitigated to some extent by careful mine
1502 operations, effluent and emission controls, and improved tailings treatment and
1503 management. Canada's National Orphaned/Abandoned Mines Initiative is supported by
1504 the Mining Association of Canada and other non-governmental and governmental bodies
1505 and is committed to the remediation of abandoned mines (www.abandoned-mines.org).
1506 However, with more than 10 000 exploration and mining sites across Canada that require
1507 varying degrees of remediation (Tremblay and Hogan 2006), including at least 1 300
1508 abandoned mines in the boreal zone (Brandt et al. 2013; see also CESD 2002; CDL 2005;
1509 Cowan et al. 2010), the problem of persistent metal contamination to waterbodies from
1510 older mining sites will continue for some time. Several long-term studies have
1511 demonstrated that water quality conditions in the boreal zone can measurably improve
1512 with enhanced emission and effluent treatments and that these improved conditions can
1513 promote biological recovery (Keller and Yan 1991; Gunn et al. 1995; Keller et al. 1998,
1514 2004, 2007). However, the recovery pathways are long and complex (Jeziorski et al.
1515 2013; Luek et al. 2013; Szkokan-Emilson et al. 2013), often confounded by the
1516 remobilization of historically deposited or stored metal concentrations in soils and
1517 sediments as described above.

1518 Recent studies have demonstrated that improved treatment of mine tailings and
1519 acid mine drainage can minimize the contamination of downstream water bodies. Sherriff

1520 et al. (2011) showed that despite high oxidation rates and metal concentrations in tailings
1521 ponds of a copper-zinc mine in Manitoba (closed in 2002), metal concentrations in a
1522 downstream lake receiving tailings outflow water were stable indicating that the tailings
1523 management activities at that site were effective to date. Kachhwal et al. (2011) also
1524 demonstrated that current water cover technologies and other tailings management efforts
1525 at a nickel-copper mining site in northwestern Ontario (closed in 1998) were effective at
1526 preventing downstream export of metals in the final effluent. Water cover technologies in
1527 rehabilitation efforts at uranium mining sites in northcentral Ontario (closed as recently
1528 as the mid 1990s) have successfully reduced acid generation such that only limited
1529 effluent treatment is required for pH control and radium removal (Davé 2011). A nickel-
1530 copper mining facility currently operating in Quebec has employed new technologies in
1531 waste management and water treatment, and reports that water discharge to downstream
1532 environments have metal and byproduct concentrations well below water quality limits
1533 (Kratochvil 2010). The use of natural and constructed wetlands for bioremediation of
1534 mining wastes is a reclamation technology that would appear to be well suited for aquatic
1535 systems in the boreal zone given the prevalence of natural wetlands (Sobolewski 1996;
1536 Zhang et al. 2010). Other biological remediation and ecological engineering technologies
1537 are also being explored and developed to prevent the formation or migration of mining
1538 waste leachates (Johnson and Hallberg 2005). Kalin et al. (2006) demonstrated that
1539 biological remediation efforts in a lake in northwestern Ontario affected by acid mine
1540 drainage were successful at preventing changes to water quality in the next lake
1541 downstream from the affected lake. Regardless of these demonstrated successes,
1542 biological remediation and other ecological engineering technologies have not been

1543 widely adopted by the Canadian mining industry (Kalin and Wheeler 2011), possibly
1544 because these technologies may be limited by cold temperatures in the boreal zone, or by
1545 the relatively low volumes of water that can be reasonably treated in this fashion.

1546 Recent national mining impact assessments continue to report degraded water
1547 quality near mining sites (Ptacek et al. 2004; Environment Canada 2012b), indicating that
1548 mitigation measures for water quality are not always completely effective. Degraded
1549 water quality from mining activities can pose risks of harm to aquatic biodiversity, and
1550 national reporting regulations under the Environmental Effects Monitoring Program are
1551 intended to track these risks and impacts (Environment Canada 2012b; Kreutzweiser et
1552 al. 2013). To the extent that metal mining developments and their associated tailings
1553 ponds, waste rock, effluents, and emissions are likely to increase, the potential for
1554 contamination of regional water bodies could also increase and would be incremental to
1555 the legacy effects of many older sites. From the published studies, it was not always clear
1556 to what spatial extent the contamination of water bodies extended beyond mining or
1557 processing sites. Most studies reported impacts within a few kilometers of the mining
1558 sites and many showed declining metal concentrations or water quality issues with
1559 distance from the mining activities. Given the extensive distribution of metal mining
1560 across the boreal zone, the persistent nature of metal contamination, and the fact that
1561 mining and processing often occur in or near water-abundant areas, there is potential for
1562 significant cumulative impacts at regional scales. This potential for cumulative impacts
1563 has not been rigorously assessed in the published literature and remains an information
1564 gap. There are indications that the mining sector is responding to the challenges of
1565 environmental stewardship by implementing various beneficial management practices

1566 and mining reforms across Canada (MAC 2010; PDAC 2012). This includes new
1567 reporting requirements for improved tailings management, environmental performance,
1568 and biodiversity conservation management (MAC 2010).

1569

1570 **3.6 Conventional oil and gas and oil sands development**

1571 ***3.6.1 Introduction***

1572 Fossil fuels account for the greatest share of Canada's production of primary
1573 energy, dominated by crude oil (41.4%), natural gas (36.5%), and coal (9.2%)
1574 (NRCan2011). The oil sands is an important component of our crude oil, representing
1575 97% of the crude reserves and a little over half of crude production in 2010 (NRCan
1576 2011), with the majority from surface extraction (King and Yetter 2011). Oil and gas
1577 developments are located primarily in the sedimentary regions of the western boreal zone
1578 (Fig. 9). It has been estimated that the oil and gas well sites, pipelines, and seismic lines
1579 (Fig 2F) have an industrial footprint of 46 million ha (Timoney and Lee 2001). Brandt et
1580 al. (2013) calculated from the Seismic Data Listing Service that there were about 441 000
1581 km of pipelines and 1.7 million km of seismic lines. There are more than 220 000 active
1582 and inactive well sites drilled by the oil and gas industry in boreal Canada and this
1583 number is increasing at the rate of approximately 10 000 new wells per year (Wells et al.
1584 2010; Brandt et al. 2013). Pasher et al. (2013) estimate that well sites and oil and gas
1585 infrastructure cover approximately 114 377 ha. Oil and gas developments can be
1586 characterized as being intensive (surface extraction of oil sands bitumen) or extensive
1587 (individual well sites). Similar to mining, water resources are impacted by oil and gas
1588 development at all stages of the development cycle, including exploration (seismic lines),

1589 drilling (roads and well pads, water consumption), and production (water removals and
1590 contamination).

1591 Impacts on water resources differ depending on the mode of extraction. In older
1592 conventional oil fields, water is pumped down the well to stimulate production. In the oil
1593 sands, *in-situ* extraction uses steam to liquefy the viscous bitumen so it can be pumped to
1594 the surface (Gosselin et al. 2010; CAPP 2013). Surface extraction of oil sands occurs
1595 where bitumen reserves are within 70 m of the surface of the ground. After the removal
1596 of soils and overburden, bituminous sands are removed with conventional mining
1597 techniques (e.g., truck and shovel) (King and Yetter 2011; Mikula 2012). Bitumen is
1598 extracted by mixing with hot water and mechanical agitation. Typically, the bitumen is
1599 about 10% by weight of the raw ore and 12 volumes of water, sand, silt, and clay tailings
1600 are created for each volume of bitumen produced (Kasperski and Mikula 2011). The
1601 resulting tailings are pumped to tailings ponds where the sand settles quickly leaving a
1602 suspension of fine particulates and dissolved organics, minerals, and salts (Kasperski and
1603 Mikula 2011; CAPP 2013).

1604 Parts of the boreal zone in Alberta and British Columbia have natural gas
1605 contained within shale that can only be accessed by a technique known as hydraulic
1606 fracturing (fracking) (National Energy Board 2009). Pressurized hydraulic fracturing
1607 fluid, which contains large volumes of water, mixed with sand and chemical additives, is
1608 pumped into the wellhead at high pressure. This creates cracks in the rock beds, allowing
1609 for increased extraction and recovery of gas (National Energy Board 2009; CAPP 2013;
1610 CSUG undated). Although currently shale gas extraction is limited to only a few places in
1611 the western boreal, this component of the oil and gas industry is likely to increase in the

1612 future. Although environmental impacts of shale gas extraction are known (Council of
1613 Canadian Academies 2014) , boreal-specific studies on the impacts on surface and
1614 groundwater resources are presently lacking.

1615

1616 ***3.6.2 Water quantity***

1617 Water is a critical component of oil and gas production in Canada. For example,
1618 about 75% of Alberta’s oil production is water assisted (CAPP 2010). In 2010, 9% of all
1619 water allocated within the province of Alberta was licensed for use by the oil and gas
1620 industry (CAPP 2012). The source of water for oil and gas extraction may be from
1621 surface or groundwater sources (CAPP 2010, 2012; Government of Alberta 2013b).

1622 When surface water sources are used, the main concern for water quantity is the
1623 reduction of water supplies to downstream users and ecosystems (DFO 2010). Given that
1624 river flows are variable, removals have more impact on in-stream processes when water
1625 flow rates are low (DFO 2010). Groundwater can be sourced from shallow or deep
1626 sources or from fresh or saline sources (Government of Alberta 2012b, CEMA 2012).

1627 The use of groundwater for oil and gas extraction has the potential to deplete local
1628 aquifers that are needed for other important purposes (e.g., drinking water). These
1629 groundwater depletions may also alter pressure within the aquifer which influences local
1630 or regional groundwater flow patterns and recharge of surface waters and wetlands
1631 (Manitoba Environment 1993; Griffiths et al. 2006; Government of Alberta 2012b).

1632

1633 ***3.6.2.1 Oils sands example***

1634 All aspects of the oil sands developments, including surface mining, bitumen upgrading,
1635 and *in-situ* extraction, are dependent upon water (Gosselin et al. 2010; CEMA 2012). It is
1636 estimated that 523 million m³ of water is used per year in the oil sands extraction process
1637 (Griffiths et al. 2006; Alberta Environment 2007). The production of 1 m³ of synthetic
1638 crude oil (upgraded bitumen) from surface extraction requires about 12 m³ of water, of
1639 which about 75% is recycled, meaning that 3 m³ must be drawn from licensed sources
1640 (Kasperski and Mikula 2011; CAPP 2012; CEMA 2012). Water used to extract surface
1641 mined bitumen comes primarily (50-75%) from surface waters of the Athabasca River.
1642 Groundwater is also extracted during surface extraction, to decrease the groundwater
1643 table to access the surface deposits, and occasionally as an alternative water source for
1644 processing when river flows are low (King and Yetter 2011). Groundwater is used almost
1645 exclusively in *in-situ* extraction. About 0.5 m³ of groundwater is required to produce 1 m³
1646 of bitumen (CAPP 2012; CEMA 2012). All bitumen extraction and production facilities
1647 are required to recycle water as much as possible; for example, 85% of water used in the
1648 surface extraction process is recycled (CAPP 2008) and 90-95% of water used in *in-situ*
1649 operations is recycled (Humphries 2008).

1650 Current allocation of water from the Athabasca River, for all usages, is 3.5% of
1651 the total annual average river flow, with allocations for oil sands mining projects
1652 accounting for about 2.2% of total flow (AMEC 2007). Less than 10% of the water used
1653 for the oil sands industry is returned to the Athabasca River (Richardson 2007). Despite
1654 recycling, much of the water ends up in the engineered tailings ponds or evaporates from
1655 the pond's surface (Griffiths et al. 2006; CEMA 2012). Several reports have found no
1656 trend or mildly increasing flow in the Athabasca River prior to oil sands development in

1657 the mid-1970s and declines in flow following development (Alberta Environment 2004;
1658 Schindler et al. 2007; Squires et al. 2009). In light of observed declines in flow rates of
1659 the Athabasca River, the percentage of flow removed increases even if water use remains
1660 constant. Flows within the Athabasca River are lowest during winter, due to headwaters
1661 being frozen (Kim et al. 2013); hence there is a high likelihood of water shortages during
1662 winter in the future (Mannix et al. 2010) or at other periods of low precipitation (Kim et
1663 al. 2013). The water management framework for the Lower Athabasca River (Ohlson et
1664 al. 2010) requires water removals be adjusted to meet in-stream requirements (DFO
1665 2010).

1666 Unlike river flows, groundwater flows are not easily observed. Groundwater
1667 dynamics within the oil sands regions are extremely complicated (CEMA 2012;
1668 Government of Alberta 2012b). The surficial geology in the Athabasca oil sands region
1669 consists of Holocene and Quaternary sediments in a complex distribution of tills, outwash
1670 sands, and buried valleys (King and Yetter 2011; Government of Alberta 2013a), creating
1671 a complex hydrogeologic system of aquifers (geologic formations that are permeable and
1672 contain or transmit groundwater) and aquitards (geologic formations with low
1673 permeability to groundwater) (Andriashek and Atkinson 2007). Many of the
1674 hydrogeology investigations date back to the 1970s and 1980s. However, there have been
1675 more efforts to better characterize the subsurface hydrogeology. For example,
1676 Andriashek and Atkinson (2007) acquired and interpreted more than 35 000 new
1677 borehole logs from the oil sands industry to construct a three-dimensional model of the
1678 subsurface including major buried aquifers contained within buried valleys and channels
1679 north of Fort McMurray. However, even with higher density of data, the narrow form

1680 and discontinuous nature of many of the subsurface features that may contain aquifers or
1681 be natural pathways of water and contaminants means that they remain undetected
1682 (Andriashek and Atkinson 2007).

1683 Groundwater needs to be removed for surface oil sand extraction in order to lower
1684 the groundwater table to prevent seepage into the open-pit mine. This creates a draw-
1685 down zone that dries out the surrounding wetland and upland areas (King and Yetter
1686 2011; Government of Alberta 2012b). Draw-downs of 15 to 25 m have been observed
1687 near active mining areas compared to 1 to 4 m in natural water level fluctuation
1688 (WorleyParsons 2010 in Government of Alberta 2012b). Groundwater removals in both
1689 surface and *in-situ* extraction can depressurize aquifers (King and Yetter 2011; CEMA
1690 2012). Control of the hydraulic pressure within the basal aquifer must be maintained to
1691 prevent hydraulic connection between the basal aquifer and groundwater within overlying
1692 sediments (King and Yetter 2011). There are limits set on the volume of groundwater that
1693 can be extracted (Water Conservation and Allocation Guideline for Oilfield Injection;
1694 Alberta Environment 2006). Non-saline groundwater can be drawn-down in the
1695 production aquifer to 35% during the first year of operation and no more than 50% over
1696 the life of the project (Alberta Environment 2006). In 2009, the total licensed volume for
1697 groundwater removal was approximately $52 \times 10^6 \text{ m}^3$ per year while the average annual
1698 volume of water actually withdrawn was 57% of that ($30 \times 10^6 \text{ m}^3$ per year)
1699 (Government of Alberta 2012b). Over half of those licenses were in shallow
1700 groundwater deposits (Government of Alberta 2012b).

1701 Groundwater and surface water are interconnected. The depressurization of
1702 aquifers that accompanies groundwater removals affects local and regional groundwater

1703 flows patterns (King and Yetter 2011; CEMA 2012). Changes in both surface water and
1704 groundwater affect local and regional recharge and discharge to the wetlands and shallow
1705 lakes common in the area (Devito et al. 2005). Schmidt et al. (2010) showed that a
1706 combination of stable isotope and radon mass balance approaches provides information
1707 on water flow path partitioning that is useful for evaluating surface-groundwater
1708 connectivity (i.e., connectivity of lakes and rivers to underlying aquifers).

1709

1710 ***3.6.3 Water quality***

1711 Water quality concerns with oil and gas development are related to both surface
1712 and groundwater contamination. Contamination can occur both at the well or excavation
1713 site (e.g., oil and hydraulic fracturing fluid spills), from tailings ponds (e.g., overflows or
1714 breaches, groundwater seepage), from pipelines during transport, or from groundwater
1715 contamination (e.g., saltwater intrusion, mixing of water with hydrocarbons, groundwater
1716 heating from steam extraction) (National Energy Board 2009; CAPP 2010, 2012;
1717 Environment Canada 2010; Foote 2012; Council of Canadian Academics 2014).

1718 Although some of these are specific to different oil and gas extraction methods, examples
1719 from surface and *in-situ* extraction of the oil sands are given below.

1720

1721 ***3.6.3.1 Oil sands example***

1722 Water quality can be impacted at various points in oil sands processing. Water quality is
1723 usually impaired from contaminated groundwater discharging into surface bodies (i.e.,
1724 streams, rivers, lakes, wetlands) rather than direct effluent release (Gosselin et al. 2010;
1725 King and Yetter 2011). Tailings ponds are an important point source of contaminants

1726 (Government of Alberta 2013a). After bitumen has been extracted from surface mining
1727 of oil sands, the residual material composed of solids and contaminated oil sands process
1728 water is discharged to tailings ponds (CAPP 2012). Water is removed from tailings (i.e.,
1729 dewatered), producing a fine particulate suspension known as mature fine tailings.
1730 Generally, tailings contain approximately 70 to 80% weight basis water, 20 to 30%
1731 solids, and 1 to 3% bitumen (Allen 2008). The released water is reused in the oil sands
1732 extraction process. This remaining water contains elevated levels of soluble salts (due to
1733 continuous recycling of water), dissolved organics (including naphthenic acids), metals,
1734 trace elements, phenols, and low molecular weight hydrocarbons (e.g., polycyclic
1735 aromatic hydrocarbons) (Government of Alberta 2013a). All extraction wastes are
1736 contained in large tailings or settling ponds, constructed using overburden and tailings
1737 and retained by sand dykes constructed with drainage collections systems to intercept
1738 leaking water. The current volume of fluid tailings is estimated at $720 \times 10^6 \text{ m}^3$ and
1739 covers 130 km^2 (Government of Alberta 2012a), an area likely to increase with further
1740 expansion of oil sands mining operations unless regulatory requirements reduce the 40
1741 year historical inventory of tailings (Woynillowicz et al. 2005; Gosselin et al. 2010;
1742 Foote 2012). Directive 074 of Alberta Energy Regulator requires reduction of fluid
1743 tailings and conversion of tailings into trafficable deposits that are ready for reclamation
1744 five years after the deposits have ended (ERCB 2009). New technologies, such as
1745 Suncor's TROTM process (<http://www.suncor.com/en/responsible/3229.aspx>), are steps
1746 towards meeting those targets by converting fluid fine tailings to solid materials that can
1747 be used in reclamation. Reduction in fluid fine tailings will also increase the water

1748 available for recycling, reducing demands for process water from licensed sources
1749 (Kasperski and Mikula 2011).

1750 Water quality impacts to groundwater may occur due to the disruption of natural
1751 or engineered tailings pond barriers, or through vertical or horizontal connections in
1752 groundwater pathways (Gosselin et al. 2010; Government of Alberta 2012b). There are
1753 also naturally occurring geologic formations that impact surface water quality, including
1754 soluble hydrocarbons, salts, and trace elements that leach from bedrock, groundwater,
1755 and wetlands (Government of Alberta 2012b). The transport and persistence of various
1756 contaminants from tailings ponds and separating natural background levels of
1757 contaminants from process-affected contamination continues to be an important research
1758 focus. Leakage rates from tailings containment structures are gaps in knowledge.

1759 A class of contaminants of particular concern because of their toxicity to aquatic
1760 organisms (Kreutzweiser et al. 2013) is naphthenic acid (NA), which is the primary class
1761 of hydrocarbons in the process water. Ambient concentrations of NA in northern Alberta
1762 rivers within the Athabasca Oil Sands region are generally below 1 mg/L, while
1763 concentrations of NA in tailings pond waters can reach over 100 mg/L (Headley and
1764 McMartin 2004). NAs are mobile and persistent in shallow groundwater of the oil sands
1765 mining area. Therefore, little attenuation, beyond dispersive dilution, can be anticipated
1766 along groundwater flow paths in the reclaimed landscape (Ferguson et al. 2009; Oiffer et
1767 al. 2009; Barker et al. 2010). Gervais (2004) found that low molecular weight NA could
1768 be biodegraded under aerobic but not anaerobic conditions. However, Janfada et al.
1769 (2006) found preferential sorption of the individual NAs in soil, which could be
1770 important from a toxicity standpoint because different NA species have varying degrees

1771 of toxicity. Ahad et al. (2013) highlighted the need for accurate characterization of the
1772 diversity of NA species in order to quantify potential seepage from tailings ponds. Savard
1773 et al. (2012) developed a new method of carbon isotopic analysis of carboxyl groups to
1774 distinguish mining-related contaminants from natural background organic acids. These
1775 data were combined with a two-dimensional conceptual model to simulate groundwater
1776 flow and mass transport between tailings ponds and the Athabasca River. Estimates
1777 suggest that mining-related acid extractable organics (e.g., NAs) may be reaching the
1778 river, but groundwater contamination is more of an issue at the local scale (Savard et al.
1779 2012).

1780 Metal contamination is another potential concern. Savard et al. (2012) developed
1781 lead and zinc isotopic methods to discriminate natural from mining-related metal inputs.
1782 They found extraction-related metals were attenuated along the groundwater flow path,
1783 with practically no load being delivered to the Athabasca River (Savard et al. 2012).

1784 Oiffer et al. (2009) also found minimal trace metal mobilization. Barker et al. (2010)
1785 found that process affected groundwater usually contains little trace metals, but mildly
1786 anaerobic process waters may leach toxic metals from the aquifer material and so could
1787 attain undesirable levels of selenium, arsenic, among other metals. However, studies of
1788 current plumes show little accumulation of such toxic constituents (Barker et al. 2010).

1789 Gibson et al. (2011) assessed the potential for labeling process-affected water
1790 from oil sands operations using a suite of isotopic and geochemical tracers, including
1791 inorganic and organic compounds in water. Although selected isotopic and geochemical
1792 tracers were found to be definitive for labeling water sources in some locations, overall it
1793 was unreliable to attempt any universal labeling of water sources based solely on

1794 individual tracers or simple combinations of tracers. They concluded that understanding
1795 the regional hydrogeological system and interpretation of tracer variations in the context
1796 of a systems approach on a case-by-case basis offers the greatest potential for
1797 comprehensive understanding and labeling of water sources and pathways (Gibson et al.
1798 2011).

1799 *In-situ* oil sand extraction creates different water quality issues. *In-situ* wells are
1800 encased to prevent groundwater from moving up or down through vertical layers and to
1801 prevent contamination (saline water or bitumen) from moving into the groundwater
1802 (CEMA 2012). However, breaches in well encasements can degrade groundwater through
1803 disturbance and mixing of saline and freshwater aquifers. Water quality may also be
1804 impaired from improper waste disposal of salts from the water treatment and steam
1805 production process (CEMA 2012). Salt disposal is usually by deep injection, raising
1806 concern for subsurface contamination. Steam-assisted extraction leads to increases in
1807 groundwater temperatures; any increase in temperature is expected to result in negative
1808 effects because of the potential for enhanced solubility and/or mobility of chemical
1809 constituents that are stable at naturally-occurring temperatures (Gosselin et al. 2010;
1810 CEMA 2012). Dewatering in the subsurface can cause increased oxygen in the
1811 subsurface, which could cause oxidative release of some toxic elements, such as arsenic,
1812 into groundwater (Birkham et al. 2007; Gosselin et al. 2010; Holden et al. 2013).

1813 Monitoring is essential to quantifying impacts of oil sands development and
1814 efficacy of mitigation measures. The Regional Aquatics Monitoring Program (RAMP)
1815 was established in 1997 to assess water quality in rivers and lakes of the oil sands region,
1816 including the Athabasca River Delta. RAMP monitors different elements and compounds

1817 within surface waters, including major ions, nutrients (ammonia, total nitrogen, total
1818 phosphorus, dissolved organic carbon), general organics (NAs and phenolics), total
1819 metals (e.g., aluminum, arsenic, cadmium, selenium, and zinc), polycyclic aromatic
1820 hydrocarbons (PAHs) and other constituents like phthalates, acrylamide, petroleum
1821 hydrocarbons, vanadium and dissolved oxygen.

1822 A recent RAMP study (RAMP 2009) attributed the contaminants found in the
1823 Athabasca River and its tributaries to a natural origin. Detailed examination of the results
1824 revealed that there were deviations from the baseline at about 25% of the sites for one or
1825 more of the water quality parameters evaluated, but there were no consistent trends for
1826 which constituents were elevated. Other studies (e.g., Headley et al. 2005; McMaster et
1827 al. 2006a; Conly et al. 2007) also concluded that high concentrations of arsenic (Conly et
1828 al. 2007), metals (Headley et al. 2005), and hydrocarbons (McMaster et al. 2006a) were
1829 of natural origin. Several studies have been critical of the RAMP conclusions. Timoney
1830 and Lee (2009) and Kelly et al. (2009) showed higher PAHs downstream of the oil sands
1831 that were not attributed to natural oil sand. In a more recent paper, Kelly et al. (2010)
1832 showed that elevated priority pollutants (13 elements as identified by the United States
1833 Environmental Protection Agency: antimony, arsenic, beryllium, cadmium, chromium,
1834 copper, lead, mercury, nickel, selenium, silver, thallium, and zinc) were higher in the
1835 snowpack near oil sands developments than at more remote sites. They also showed that
1836 summer stream water priority pollutants were greater near developed and downstream
1837 areas than undeveloped and upstream areas.

1838 Groundwater quality monitoring has been limited (Gosselin et al. 2010), with the
1839 Lower Athabasca Groundwater Management Framework (Government of Alberta 2012b)

1840 assessing groundwater quality data as fair to poor. There are limited spatial and temporal
1841 data for most aquifers. In particular, more groundwater quality monitoring data are
1842 needed for buried valley and buried channels to sufficiently frame baseline conditions
1843 (Government of Alberta 2012b).

1844

1845 **3.6.4 Prognosis**

1846 Water is necessary for oil and gas extractions, particularly in the oil sands. Oil and
1847 gas companies are keenly aware of the need to conserve and recycle the water resources
1848 that are vital to the extraction process (CAPP 2012). The oil sands highlight the need for
1849 water management frameworks to be developed and/or implemented.

1850 Management plans will likely need to be both sound and adaptive to meet future
1851 water needs of an expanding oil sands industry. In 2011, Canada's oil sands production
1852 was 1.7 million barrels per day (Government of Alberta 2013b) and is forecasted to
1853 increase to be 4.3 million barrels per day by 2030 (Prebble et al. 2009). These increases
1854 in oil sands extraction rates, coupled with changes to water flows with a changing climate
1855 (Price et al. 2013) will place increasing pressure on the water resources of the Lower
1856 Athabasca River basin.

1857 More surveys of water resources would support water resource decision making.
1858 Some of the fiercest criticisms of oil sands environmental management have been
1859 directed at the water monitoring program of RAMP, with recent reviews identifying
1860 many gaps in the RAMP monitoring design in the oil sands (Environment Canada 2010;
1861 Main 2011). Gaps included lack of sufficient temporal data to determine historical trends
1862 for key water indicators, limited monitoring coverage outside of the active mineable oil

1863 sands area, a lack of hydrologic and geologic data, and the absence of a regional
1864 groundwater model. As part of a response to these concerns, Environment Canada (2011)
1865 put forth a conceptual framework for an integrated oil sands environmental monitoring
1866 plan that would be holistic and comprehensive, scientifically rigorous, adaptive and
1867 robust, inclusive and collaborative, and transparent and accessible. In early 2012,
1868 Alberta's Minister of Environment and Sustainable Resource Development and Canada's
1869 Environment Minister announced the Joint Canada-Alberta Implementation Plan for Oil
1870 Sands Monitoring (<http://environment.gov.ab.ca/info/library/8704.pdf>). The plan, to be
1871 implemented by 2015, is aimed at enhancing the oil sands monitoring program for air,
1872 land, water and biodiversity. A key aspect of the plan is transparency, with data to be
1873 publicly accessible through the Oil Sands Information Portal
1874 (<http://environment.alberta.ca/apps/osip/>).

1875 As highlighted in the Environment Canada review of RAMP, understanding
1876 groundwater resources both spatially and temporally is a significant knowledge gap, one
1877 echoed in the recently released Lower Athabasca Groundwater Management Framework
1878 (2013). Groundwater quantity and quality issues are only going to become more
1879 important in the future. As of October 2011, groundwater allocation had risen by
1880 approximately 40% from 2009 values to $72.8 \times 10^6 \text{ m}^3$ per year (WorleyParsons 2012).
1881 Greater water demands coupled with requirements for maintaining in-stream flow
1882 requirements are likely to result in greater withdrawals of groundwater. Groundwater is a
1883 hidden resource and, because of uncertainties in its distributions, flows and recharge
1884 rates, may be vulnerable to over exploitation (Struzik 2013). Efforts such as the aquifer
1885 risk assessment methods presented in the Lower Athabasca Region Groundwater

1886 Management Framework for North Athabasca Oil Sands Area (2013) are important first
1887 steps in identifying vulnerable groundwater aquifers.

1888 Groundwater flows at much lower velocities (m per year) than surface water (m
1889 per second), and if contaminated, the time scale for recovery will be much longer
1890 (Gosselin et al. 2010). Thus, groundwater monitoring will be necessary for decades after
1891 operations cease. Much of the focus on groundwater research, and gains in
1892 understanding, has been on contaminant hydrogeology (Lyness and Fennell 2010).
1893 Knowledge of contaminant persistence and transfer is essential, but new collaborative
1894 efforts could be aimed at fingerprinting the natural and mining-related sources of organic
1895 and metallic contaminants in connection with the groundwater flow systems (Savard et
1896 al. 2012).

1897 Surface and groundwater systems are inter-related. As reviewed above, recent
1898 studies have significantly expanded our understanding of shallow groundwater processes
1899 in this hydrogeologically complex terrain. However, much still needs to be done,
1900 particularly in understanding the pivotal role of wetlands, ponds and lakes as connectors
1901 between surface and groundwater systems (e.g., Ferone and Devito 2004; Smerdon et al.
1902 2005; Devito et al. 2012). Further development of isotope and mass balance approaches,
1903 along with other remote sensing platforms that detect soil moisture through dielectric
1904 properties (e.g., Canadian Space Agency's Radarsat; Parks and Grunsky 2002) or gravity
1905 fields (e.g., NASA's Gravity Recovery and Climate Experiment (GRACE); Schmidt et al.
1906 2008), may help to better understand the complex hydrogeologic processes and the
1907 patterns they create in this landscape.

1908 Modeling surface and groundwater hydrology and contaminant persistence and
1909 transport will continue to be an important tool in understanding impacts of current and
1910 future scenarios. However, even the best models can be inaccurate and imprecise.
1911 Monitoring data are essential to improve our understanding of hydrologic processes and
1912 for calibrating and validating model outputs. Site-level modeling is often a patchwork of
1913 different models and development of regional groundwater models will be an essential
1914 step in understanding cumulative effects. Initial steps are being taken to develop
1915 standardized modeling practices in the Oil Sands (Jones and Mendoza 2013).

1916 Cumulative effects assessments have been identified as essential in dealing with
1917 regional water issues (Government of Alberta 2012b). There are multiple disturbances
1918 within the Lower Athabasca watershed including surface and extensive oil sands
1919 developments, mines, and forest management disturbances that contribute to changes in
1920 water quantity and quality. Despite recognition of the importance of quantifying
1921 cumulative effects, a clear approach to achieve this has still to be developed.

1922 A wealth of information has been collected on a variety of topics by consultants,
1923 academics, and government scientists related to oil sands development. The CEMA Oil
1924 Sands Environmental Management Bibliography
1925 (<http://osemb.cemaonline.ca/rrdcSearch.aspx>) and the Oil Sands Research and
1926 Information Network (OSRIN) library (<http://www.osrin.ualberta.ca/>) are examples of
1927 portals to access this information. Recent reports have done thorough reviews of the oil
1928 sands environmental and health impacts (e.g., Gosselin et al. 2010, commissioned by the
1929 Royal Society), but not without criticism (Timoney 2012) or rebuttal (Hrudey et al.
1930 2012). These reviews and critiques highlight the immense scrutiny placed upon the oil

1931 sands industry both from within Canada, but also around the world, to demonstrate
1932 environmental stewardship and leadership in developing the oil sands. Progressive
1933 management frameworks, recognition of the importance of cumulative effects, risk
1934 assessments, and research databases are all steps in the right direction to addressing
1935 criticism, but benefits will only become realized once fully implemented.

1936

1937 **3.7 Peat mining**

1938 ***3.7.1 Introduction***

1939 In some areas of the boreal zone, primarily in Quebec, Manitoba and Alberta, peat
1940 is commercially extracted for sale as a horticultural soil conditioner, personal hygiene,
1941 industrial absorbent, or as a biofuel source (Daigle and Gautreau-Daigle 2001; Poulin et
1942 al. 2004). In 2012, it was an industry that extracted 973 000 tonnes of peat in Canada
1943 with an economic value of approximately \$186M (NRCan 2013*d*). Peat extraction affects
1944 a relatively small portion of the boreal landscape. Environment Canada (2014) estimated
1945 that 260 km² are, or were at some point in the past, drained for peat extraction, with 140
1946 km² currently being actively managed and 110 km² no longer under production. Peat
1947 mining can pose risks to water resources at a local scale because the process almost
1948 always requires the draining and ditching of wetlands, hydrologic disruptions, and the
1949 removal of peat and associated vegetation.

1950 There are two methods for peat harvesting with different environmental impacts
1951 (Gleeson et al. 2006). Dry harvesting involves draining the peatland first to allow solar
1952 drying over a period of time. Following draining, dry peat can be extracted by sod peat
1953 production (blocks of peat are cut and extracted to dry), milled peat production (cutting

1954 and shredding of the top surface, followed by turning over to promote drying), or vacuum
1955 peat production (pneumatic removal of dry upper milled peat). The mined peat is then
1956 further dewatered for production of briquettes or pellets. The alternative method is wet
1957 harvesting, which involves removing peat without any on-site solar drying and
1958 transporting it to a plant for dewatering and thermal drying. This method allows
1959 extraction to occur in areas where drainage is impossible, greatly increases the peat
1960 production season, and is relatively cost-effective. Wet mining, however, is a new
1961 extraction method for which there is little published information on the efficiency and
1962 impacts (Gleeson et al. 2006) and is not currently used on a large scale in Canada
1963 (Environment Canada 2014).

1964

1965 **3.7.2 Water quantity.** Impacts on water quantity from peat harvesting differ depending on
1966 the extraction method (Gleeson et al. 2006). Wet harvesting removes both the peat and
1967 water, thus water drains and evaporates away from the local site. Dry harvesting can
1968 result in losses of water to natural water bodies by diversion to drainage ditches prior to
1969 harvest; when harvesting is completed, the drainage ditches are blocked and water levels
1970 are slowly increased. For both methods, the large amount of peat material that is removed
1971 results in a permanent (i.e., over thousands of years) loss of water storage capacity (Laine
1972 et al. 1995).

1973 Price et al. (2005) reviewed recent studies in peatland hydrology and found that
1974 peatland drainage often resulted in peat compression from seasonal drying and a decrease
1975 in the hydraulic conductivity of the remaining peat by over 75%, and that surface runoff
1976 from affected sites increased by about 25%. Over time, and in the absence of natural or

1977 planned drainage abatement, mined sites continue to get drier. The effects of drying can
1978 continue long after the site has been abandoned because of the altered water flow
1979 patterns, and because the developing shrub community can deposit annual leaf litter
1980 layers to the soil surfaces that restrict water flow to deeper soils (Van Seters and Price
1981 2001, 2002; Price 2003; Price and Whitehead 2004). Colonization of mined sites by trees
1982 can cause further water losses at the site by evapotranspiration (up to 25% of
1983 precipitation) and by interception (up to 32% of precipitation) (Price et al. 2003). Price et
1984 al. (2005) suggest that these impairments of hydrologic function induced by peatland
1985 drainage may have broader implications for local or regional scales where disruptions to
1986 water flow paths within peat-dominated basins could interfere with natural hydrologic
1987 patterns and linkages between uplands and wetlands. In areas where peat mining is
1988 extensive, this could cause cumulative effects and altered hydrologic regimes at a broader
1989 scale (Bedford 1999; Buttle et al. 2005).

1990

1991 ***3.7.3 Water quality***

1992 Undisturbed peatlands affect downstream water quality by interception and
1993 accumulation of inorganic elements, thereby influencing their export to receiving waters
1994 (Weis and Weis 2004; Pelster et al. 2008). Additionally, peatlands are a primary source
1995 of dissolved organic matter that is created by anoxic and acidic conditions and is exported
1996 downstream (Schiff et al. 1998; Nyman 2011). Therefore, disruptions to these
1997 biogeochemical processes in peatlands through peat harvesting activities can influence
1998 the production and hydrologic transfer of such elements (Laine et al. 1995). The changes

1999 in hydrology induced by peat mining can increase levels of suspended sediments and
2000 cause changes in water quality (Daigle and Gautreau-Daigle 2001).

2001 Of the few recent studies we found that reported effects of peat mining in or near
2002 the Canadian boreal zone on downstream water quality, most focused on suspended
2003 sediments. Peatland drainage ditches in New Brunswick (just outside the boreal zone in
2004 the hemiboreal subzone (*sensu* Brandt 2009)) delivered elevated concentrations of
2005 suspended sediments to downstream receiving waters, despite the use of sediment settling
2006 ponds in the peat mining area (Clément et al. 2009). In a separate study, St-Hilaire et al.
2007 (2006) also recorded elevated suspended sediment concentrations in downstream areas
2008 below sediment settling ponds of peat mining areas. Significantly higher concentrations
2009 of suspended sediments were found in streams draining other peat mining sites in
2010 comparison to non-mined peatlands and exceeded provincial guidelines for daily
2011 maximums to protect aquatic organisms about 70% of the time (Pavey et al. 2007).
2012 Prévost and Plamondon (1999) reported elevated concentrations of several elements in
2013 water draining from mined peatlands in Quebec than in nearby natural waters. Although
2014 these increases in sediment and element concentrations could have implications for
2015 biodiversity in receiving waters, there are few published reports on impacts of water
2016 quality changes from peat mining on aquatic organisms (Kreutzweiser et al. 2013).

2017

2018 **3.7.4 Prognosis**

2019 Peat mining is not a serious threat to water quantity and quality in the Canadian
2020 boreal zone, given that only about 0.01% of boreal peatlands is being harvested for peat
2021 (Cleary et al. 2005). However, at a local scale, peat mining can measurably impact water

2022 quantity and quality in mined areas and in downstream receiving waters, with potential
2023 for adverse effects on some aquatic biota (Kreutzweiser et al. 2013). With increasing
2024 interest in peat as a biofuel source (Telford 2009), the area of peat mining in the boreal
2025 zone could expand, with potential impacts to water resources and recovery of wetlands.
2026 Burning peat for heating has a long history in northern Europe, but it is a relatively new
2027 consideration in Canada. Some early trials for co-firing peat at a thermal generating
2028 station were, however, not very promising. Burning peat was not a clean source of
2029 energy, and other issues related to energy costs from drying and processing prior to
2030 burning are likely to limit its potential as biofuel feedstock (Gleeson et al. 2006).

2031 Given that peat accumulates slowly over thousands of years (Ali et al. 2008), peat
2032 mining may not be a sustainable industry at local scales because of the profound
2033 alterations to the hydrologic and therefore ecological functions of peatlands, and because
2034 of the time required to replace similar volumes of peat. Price et al. (2003) point out that it
2035 is still uncertain whether the hydrologic, biogeochemical (carbon storage), and ecological
2036 functions of peatlands can be restored within a reasonable time frame, and suggest that
2037 efforts should be made to retain more water on site during peat extractions to minimize
2038 overall impacts. The peat-mining industry, under the auspices of the Canadian Sphagnum
2039 Peat Moss Association, acknowledges the inherent need for responsible peatland
2040 management to improve the sustainability of their industry (www.peatmoss.com). As part
2041 of their preservation and reclamation policy, they recommend that exploitation methods
2042 minimize the affected surface area, that some plots of natural peat be retained as a buffer
2043 and recolonization source, that vegetation removed should be retained and replanted on
2044 abandoned sites, and that surface re-profiling, mulches, and seepage reservoirs be applied

2045 to reduce hydrologic impacts and to assist vegetation re-establishment. Price et al. (2003)
2046 also recommended blocking ditches, constructing peat banks or terraces, and creating
2047 shallow depressions as means to retain water in rehabilitation efforts.

2048 The International Peat Society's Strategy for Responsible Peatland Management
2049 (Clarke and Rieley 2010) outlines best management principles for sustainable peat
2050 management. The principles are based on conservation of biodiversity, mitigation of
2051 impacts on hydrology and water regulation, implications for carbon cycling and climate
2052 change, and application of the latest technologies in post-mining rehabilitation and
2053 reclamation. It also advocates for the transfer and sharing of knowledge, engagement
2054 with local interest groups, and good governance to ensure conservation of ecological
2055 goods and services from peatlands.

2056

2057 **3.8 Atmospheric emission and deposition of pollutants**

2058 ***3.8.1. Introduction***

2059 Natural resource extraction activities within the boreal zone, as well as other
2060 industrial activities outside of the boreal zone, both from within Canada and around the
2061 world, create emissions that influence boreal water bodies. These emissions initiate from
2062 point sources such as industry, or diffuse sources such as forest fires. Emissions from
2063 industry (Fig. 10) are the gaseous and particulate products of combustions (e.g., nitric
2064 oxides (NO_x) and sulphuric oxides (SO_x)) from biomass and fossil fuel burning, and
2065 volatilized bi-products from extracting and processing resources (see Fig. 6 in Brandt et
2066 al. 2013). These emissions quickly spread from local sources and are transported
2067 throughout the airshed by the prevailing winds eventually falling as wet or dry

2068 deposition. Although the pollution may become diluted in the atmosphere, impacts have
2069 been observed to occur with chronic exposure (e.g., Sudbury, eastern North American
2070 seaboard).

2071

2072 **3.8.2. Water quantity**

2073 Atmospheric emissions of pollutants do not have direct effects on water quantity.
2074 However, particulate emissions, aerosols, and greenhouse gases are known to affect
2075 atmospheric radiative balances (either absorb or reflect solar radiation depending on the
2076 particle) as well as serving as cloud condensation nuclei (Gieré and Querol 2010). These
2077 mechanisms have largely unknown (Cotton and Yuter 2009) but likely small impacts on
2078 local or regional precipitation patterns within the boreal zone.

2079

2080 **3.8.3. Water quality**

2081 Surface water quality is influenced by wet and dry atmospheric deposition of
2082 pollutants either through direct input onto surface waters or indirectly through soil-
2083 mediated effects. When sulphate- and nitrate-based emissions combine with
2084 precipitation, the resulting acidic deposition can reduce pH in soil and surface waters.
2085 Acidification synergizes metal export and accumulation in receiving waters (LaZerte
2086 1986) and confounds, often increasing, the bioavailability and toxicity of metals to
2087 aquatic organisms (Keller and Pitblado 1986; Schindler 1988). A vast literature exists on
2088 the impacts of acidification and associated metal toxicity on aquatic ecosystems, and
2089 demonstrates that when pH levels in receiving waters have dropped to or near 4, drastic
2090 and long-lasting effects on aquatic organisms can occur (e.g., Keller and Yan 1991).

2091 Atmospheric nitrogen deposition also impacts the amount of nitrogen within forest soils
2092 having impacts on forest productivity, soil nutrient cycling, and nutrient (particularly C
2093 and N) exports (Hyvönen et al. 2007).

2094 There is a long history of studying acid rain impacts in eastern North America.
2095 The eastern parts of the Canadian boreal zone and the hemiboreal subzone were
2096 traditionally the area of highest acidic deposition and acidification of surface waters
2097 (Environment Canada 2005a). From the 1960s through to the 1980s, emissions in eastern
2098 North America were high because of combustion of fossil fuels used in industry and in
2099 thermal hydroelectric generation within the industrial heartland of eastern United States
2100 and southern Canada. Since the implementation of the United States Clean Air Act
2101 (1990) and the Canada-United States Air Quality Agreement (IJC 1991), SO_x levels have
2102 dramatically been reduced, and NO_x has levelled off, or shown a slight decline. Controls
2103 on sulphur dioxide emissions in North America resulted in large reductions in sulphate
2104 concentrations and increases in pH in many eastern boreal lakes by the mid-1990s (Keller
2105 2009). Numerous studies are tracking the chemical and biological recovery of water
2106 bodies from acidification (Keller et al. 1992; Gunn 1995; Carbone et al. 1998; Doka et al.
2107 2003; Jeffries et al. 2003a, 2003b; Snucins and Gunn 2003; Clair et al. 2007; Keller et al.
2108 2007; Gray and Arnott 2009). Overall, these show that the pH of surface waters in most
2109 areas near historic industrial operations is improving, but that metal concentrations often
2110 remain elevated above reference or target levels and that biological recovery usually lags
2111 behind chemical recovery.

2112 Although improvements have been made at regional (northeastern United States
2113 and Canada) and local scales (e.g., Sudbury; Keller 2009), parts of the boreal zone in

2114 Alberta and Saskatchewan have soils with relatively low buffering capacity, and therefore
2115 water bodies in that region are at risk of acidification from atmospheric pollutants
2116 (Hazewinkel et al. 2008; Aherne and Shaw 2010; Scott et al. 2010; Whitfield et al. 2010).
2117 Deposition of SO_x and NO_x in the oil sands region of Alberta has increased over the last
2118 40 years, although levels have shown some decline over the period 2000 to 2005 with the
2119 introduction of sulphur capture technology (Alberta Environment 2008) and are low
2120 compared to areas of eastern North America (Whitfield et al. 2010). Scott et al. (2010)
2121 surveyed 259 headwater lakes within 300 km of Fort McMurray, AB. They found that
2122 60% of the lakes were classified as sensitive and 8% as very sensitive to atmospheric
2123 deposition of acidic pollutants. They determined that although acidification appears not
2124 to be significantly advanced, many dilute oligotrophic lakes with pH 6 to 6.5 are
2125 vulnerable to acidic pollutants.

2126 The absence of evidence for acidification does not necessarily imply that
2127 emissions from the oil sands region are environmentally benign, but rather suggests that
2128 the biogeochemistry of these lakes differs fundamentally from well-studied acidified
2129 counterparts in eastern North America and northern Europe (Hazewinkel et al. 2008).
2130 Model simulations suggest limited risk of acidification primarily due to sulphur retention
2131 in the basins, however drought may induce episodic depression of acid neutralizing
2132 capacity (Whitfield et al. 2010). Thus, it has become apparent that the recovery of lakes
2133 from acidification is closely linked with the responses to, and interaction with, other
2134 large-scale environmental stressors like climate change and calcium declines (Keller
2135 2009).

2136 Across the boreal zone, trends in critical loads (i.e., an estimate of an exposure to
2137 one or more pollutants below which significant harmful effects on specified sensitive
2138 elements of the environment do not occur according to present knowledge) of acidity and
2139 their exceedances (i.e., acid deposition load minus critical load) show that many areas
2140 have high sensitivity to acidification and exceedance of critical loads (Carou et al. 2008)
2141 (Fig. 11). These areas are mostly characterized by non-carbonate bedrock and shallow,
2142 coarse-textured, upland soils that have a low buffering capacity. A primary goal of *The*
2143 *Canada-Wide Acid Rain Strategy for Post-2000* is to meet the environmental threshold of
2144 critical loads for acid deposition across Canada, by pursuing further acidifying emission
2145 reductions in Canada and the United States, preventing pollution and protecting
2146 ecosystems not yet impacted by acidifying emissions. These national maps of the
2147 sensitivity of ecosystems serve as a useful means to illustrate where and by how much
2148 forest ecosystems are at risk of sulphur and nitrogen deposition damage, and therefore
2149 where and by how much emission controls will be needed in order to protect sensitive
2150 ecosystems (Carou et al. 2008).

2151

2152 **3.8.4. Prognosis**

2153 Investigations into atmospheric deposition of acidic pollutants from processing
2154 emissions and their impacts on water quality are ongoing. Emissions of acidifying
2155 pollutants have declined by about 50% nationally, and by much more in some areas
2156 (CCME 2011). However, atmospheric deposition still exceeds critical loads, and
2157 acidifying emissions are still increasing in other areas (FPTC 2010). To the extent that
2158 natural resource development activities and their related emissions are likely to increase

2159 in some areas of the boreal zone (e.g., oil sands region, Hudson Bay lowlands), the
2160 deposition of acidic pollutants on landscapes with concomitant impacts on water quality
2161 are likely to increase at local to regional scales. To further reduce emissions,
2162 consideration could be given to improve scrubbing technologies, increased processing
2163 efficiencies, and cleaner fuels.
2164

2165 **4. Cumulative effects and non-linear, threshold and** 2166 **tipping point behaviors**

2167

2168 Although much of the boreal zone remains unaffected by natural resource
2169 development activities, some regions, particularly in southern parts, have been affected
2170 substantially (Fig. 2). This anthropogenic footprint has resulted in fragmentation of the
2171 landscape (Schneider et al. 2003). Many continuing industrial activities have reduced
2172 their disturbances, but still produce measurable cumulative impacts to water quantity and
2173 quality over local to regional scales, and legacy impacts of former industrial installations
2174 in particular will persist for decades (Table 3; Niemelä 1999; Ptacek et al. 2004; Keller et
2175 al. 2007).

2176 The release of pollutants and contaminants into surface waters can occur not only
2177 at the site of disturbance but can also be transported many kilometers downstream (Culp
2178 et al. 2003; Muscatello et al. 2008; Kelly et al. 2010). Individual effects of headwater
2179 disturbances can potentially accumulate and affect downstream ecosystems and their
2180 services (Townsend et al. 2008). Multiple effects of point and diffuse disturbances within
2181 and among industrial sectors that are hydrologically connected can also accumulate over
2182 the landscape (Keller 2009; Seitz et al. 2012), leading to greater uncertainty and
2183 unpredictability of ecosystem response (Dubé 2003; Gunn and Noble 2011). It is
2184 expected that both individual and cumulative effects resulting in non-linear and tipping
2185 point behaviour will become the new “norm” in boreal ecosystems in response to
2186 intensified and/or expanded natural resource development (Kinzig et al. 2006; Foote
2187 2012).

2188 Cumulative Effects Assessments (CEAs) are a requirement under various
2189 provincial Environmental Impact Assessment laws and regulations. CEAs examine the
2190 interactions between changes in basin structure and function that accumulate and the
2191 response of the river ecosystem under different futures scenarios of development within
2192 the basin (Seitz et al. 2011). A CEA demands that proponents examine the cumulative
2193 effects associated with their proposed development alongside relevant past, present, and
2194 future projects. The CEA is often done at the project scale, not at the regional scale
2195 (Squires et al. 2009). CEAs are methodologically complex as cumulative effects are
2196 frequently interactive or synergistic in nature (Dubé 2003) and there is not a single
2197 conceptual approach that is widely accepted by both scientists and managers (Squires et
2198 al. 2009). Overcoming challenges set forth by scaling issues, diverging views, different
2199 policies and legislations, and complex ecological pathways is in itself the main challenge
2200 for those who try to carry out an effective CEA (Gunn and Noble 2011; Seitz et al. 2011).
2201 Examples of effects-based CEAs in the boreal zone include the Moose River Basin study
2202 (Munkittrick et al. 2000), Northern Rivers Basin Study (Culp et al. 2000), Northern
2203 Rivers Ecosystem Initiative (Dubé et al. 2006), and Athabasca River Basin study (Squires
2204 et al. 2009).

2205 Seitz et al. (2011) suggest that regional CEAs should take a multiple stakeholder
2206 approach, with government assuming leadership to establish management objectives as
2207 well as compliance and effectiveness monitoring at the regional basin scale. The
2208 Cumulative Effects Monitoring Association (CEMA), within the Regional Municipality
2209 of Wood Buffalo, AB, is an example of a multiple stakeholder group established to
2210 produce recommendations and management frameworks pertaining to the cumulative

2211 effects of oil sands development in northeastern Alberta, which are, once complete,
2212 forwarded to provincial and federal government regulators
2213 (<http://cemaonline.ca/index.php/about-us>). Technical and scientific work performed for
2214 CEMA is completed through the direction of working groups (Land, Reclamation, Air,
2215 Wetland, Traditional Knowledge) using smaller task groups to focus on specific issues
2216 (<http://cemaonline.ca/index.php/working-groups>). Creation of a regional CEA is,
2217 however, not part of its mandate.

2218

2219 Climate change complicates efforts to understand cumulative effects within
2220 ecosystems. For example, long-term increases in temperature and changes in
2221 precipitation, coupled with increased frequency of extreme events such as droughts and
2222 floods (as anticipated for Canada's boreal zone by Price et al. 2013), may create
2223 ecosystem instabilities that promote non-linear responses (Keller et al. 2007; Vose et al.
2224 2011).

2225 Innovative research methods that use spatially explicit data and long-term
2226 monitoring networks would support efforts to quantify cumulative effects of natural
2227 resource development activities on basin processes at large spatial and temporal scales
2228 (Creed et al. 2011a; Vose et al. 2011). Development of predictive models may also help
2229 to make informed decisions on projected impacts of future natural resource extraction
2230 activities under different climate and possible mitigation scenarios (Chapin et al. 2004;
2231 Creed et al. 2011a), and require rigorous scientific data, and a tighter coupling between
2232 management and monitoring activities (i.e., adaptive management; see Gauthier et al.
2233 2014). Furthermore, model outputs need to be better integrated into decision-making

2234 frameworks for prioritizing boreal forest management strategies that minimize impacts of
2235 natural resource development on boreal aquatic ecosystems and the services they provide.
2236 This is an active research focus of the Canadian Network of Aquatic Ecosystem Services
2237 in the Healthy Forests, Healthy Waters Theme (CNAES 2014).

2238

2239 **5. Improving water stewardship**

2240 Natural resource industries have and will continue to require access to water, a
2241 need that will likely continue to expand (UNEP 2009; NRTEE 2011). This review has
2242 identified that water quantity and quality has been impacted by industrial development in
2243 some areas of the boreal zone. Although some natural resource developments have
2244 localized, short-term effects on water resources (e.g., forest harvesting at stand or sub-
2245 basin scales), others have localized, long-term effects (e.g., mining activities and their
2246 effluents), and still others have regional, long-term impacts (e.g., processing emissions,
2247 hydroelectric power generation, oil sands development). Certain areas of the boreal zone
2248 are at increased risk of cumulative effects of multiple natural resources development on
2249 water resources (e.g., Lower Athabasca Watershed). Effectively managing impacts of
2250 natural resource development across multiple spatial and temporal scales of Canada's
2251 boreal zone is challenging. Improving water stewardship performance in the natural
2252 resource sector could include the following considerations.

2253

2254 **5.1 Awareness of water resource values**

2255 The places where humans benefit from ecosystem services is often different from
2256 the places that produce those services, thus feedbacks to ensure the continued provision

2257 of these services is often missing (Brauman et al. 2007; Keeler et al. 2012). This is
2258 particularly relevant for the water resources of the boreal zone, where water-related
2259 ecosystem services generated in source areas are often hundreds of kilometers away from
2260 the downstream human settlements that benefit from those services. The lack of a clear
2261 connection between the locations of the supply of services and the beneficiaries of those
2262 services may reduce the recognition of the value of water-related ecosystem services
2263 (Brauman et al. 2007).

2264 Accurate valuation of water resources may increase the public awareness of their
2265 value and encourage more efficient water use (Hoover et al. 2007). Calculation of a water
2266 footprint, which is the volume of water appropriated to produce a product taking into
2267 account the water consumed and polluted in the different steps of the supply chain, is one
2268 approach to water valuation (Hoekstra and Mekonnen 2012; ISO 2012; Launiainen et al.
2269 2013). Although there is some criticism of water footprint estimates, which vary
2270 dramatically depending on the methodology used, the concept has had notable success in
2271 raising awareness about water use by providing a previously unavailable and seemingly
2272 simple numerical indicator of water use (Chenoweth et al. 2013).

2273 Economic valuation of water-related ecosystem services can be a powerful tool
2274 for policy evaluation because it provides a common metric with which to make
2275 comparisons between natural resources and natural capital (Heal 2000). However, values
2276 associated with ecosystem goods and services are challenging to estimate. Frequently,
2277 there are insufficient data to attempt to quantitatively incorporate all environmental value
2278 information into prices or to construct accurate full cost accounting (Adamowicz et al.
2279 2007).

2280 As one example within the boreal zone, Anielski and Wilson (2005) estimated
2281 Canada's natural capital based on boreal ecosystem services at a total non-market value
2282 of \$93.2B per year based on 2002 Canadian dollars, approximately 2.5 times greater than
2283 the net market value of boreal natural capital extraction (Anielski and Wilson 2005).
2284 They later revised the ratio upward to 13.8, primarily from a jump in nonmarket
2285 ecological service values due to revaluation of stored carbon in forest and wetlands
2286 (Anielski and Wilson 2009). The majority of the non-market values of the boreal zone
2287 was attributed to water-dependent services, in particular, the ecosystem services provided
2288 by wetlands (i.e., groundwater aquifer recharge, water purification, flood control, and
2289 biodiversity) (Anielski and Wilson 2005). The valuation of ecosystem services continues
2290 to be an active and important field of investigation for environmental economists (Liu et
2291 al. 2010; Molnar and Kubiszewski 2012). Considerable research effort is being focused
2292 on improving methodologies so that results are transparent and defensible (Liu et al.
2293 2010; Salles 2011).

2294 **5.2 Governance of water resources**

2295 Water is one of the most challenging natural resources to govern and manage due
2296 to the division of power and shared responsibilities for management among all levels of
2297 government (de Loë 2008; NRTEE 2011; Canada West Foundation 2011). Water
2298 governance in Canada has been founded on a collection of statutes and policies, involving
2299 all levels of government (Saunders and Wenig 2007; Hipel et al. 2011). Water
2300 governance within the boreal zone is primarily under provincial or territorial
2301 responsibility that includes a mix of federal and provincial legislation, even though many
2302 boreal basins cross multiple jurisdictions and ecozones.

2303 The current federal water policy dates back to 1987 ([Environment](#) Canada 1987)
2304 and the federal wetland conservation policy to 1991 (Government of Canada 1991).
2305 These policies were never fully implemented and have been criticized for not reflecting
2306 new pressures on water resources from expanding industries, water diversions,
2307 international water export, and climate change (Morris et al. 2007; Muldoon and
2308 McClenaghan 2007; Sanders and Wenig 2007; NRTEE 2011). National consistency on
2309 environmental issues concerning water could be achieved through the Canadian Council
2310 of Ministers of the Environment (CCME) (de Loë 2008; Hipel et al. 2011). The CCME
2311 serves as a principal forum for members to develop national strategies, norms, and
2312 guidelines for setting environmental standards (Hipel et al. 2011).

2313 There have been repeated calls from a diverse range of groups and sectors for
2314 renewed federal action on water (Morris et al. 2007; Pollution Probe 2008; de Loë 2008,
2315 2009; McLaughlin 2009; Hipel et al. 2011; NRTEE 2011). Constitutional and practical
2316 considerations require that leadership in water governance come from both federal and

2317 provincial governments (Morris et al. 2007; McLaughlin 2009). However, Hipel et al.
2318 (2011) suggest that national water governance is best led by the federal government given
2319 its experience in cooperatively managing water resources in shared basins with the
2320 United States *via* the International Joint Commission that was established under the
2321 Boundary Waters Treaty of 1909.

2322 There are differing opinions as to the form in which federal action on water could
2323 take, whether it be a revised federal water policy, water framework, or pan-Canadian
2324 water strategy. Despite differences in the name attached to the action, most proposals
2325 share a similar vision and approach (Bakker 2007; Morris et al. 2007; de Loë 2008, 2009;
2326 NRTEE 2011). Their shared vision is to achieve a comprehensive and coordinated
2327 approach to water governance in Canada as a platform for addressing water-related
2328 challenges and opportunities that demand a national perspective. A guiding principle, as
2329 proposed by NRTEE (2011), is that water has economic, environmental and social values
2330 and should be managed in trust without harm to its sustainability or to that of the
2331 ecosystems in which it occurs or flows through. Proponents of a national water strategy
2332 suggest that a key element of the strategy could be a national water council that would
2333 include basin management authorities. These authorities would work with the provinces
2334 in undertaking the day-to-day water management functions. The national water strategy
2335 could focus on priority areas, such as enhancing national capacity for freshwater
2336 protection, responding to the impacts of climate change, ensuring sustainable natural
2337 resource development, securing safe drinking water for all Canadians, protecting aquatic
2338 ecosystems and aboriginal water rights, promoting water conservation, circumventing

2339 inter-jurisdictional and trade conflicts, and supporting world class water science (Morris
2340 et al. 2007).

2341 Some proponents of national action on water resources (e.g., Bakker 2007; Morris
2342 et al. 2007) suggest that Canada's water strategy could follow the European Union (EU)
2343 Water Framework Directive (Commission of the European Communities 2007). Since
2344 2000, the EU has applied a collaborative approach that facilitates action across political
2345 and cultural borders by investing in science, knowledge sharing, and a common
2346 operational approach to water management (Lagacé 2011). Under the guiding principles
2347 of harmonization and subsidiarity, the goal of the directive is to establish common
2348 standards and practices across the EU that safeguard water quantity and quality of for the
2349 future (Lagacé 2011).

2350 Proponents of a national water strategy for Canada have suggested that it should
2351 include the following themes. First, water resources within basins would be treated as a
2352 whole within an integrated system, acknowledging that many interconnected water-
2353 related factors must be considered and that the trade-offs among competing stakeholder
2354 and environmental requirements be taken into account (Global Water Partnership and
2355 International Network of Basin Organizations 2009; Hipel et al. 2011). Second, water
2356 governance and related policies and management would be adaptive in order to handle
2357 the largely unpredictable behavior of the environment and society brought about by
2358 intrinsic complexity, uncertainty, and interconnectedness (Hipel et al. 2011) and by
2359 continually learning (through on-going assessments) from the success (or failure) of
2360 policies that are implemented. Third, water governance and management would be
2361 collaborative, integrating and coordinating federal, provincial, territorial, First Nation

2362 policies as well as a broad range of other stakeholders (Berkes and Folk 1998; Blatter and
2363 Ingram 2001; Berkes 2002; Finger et al. 2006; Bouleau 2008).

2364 Despite widespread agreement in terms of the purpose and approach of a national
2365 water strategy, its development and implementation would not be straightforward
2366 (Bakker 2007). Experiences in jurisdictions such as Brazil, Australia, South Africa, and
2367 the EU that have developed or are developing water strategies illustrate these difficulties
2368 (de Loë 2008). Regardless of the challenges, de Loë (2008) points out that the benefits of
2369 a national water strategy would include: (1) an improved position for meeting growing
2370 international expectations and obligations; (2) a stronger capacity to respond to threats
2371 (e.g., climate change, extractions, contaminants, legacy effects) and opportunities (e.g.,
2372 funding programs); (3) a clarification of responsibilities among jurisdictions; (4) a greater
2373 consistency in responding to concerns; and (5) more effective decision making when
2374 problems transcend jurisdictional boundaries.

2375 Although a formalized national water strategy could improve effective and
2376 consistent decision making for water stewardship (Bakker 2007), its absence is not
2377 necessarily a barrier to making progress in managing boreal water resources. A number
2378 of provinces and territories have recently developed water strategies including Quebec
2379 (2002), Alberta (2003), Manitoba (2003), British Columbia (2008) and Nova Scotia
2380 (2010) and NWT (2010) (Morris et al. 2007; de Loë 2008; NRTEE 2011). Most
2381 provinces or territories have some form of basin management organization, although the
2382 roles and responsibilities of these organizations vary among jurisdictions.

2383

2384

2385 **5.3 Innovative water management**

2386 Several reports have illustrated that water supply challenges can often be traced to
2387 misallocation rather than to water scarcity (Environment Canada 2004; Schoengold and
2388 Zilberman 2007; Muldoon and McClenaghan 2007; Hipel et al. 2011), although this may
2389 change in the future as a result of the anticipated consequences from climate change.
2390 Misallocation can be the result of poor or insufficient baseline data (NRTEE 2009).
2391 Accurate, complete, and current water-quantity data are critical in establishing water-
2392 management systems in which water is effectively allocated and efficiently used
2393 (McLaughlin 2009). Sustainable water management requires current and accurate
2394 information on water supply and demand, and reasonable estimates of how supply and
2395 demand may change in the future.

2396 On the supply side, current data on water quantity, monitoring capacity, and
2397 reporting protocols are available through the National Hydrometric Program operated by
2398 the Water Survey of Canada, and often supplemented by provincial monitoring.
2399 However, the program has been criticized for lacking a long-term plan for expanding and
2400 maintaining monitoring stations, and lacking clear directions for adapting to water-
2401 quantity threats such as climate change (NRTEE 2011). These concerns are particularly
2402 relevant in the boreal zone where there is a low density of monitoring sites (Fig. 6),
2403 increasing industrial development, and high vulnerability to climate change (Schindler
2404 and Lee 2010).

2405 On the demand side, there are gaps in availability and discrepancies in the quality
2406 of data collected across provinces (Hipel et al. 2011; NRTEE 2011). Demand estimates
2407 are often based on water allocation permits, but the actual quantity taken or returned to

2408 water bodies is typically unknown (NRTEE 2011). However, overall trends in water
2409 demand during the period 1981 to 2005 indicate that water extraction or use did not
2410 increase at the same rate as economic growth, reflecting improved water-use efficiencies
2411 and conservation among most natural resource industries (NRTEE 2011). These
2412 efficiencies may offset increasing water demand from expanding natural resource
2413 development in Canada, such that the overall demand for water by the natural resource
2414 sector at the national scale is expected to hold steady or increase marginally over the next
2415 couple of decades (NRTEE 2011). There may be exceptions at a regional scale. For
2416 example, the largest increases in industrial water use in Canada's boreal zone by 2030
2417 that are expected to present water supply challenges are associated with oil sands
2418 development (both *in-situ* and surface mining) (Bruce et al. 2009; Hipel et al. 2011).

2419 There is growing recognition across Canada's natural resource sectors that water
2420 stewardship be considered in business planning and that boreal ecosystem security be
2421 incorporated in water allocation decisions (McLaughlin 2009). Economic instruments
2422 could facilitate these considerations while addressing water stewardship performance
2423 outside of a government framework (Brauman et al. 2007). Economic instruments are
2424 emerging tools for water management that complement, rather than replace, traditional
2425 government control approaches (Shrubsole and Draper 2007). Economic instruments
2426 provide incentives for behavioural change, generate revenue for financing environmental
2427 initiatives, promote technological innovations, and reduce wasteful water usage at the
2428 lowest cost to society (Shrubsole and Draper 2007; NRTEE 2011). Economic instruments
2429 include water surcharges, tradable water permits, subsidies, and financial incentives
2430 (Shrubsole and Draper 2007).

2431 Water pricing on a volumetric basis can help achieve water reduction objectives,
2432 with modest impacts to most sectors and the national economy (NRTEE 2011;
2433 McLaughlin 2009). Typically, industrial water-use systems run on a cost-recovery basis,
2434 with fees that are set very low and that do not reflect the true value of the water being
2435 used (Bruce et al. 2009). For example, the average cost of gross water use is about \$0.13
2436 per m³ across all natural resource sectors. To achieve a 40% reduction in water intake
2437 would require an increase in price ranging from \$0.50 to \$0.70 per m³ (NRTEE 2011).
2438 Charging for the use of water at prices that more closely reflect the actual value of water
2439 used in natural resource development could promote responsible water use (McLaughlin
2440 2009). For example, incentives through water pricing can lead to innovation in
2441 technologies for reducing water uptake by increasing recirculation or recycling
2442 (Adamowicz et al. 2007).

2443 Water trading (the reallocation of water permits from license holders with a
2444 surplus of water to those with a need) is another economic instrument that exists in many
2445 parts of the world, but within Canada, it exists only in Alberta and only at a very limited
2446 scale (Adamowicz et al. 2007; Shrubsole and Draper 2007; NRTEE 2011). It can be a
2447 transformational strategy in the sense that trading water rights would be a fundamental
2448 shift in water management such that water managers and regulators would become
2449 market managers that could potentially drive up water prices. Such a transformational
2450 strategy warrants caution, time, and careful thought to implement (NRTEE 2011).
2451 Practical experience in water markets remains limited, with other countries having
2452 demonstrated both successes and failures in implementation (NRTEE 2011). A move
2453 toward a national water trading scheme would require an understanding of the level of

2454 acceptance for, and potential negative implications of, water trading, and appropriate
2455 institutional and legal safeguards would need to be included (Adamowicz et al. 2007).

2456 Financial incentive programs are another form of economic instruments that can
2457 result in improved water stewardship performance. For example, the Pulp and Paper
2458 Green Transformation Program of Natural Resources Canada funded projects that
2459 promoted technologies for improved environmental performance in the pulp and paper
2460 sector through increased efficiencies (NRCan 2010). Incentive programs can be more
2461 effective than tax subsidies or other mechanisms that discount the product at the expense
2462 of the ecosystem (Anderson et al. 2010).

2463 Voluntary initiatives can also promote water stewardship, and are taken on by
2464 industries in the absence of government intervention. Voluntary initiatives are guided by
2465 a shared commitment among participating organizations to achieve a desired outcome
2466 such as increasing requirements to respond to market and customer expectations,
2467 maintaining a social license to operate, enabling internal management and performance
2468 improvement, demonstrating environmental responsibility, and internally addressing
2469 knowledge gaps (NRTEE 2011). The NRTEE (2011) provided several examples of
2470 voluntary initiatives among natural resource sectors that included: 1) industry-driven
2471 performance initiatives with an emphasis on improved water-use efficiencies (e.g., Oil
2472 Sands Leadership Initiative, Global Social Compliance Programme, International Council
2473 on Mining and Metals); 2) standards and certification programs to improve
2474 environmental and social management practices, to promote brand recognition, and to
2475 demonstrate responsible resource development (e.g., Alliance for Water Stewardship,
2476 ISO 14046 Water Footprint, Forest Stewardship Council Certification for Ecosystem

2477 Services); 3) international reporting initiatives to promote transparency and
2478 accountability among natural resource sectors (e.g., Global Reporting Initiative, Carbon
2479 Disclosure Project – Water Disclosure); and 4) accounting and management initiatives to
2480 identify risks and areas for performance improvement, and to demonstrate corporate
2481 social responsibility (e.g., WBCSD Global Water Tool, Global Environmental
2482 Management Initiative).

2483

2484 **5.4 Integrated water management**

2485 Effective water management can be attained through adaptive and integrated
2486 ecosystem-based approaches that incorporate collaborative dialogue from multiple
2487 stakeholder groups (Berkes and Folk 1998; Blatter and Ingram 2001; Berkes 2002; Finger
2488 et al. 2006; G3 Consulting Ltd. 2009). Under this scheme, water is often managed at
2489 tertiary (local to regional) basin scales, but is considerate of implications at regional,
2490 national, and international scales. Integrated regional water planning brings together local
2491 management plans and integrates them over large areas that include multiple sectors.
2492 Undertaking integrated water management is challenging in that it requires knowledge of
2493 regional surface water and groundwater hydrology. Furthermore, natural resource sectors
2494 are often disproportionally dispersed over the region, and each sector may have disparate
2495 water stewardship standards and water allocation requirements such that the cumulative
2496 effects on water quantity and quality at a regional scale are unknown (NRTEE 2009).

2497 Creed et al. (2011b) suggest that integrating water management plans at regional
2498 scales could be improved by following a common set of management-oriented hydrologic
2499 principles. They proposed six principles and complementary management actions that

2500 link natural resource development to water stewardship. These principles are: 1)
2501 determine hydrologic system boundaries and consider the entire hydrologic system where
2502 development or management actions take place; 2) conserve critical hydrologic features
2503 by minimizing disturbance to areas involved in the source, movement, and storage of
2504 water; 3) maintain connections between hydrologic features by minimizing disruptions to
2505 water, sediment, and nutrient flows; 4) respect the temporal variability in hydrologic
2506 processes, over short-term (i.e., day-to-day operations) and long-term time scales (i.e.,
2507 100-year planning horizons) (e.g., instream flow requirements); 5) respect the spatial
2508 heterogeneity in hydrologic processes among different scales within a basin (e.g., stand,
2509 hillslope, basin) and among different hydrologic regions (e.g., discharge dominated vs.
2510 evapotranspiration dominated); and 6) maintain redundancy and diversity of hydrologic
2511 form and function within basins. Maintaining undisturbed sub-basins or other contiguous
2512 areas of forest and peatlands within a basin can also help to protect water resources
2513 because they prevent or slow human-caused alteration of hydrology from industrial land-
2514 use activities and can prevent or slow the spread of pollutants and invasive species (Wells
2515 et al. 2010; Andrew et al. 2014).

2516 Integrated water management at a regional scale requires knowledge of the
2517 fundamental hydrologic processes at work. As previously discussed (see section 2.1), key
2518 hydrologic processes differ across the boreal zone and these differences may influence
2519 the effectiveness of water management strategies. Therefore, integrated water
2520 management plans at regional scales across the boreal zone will need to adjust for these
2521 differences. However, this is increasingly challenging because knowledge gaps remain
2522 with regard to some of the underlying hydrologic processes, such as surface water and

2523 groundwater connections and their susceptibilities to natural resource development
2524 activities under differing regional conditions. In conjunction with these knowledge gaps,
2525 additional uncertainties are likely to arise from climate change impacts (Bruce et al.
2526 2009; McLaughlin 2009). Further to these challenges for integrated water management,
2527 Canada's water science capacity to provide current and reliable data appears to have
2528 declined in recent decades (McLaughlin 2009) and resources for systematic water-related
2529 research and monitoring have not kept pace with expanding natural resource development
2530 (Bruce et al. 2009). In the absence of current data and regional-specific information, a
2531 precautionary approach to water resource management is required (Rosenberg
2532 International Forum 2013).

2533 Even with adequate knowledge of hydrologic processes within the basin being
2534 managed, reliable data on water demand and the potential impacts of natural resource
2535 development across all sectors are required for integrated water management. Providing
2536 these data to managers for basins in the boreal zone, particularly in remote areas, is
2537 difficult because data sources are often incomplete and inconsistent (McLaughlin 2009).
2538 Others have pointed out that reliable measurements or estimates of basic water inventory
2539 metrics, such as wetland areas, lake volumes, river and glacial runoff, groundwater
2540 resources, and pollutant loadings are often scarce or absent at regional scales, thereby
2541 precluding the application of effective integrated water management strategies (Nowlan
2542 2007; Bruce et al. 2009; Hipel et al. 2011).

2543 In addition to regionally-specific information on hydrologic processes, water
2544 resource inventories, natural resource development activities, water demands, and risks to
2545 water quantity and quality, integrated water management approaches require data

2546 processing tools (Morris et al. 2007; Bruce et al. 2009). A critical gap is tools to
2547 synthesize information spatially and temporally, and to model cumulative effects, to
2548 make predictions about future scenarios and to forecast risks. Spatially distributed
2549 hydrologic models are available and can provide some assistance, but they are often
2550 incomplete and difficult to parameterize (Creed et al. 2011a). Although there have been
2551 advances in surface water models and improvements to groundwater models (Beckers et
2552 al. 2009), coupled surface water – groundwater models are still in their infancy and are
2553 important for understanding risks to water resources from natural resource development.
2554 Models that link natural resource development and water resources to economic
2555 endpoints are particularly uncommon (Bruce et al. 2009). Bruce et al. (2009) also point
2556 out that uncertainties in model predictions due to model assumptions and limited
2557 calibrations need to be clearly communicated to end-users.

2558 There are several emerging examples of integrated water management in the
2559 boreal zone. Alberta Environment and Sustainable Resource Development has river
2560 management frameworks for many of their key boreal basins. For example, the Lower
2561 Athabasca Surface Water Quality Management Framework (Government of Alberta
2562 2012c) and the Lower Athabasca Groundwater Management Framework (Government of
2563 Alberta 2012b) are comprehensive in their scope in assessing water resources and
2564 stressors to these water resources in the basin. Both of these frameworks are new,
2565 however, and it will take many years for them to be implemented and even more years to
2566 determine their effectiveness.

2567 A further example of integrated water management planning is provided by a
2568 regional river planning initiative, the Mackenzie River Basin Board (MRBB), which has

2569 been in existence for over a decade. The Mackenzie River is the largest north-flowing
2570 river in North America and the longest river in Canada, draining nearly 20% of the
2571 country (Hipel et al. 2011). It is a unique global resource providing ecosystem services
2572 locally (biodiversity) and globally (climate regulation through carbon sequestration and
2573 freshwater flows to Arctic Ocean). Although it is among the least fragmented large-scale
2574 basins in North America, it is increasingly at risk from climate warming and from
2575 development pressures including oil and gas, mining, and hydroelectric power generation
2576 (Hipel et al. 2011; Rosenberg International Forum 2013). Initially, the basin's planning
2577 governance was complicated by jurisdictional discontinuity (de Loë 2010; Rosenberg
2578 International Forum 2013) and therefore a need for an effective transboundary water
2579 management plan was recognized (de Loë 2010). In response, the Mackenzie River Basin
2580 Transboundary Agreement was developed in 1997 among the governments of Canada,
2581 Saskatchewan, Alberta, British Columbia, Yukon and Northwest Territories. It was
2582 founded on guiding principles of equitable utilization, prior consultation, sustainable
2583 development, and maintenance of ecological integrity
2584 (www.mrbb.ca/information/31/index.html). Although the agreement has been in place
2585 since 1997, a recent report concluded that there has been little effective follow-through
2586 on the Master Agreement (Rosenberg International Forum 2013). However, the report did
2587 acknowledge that the MRBB was an appropriate model for integrated resource
2588 management, and could provide effective management if it was fully implemented
2589 (Rosenberg International Forum 2013). That report also suggested that the MRBB's
2590 effectiveness could be strengthened if it had increased authority for decision-making and
2591 an expanded scientific knowledge foundation for management actions with research

2592 priorities set through a science advisory panel (Rosenberg International Forum 2013). It
2593 is likely that this model could be adapted for additional basin management boards in the
2594 boreal zone. The experience of the MRBB highlights that integrated water management
2595 authorities will need to be properly resourced and empowered to turn their visions into
2596 effective action (Sanders and Wenig 2007; Wells et al. 2010; Rosenberg International
2597 Forum 2013).

2598

2599 **6. Summary and Conclusion**

2600 The boreal zone hosts more than half of Canada's total renewable supply of
2601 freshwater and the majority of Canada's wetlands (bogs, fen, swamps, marshes and
2602 shallow open waters). Boreal water-related ecosystem services are important to the well-
2603 being of society. Natural resource industries within the boreal, which contribute
2604 significantly to Canada's GDP and employ hundreds of thousands of Canadians, are
2605 dependent on a renewable supply of water. Without water to draw, we can hew no wood,
2606 mine no coal, or extract no oil (McLaughlin 2009).

2607 The boreal zone spans the continent and, consequently, there is substantial
2608 variation in the dominant hydrologic processes across this large biome. This variation is
2609 important to understanding how disturbances may (or may not) affect water resources.
2610 The western boreal zone is characterized by a sub-humid climate (precipitation \cong
2611 evapotranspiration), deep, permeable soils and surficial geology, and relatively low relief.
2612 Thus, hydrologic processes are dominated by vertical flows and evapotranspiration, with
2613 little runoff generation. In contrast, the eastern boreal zone is characterized by a humid
2614 climate (precipitation $>$ evapotranspiration), shallow confining layers or bedrock, and
2615 relatively high relief. Here the dominant hydrologic processes are lateral flows, with
2616 substantial runoff. These differences in hydrologic regimes must be taken into account to
2617 determine how a particular disturbance affects water quality or quantity.

2618 The literature on the risks to water quantity and quality posed by natural resource
2619 development, including roads, forest management activities, pulp and paper operations,
2620 mining, oil and gas extraction, peat mining and electric power generation, were reviewed
2621 and prognoses developed. Natural resource sectors have and continue to be responsible

2622 for significant changes to boreal water quantity and quality through extractions,
2623 diversions, hydrologic disruptions, effluents, seepages, and emissions. In general, the
2624 evolving system of regulations, guidelines, and standards for natural resource
2625 development are improving the prognosis for water resources for some sectors in some
2626 areas, but in other areas legacy impacts persist, deposition of atmospheric pollutants to
2627 receiving waters continues, and expanding natural resource development poses new risks
2628 to both water quantity and quality. Additionally, the potential confounding influences of a
2629 changing climate may exacerbate natural resource development impacts on water
2630 resources.

2631 The cumulative effects of natural resource development activities over space and
2632 time, coupled with the uncertainties related to climate change, have not been adequately
2633 examined in the published literature, and therefore critical information gaps remain
2634 (Gunn and Noble 2011). As natural resource development increases into previously
2635 unmanaged areas, the resilience of aquatic ecosystems in the boreal zone may be
2636 challenged, at least at local or regional scales. An integrated approach to natural resource
2637 management that includes economic and ecological valuation of water resources,
2638 improved water governance, and conservation-based basin management will lead to
2639 enhanced management of water resources (Schindler and Lee 2010). Continued efforts in
2640 coordination between regional, provincial, and federal governments and a unifying
2641 national water strategy could enhance the sustainability of Canadian boreal water
2642 resources for future generations.

2643

2644

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2646

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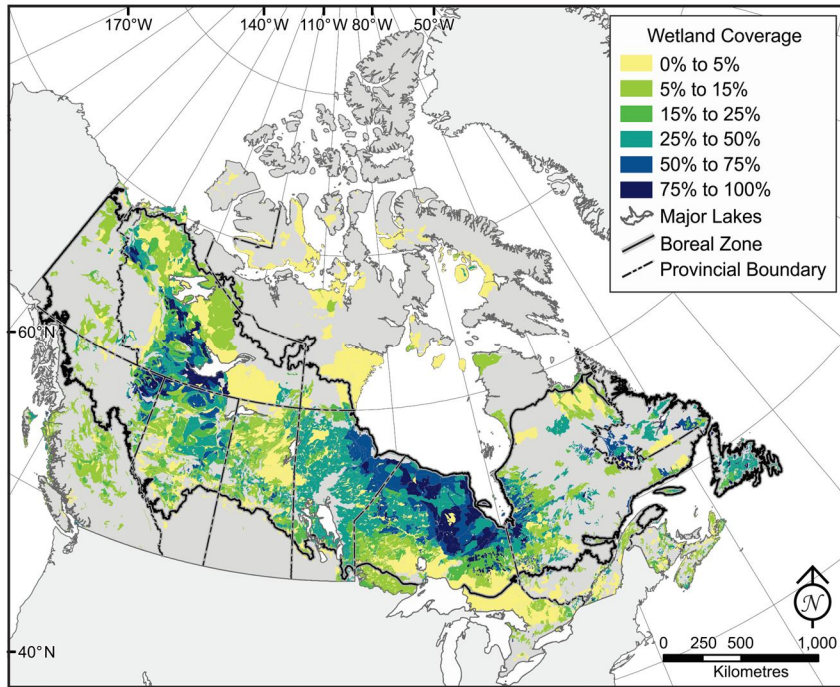
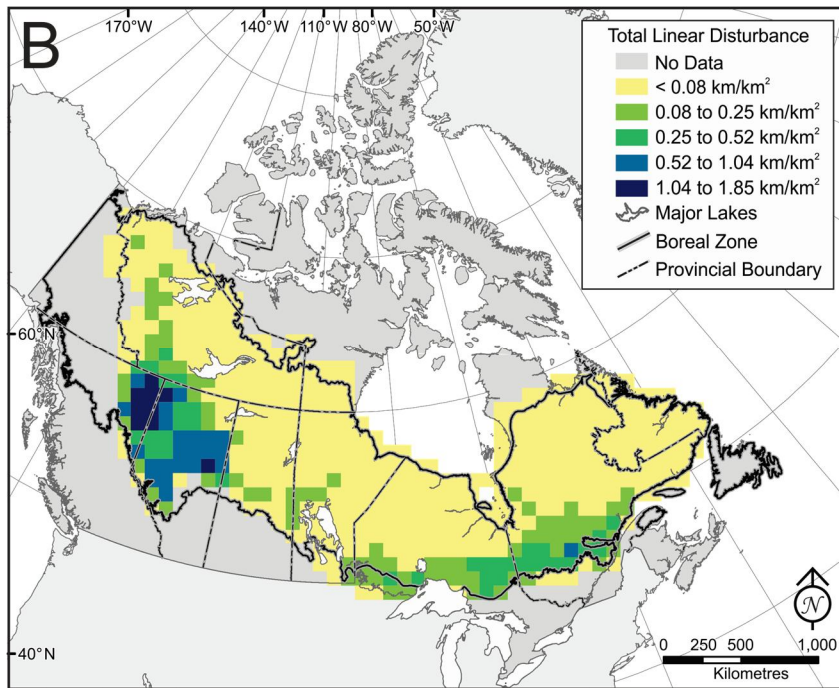
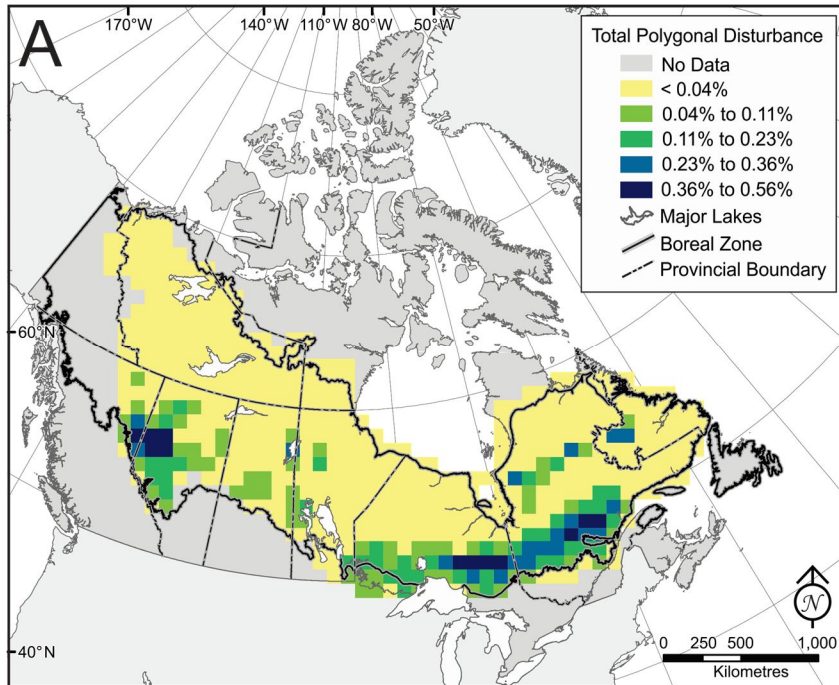
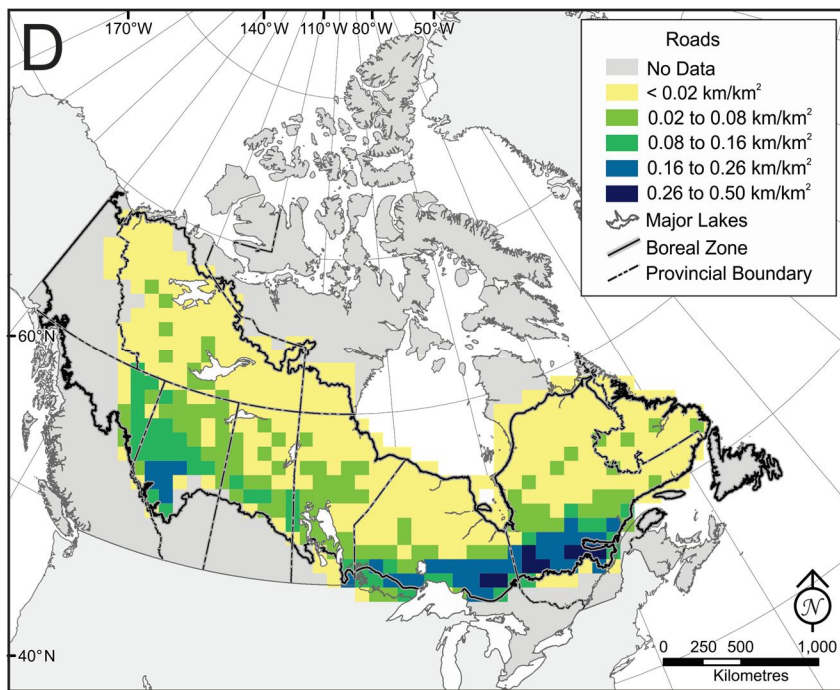
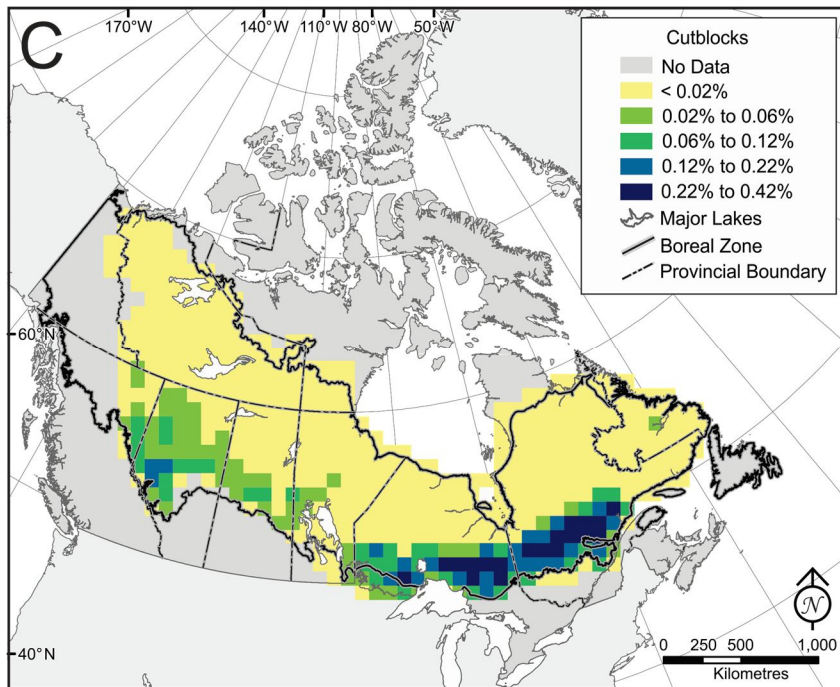


Figure 1:





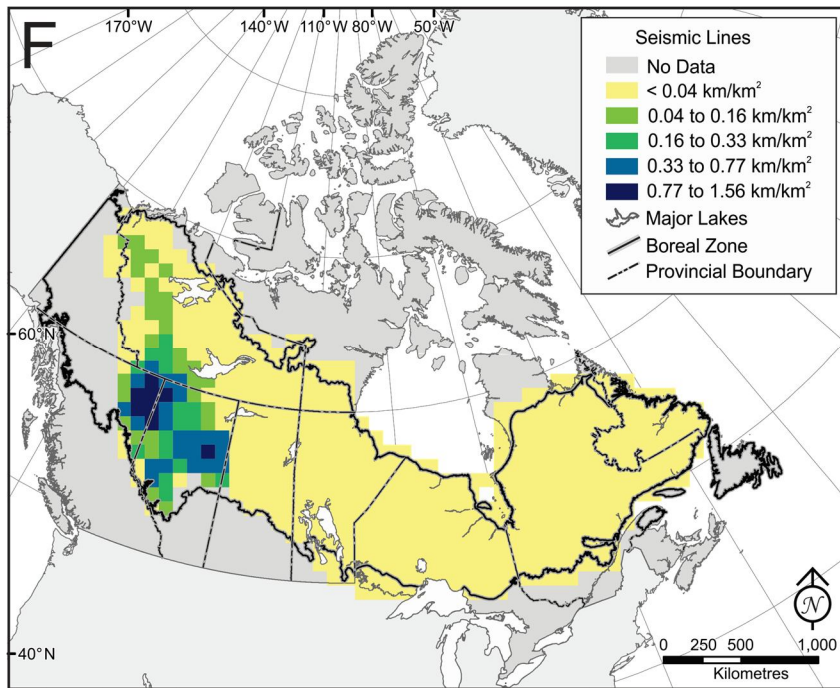
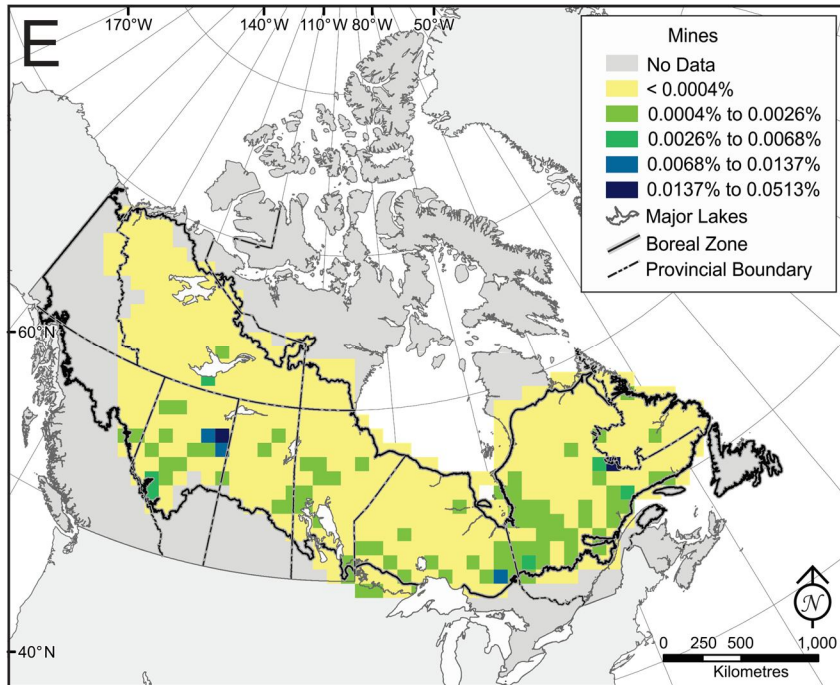


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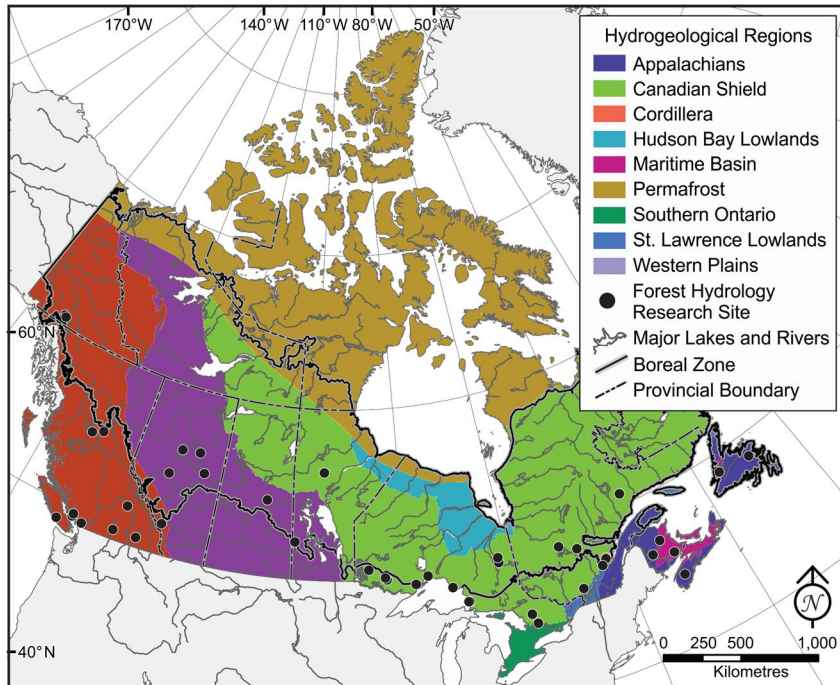
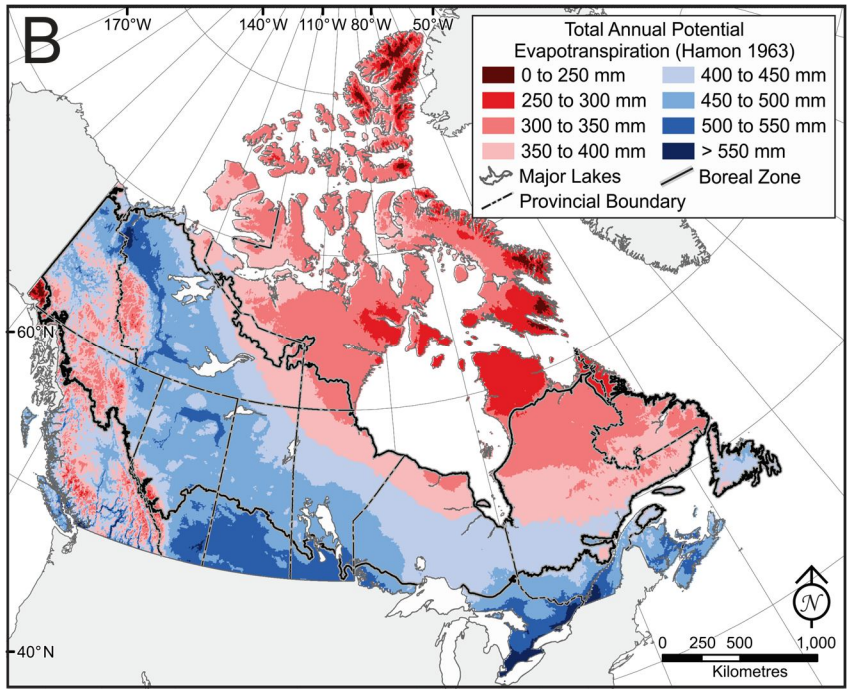
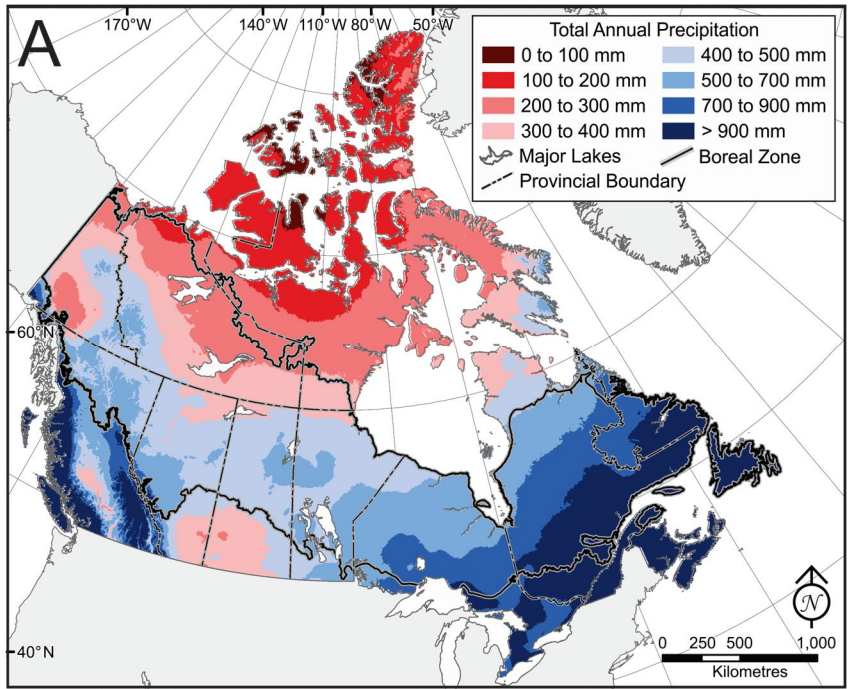


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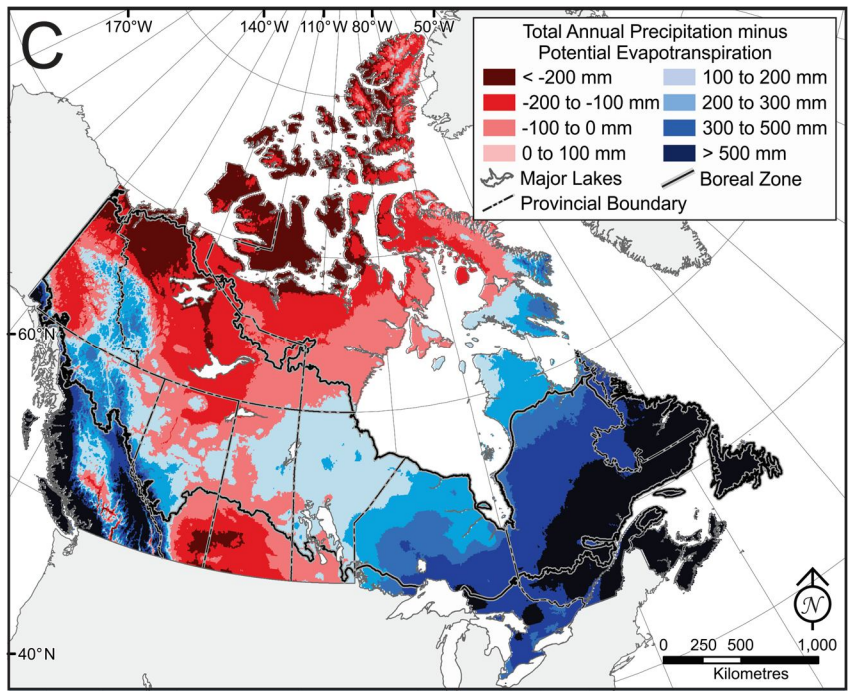
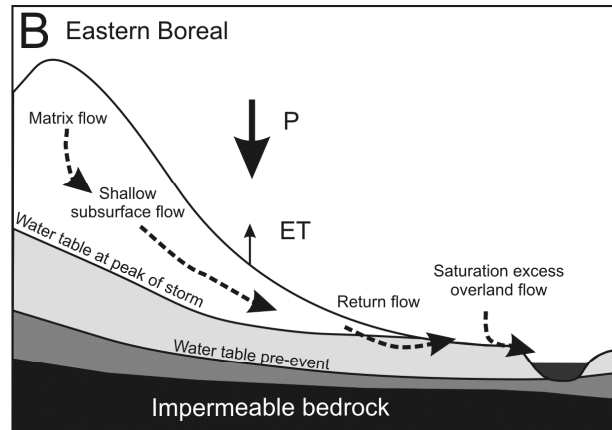
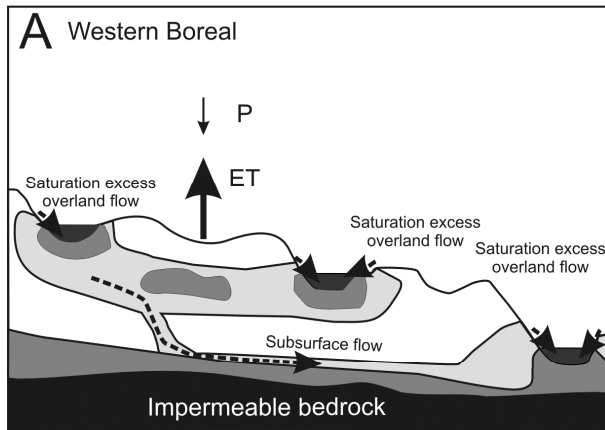


Figure 4:



- Surface water
- Water table pre-event
- Water table at peak of storm

Figure 5:

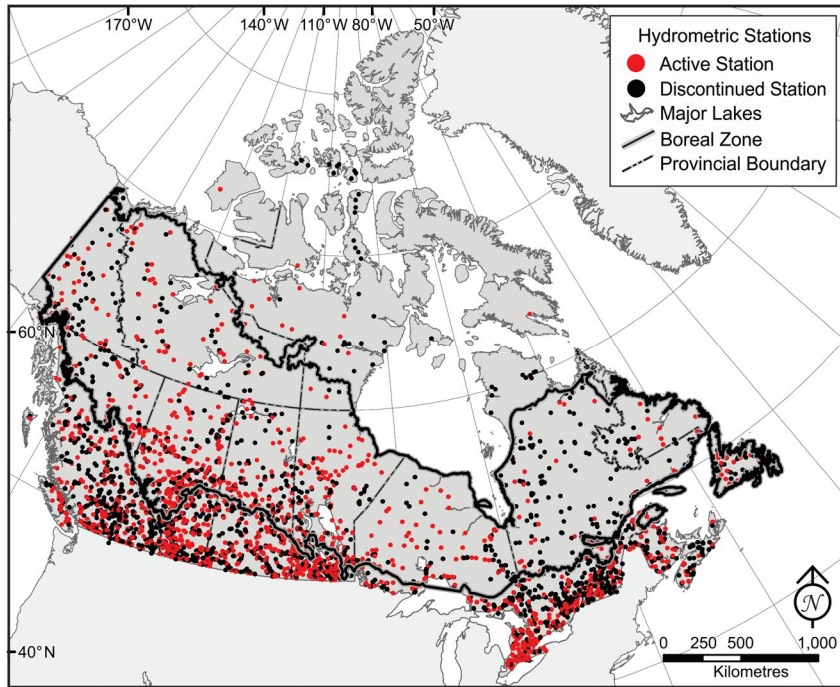


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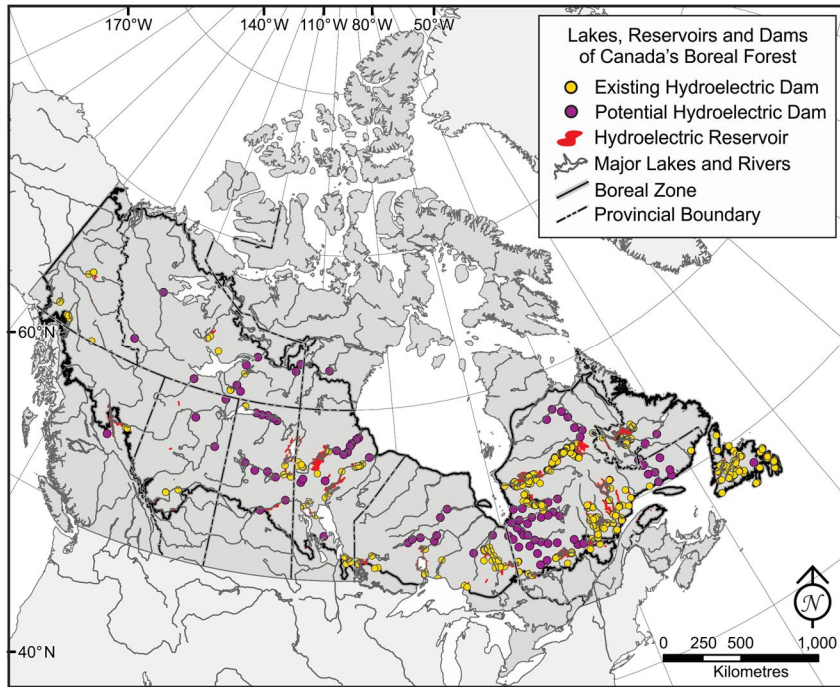


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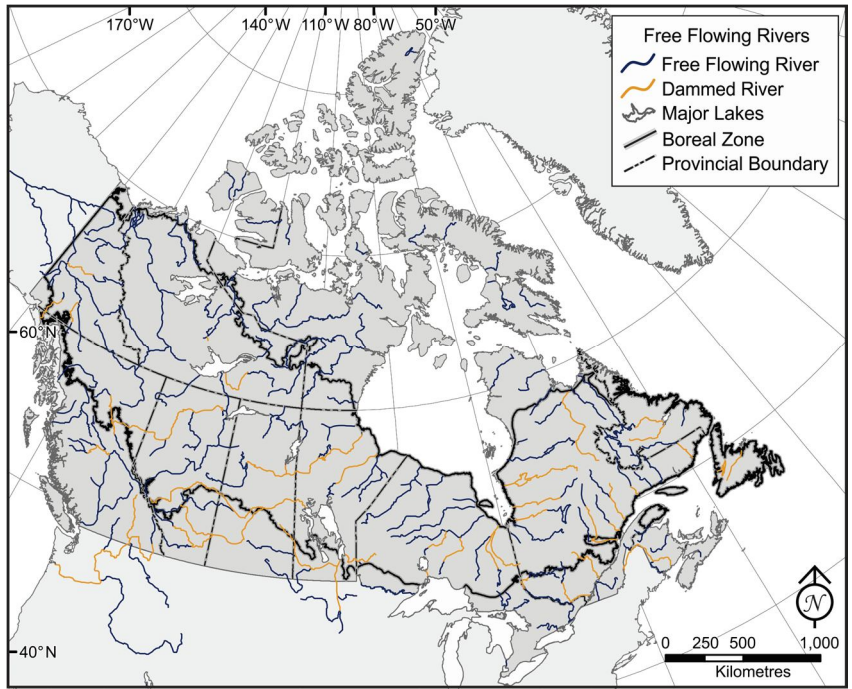


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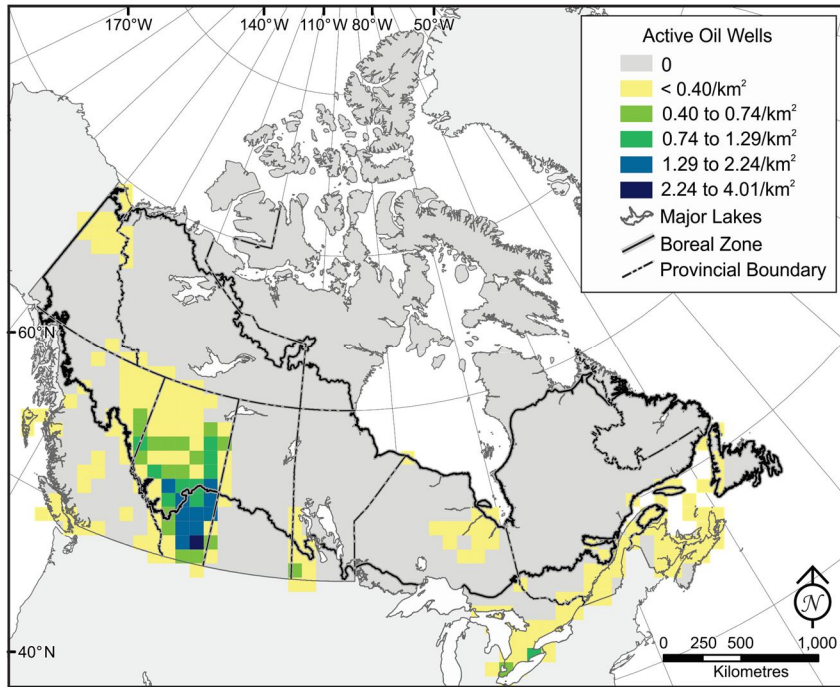


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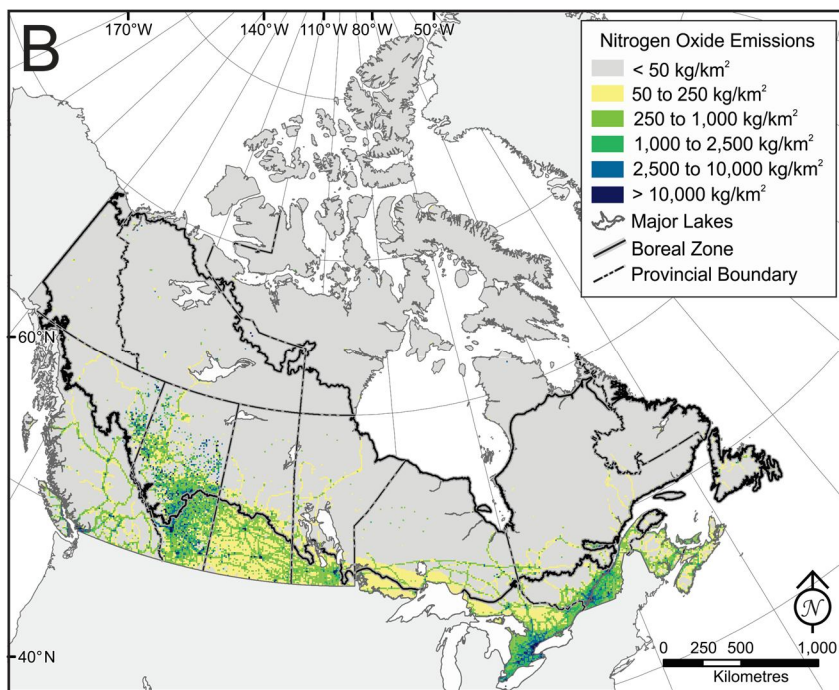
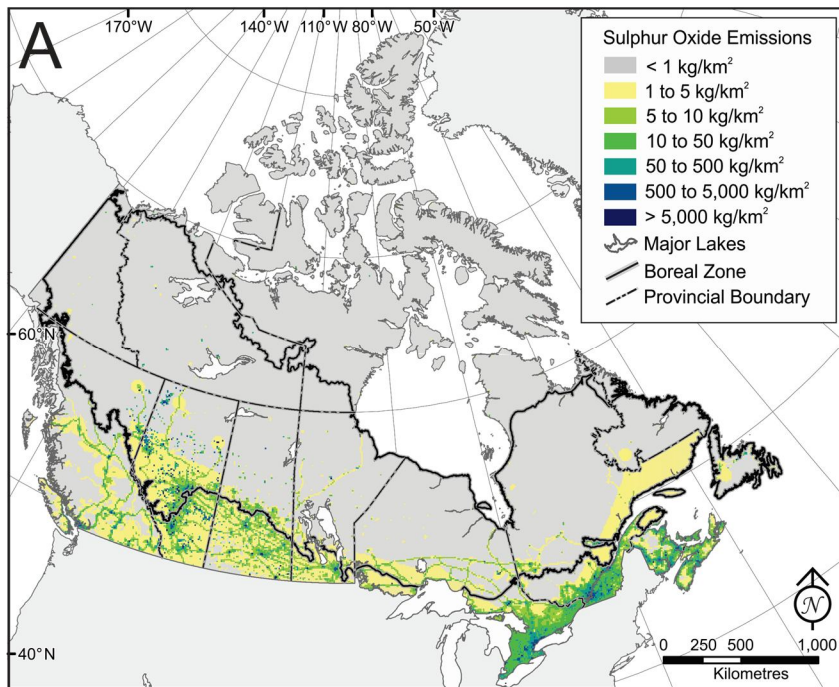


Figure 10:

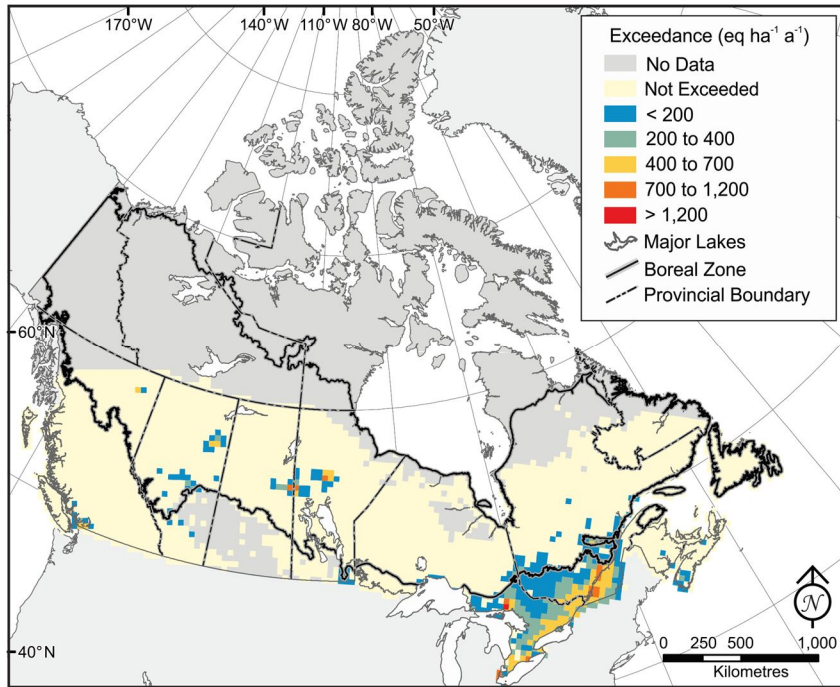


Figure 11:

Table 1: Boreal aquatic ecosystem services.

Service	Function		
	Wetland	Surface waters	Groundwater
Provisioning services			
Fibre and fuel/energy	Treed peatlands for forestry. Peat mining for peat moss products and bioenergy.	Water used in hydroelectric power generation and for cooling in thermal power generation.	NA ¹ .
Food	Used as food for people and domestic animals/wood, fur and medicine	Used as food for people and domestic animals, fur and medicine	NA ¹ .
Fresh water	Public and industrial water supply is obtained from reservoirs draining peatlands. Peatlands lost as freshwater source when drained for agriculture or forestry. Water quality compromised by waste disposal or landfill. Flooding of wetlands during reservoir creation lead to methyl mercury production.	Surface water used for public water supply, irrigation and resource extraction and processing. Transport of nutrients, organic matter and contaminants. Water quality compromised by waste disposal.	Groundwater used for public water supply, irrigation and resource extraction and processing. Surface water recharge. Transport of nutrients, organic matter and contaminants. Water quality compromised by waste disposal.
Regulating services			
Climate regulation	Regulation of greenhouse gases, regulation of climatic processes	Temperature moderation.	Temperature moderation indirectly through connections to surface water.
Water regulation	Water storage, groundwater recharge and discharge	Water storage, groundwater recharge and discharge	Water storage, groundwater recharge and discharge
Water purification and waste treatment	Retention, recovery and removal of excess nutrients and pollutants	Retention, recovery and removal of excess nutrients and pollutants	Retention, recovery and removal of excess nutrients and pollutants
Erosion protection	Peat blanket protecting the underlying soils from erosion	Local scouring of shoreline and re-deposition.	Maintenance of base flows.
Cultural Services			
Recreational and aesthetic	Opportunities for recreation and tourism; appreciation of nature	Opportunities for recreation and tourism; appreciation of nature	NA ¹ .
Spiritual and inspirational	Personal feelings and well-being; religious significance	Personal feelings and well-being; religious significance	Personal feelings and well-being; religious significance
Educational	Opportunities for education, training and research	Opportunities for education, training and research	Opportunities for education, training and research
Supporting services			
Biodiversity	Habitats for species	Habitats for species	Habitats for species
Soil formation	Accumulation of organic matter	Accumulation of sediments.	NA ¹ .
Nutrient cycling	Storage, recycling, processing and acquisition of nutrients	Transport and instream processing of nutrients	Transport and modification of nutrients.

¹. Ecosystem service is not applicable.

Table 2: Water use in Canada's natural resource sectors. Sources: NRTEE (2010), Statistics Canada (2010) and NRCan (2011).

	Forestry and Pulp and Paper	Electrical power generation		Mining	Oil & Gas
		Hydro Power	Thermal Power		
Contribution to Gross Domestic Product (%)	1.9	2.3		2.8	3.4
National Gross Water Use (%)	5	Unknown	64	4	9.6 ¹
National Consumptive Water Use (%)	2	Unknown	12	3	6.3 ¹
Regions Most Implicated	All provinces	BC, AB, MB, ON, QC, Atlantic	Territories, AB, ON, Atlantic	BC, AB, SK, Territories	AB, BC, SK

¹Petroleum and coal product manufacturing

Table 3: Summary of impacts of natural resource development on water resources in the boreal.

Development type	Local area impacted	Cumulative boreal area Impacted	Primary effects on water quantity	Primary effects on water quality	Length of effect (years)	Key references
Roads	Narrow, linear disturbance	Large (when you incorporate density) 93×10 ⁶ ha (Anielski and Wilson 2005) 600 000 km (Pasher et al. 2013)	Change to volume, timing and routing. Raise upslope water table, lower downslope water table	Sediment, contaminants and nutrients	100+	Trombulak and Frissel 2000; Whemple and Jones 2003; Partington and Gilles 2010
Forest Management	Large, irregular patches	Large 14.4×10 ⁶ ha (Pasher et al. 2013)	Increase (“watering up” from vegetation removal) or decrease (increase solar heating or draining as occurs in treed wetlands)	Change in nutrient export.	5-10 yr	Buttle and Metcalfe 2000; Buttle et al. 2000, 2005, 2009; Mallik and Teichert 2009; Smith et al. 2007; See also Table 5
Pulp and Paper Industry	Small site disturbance but discharge of effluents has greatest impact over larger scale	Small to Medium (when incorporate downstream effects) 17 mills × 100 ha approximate footprint = 1 700 ha but downstream impacts 20+ km	Small decrease (most flow returned)	Contaminants and nutrients Temperature increase	>10 yr	Chambers 2001; NCASI 2010
Electricity Generation	Medium (thermal) to large (hydroelectric)	Large (including downstream effects) 130 000 km of rivers affected by large dams (McAllister 2000) 5.2 × 10 ⁶ ha reservoir area (Lee et al. 2012)	Increase upstream (flooding) and decrease downstream (lower and irregular flows)	MeHg and nutrient export (hydroelectric) Temperature alterations (thermal)	100+	Baxter 1977; Urquizo et al. 2000; Dynesius and Nilsson 1994; Urquizo et al. 2000
Minerals and Metal Mining	Medium to large depending on mine size and mode of extraction	Large 123 active + 1300 abandoned mines × 1000 ha approximate footprint = 1.4 × 10 ⁶ ha	Decrease in stream flow due to removals. Increase where tailings ponds created Groundwater table declines	Contaminants and nutrients	1000+	Ptacek et al. 2004; Keller et al. 2007

			due to dewatering			
Oil and gas and oil sands development	Small (individual wells) to large (oil sand intensive surface extraction)	Large (when you incorporate density) 46 ×10 ⁶ ha (Timoney and Lee 2001) with 56 000 ha of actively mined oil sand (Alberta Environment http://environment.alberta.ca/02863.html)	Decrease in stream flow due to removals. For surface extraction of oil sand end pit lakes created, but groundwater table declines due to dewatering	Contaminants and nutrients	1000+	Woynillowicz et al. 2005; Gosselin et al. 2010; Kasperski and Mikula 2011; CEMA 2012; Government of Alberta 2012b, 2013a
Peat mining	Small local impact although draining may affect larger area	Small 17 000 ha (Daigle and Gautreau-Daigle 2001)	Decrease (during drainage) and increase (when drains plugged)	Nutrients	1000+	Daigle and Gautreau-Daigle 2001; Price et al. 2005; Gleason et al. 2006

Table 4. Description of boreal hydrology research sites in boreal including information on duration, treatments, and key references.

Site	Location	Area	Project Duration	Pre-harvest years	Year of Harvest	Post-harvest	Forest Operations	Research Infrastructure	Hydrologic and water quality metrics	Primary Reference (study design)
Wolf Creek, YT	60°32'N 135°07'W	14.5 – 195 km ²	1992-present	NA	NA	NA	undisturbed system	-3 met. stations -piezometers -soil moisture meas. -3 stream gauges (nested) -stream chemistry -25 snow courses (monthly) -water quality (weekly) -contaminant sampling: water, snow & fish	-daily avg. SWE	Pomeroy & Granger 1999
Spring Creek, AB	54°56'N 117°43'W	1,860 – 1,971 ha	1966 - 2000	1966 - 1993	1994	1977 - 2002	control and clearcut	-Discharge -Air Temp -Precipitation -Other Met Variables - soil moisture and temp	- evapotranspiration - sediment yield	Swanson and Rowell 2001 Martz 1988 Watertight Solutions Ltd. 2005
Tri Creek, AB	53°16'N 117°14'W	1,480 – 2,820 ha	1969 - 1985	3	1974 - 1983	1983 - 1985	yes, unknown treatments	-Discharge -Air Temp -Precipitation -Water Chemistry -Water Temp -Sediment -Other Met Variables	NA	Currie 1976 Stevens 1979 Bergstrom and Nip 1988 Nip 1991 Van Der Vinne 1992 Andres et al. 1987
Hydrology, Ecology and Disturbance (HEAD) (aka USRA Utikuma Study Research Area), AB	56°06'N 116°32'W	0.05 – 5 km ²	2001-2004	NA	NA	NA	undisturbed system	@ 24 lakes -lake level -standard limnological variables: TP, TN, chlorophyll a, DOC, conductivity (3 x/yr, May, Jul, Aug) @ 3 intensively studied lakes + catchments -met station: T, P, wind -hundreds of piezometers (some w. continuous head meas.) (manual meas. 1-2x per month during ice free season, otherwise every two months) -some deep groundwater wells -lake w seepage meters -lake levels, inflow and outflow meas.	-no flow data -groundwater fluctuation at one intensively measured site -saturated and inundated areas measured from ERS satellite	Smerdon et al. 2005
Terrestrial and Riparian Organisms, Lakes, and Streams (TROLS), AB	54°46'N 111°59'W	189-5669 ha	1994-1998	1994-1995	1996-1997	1998	-3 controls with little to no logging -8 logged basins ranging from 9-35% clear cut	-12 lakes with limnological & aquatic ecology -3 WSC stream gauges -met. data from nearby EC sites -water quality measurements (2x/month) -chemical analysis -aquatic ecology -fish studies	-median total Phosphorus concentration [TP] used to illustrate influence of surface and subsurface hydrologic connections and impacts of logging on lakes	Devito et al. 2000 Prepas et al. 2001a

Site	Location	Area	Project Duration	Pre-harvest years	Year of Harvest	Post-harvest	Forest Operations	Research Infrastructure	Hydrologic and water quality metrics	Primary Reference (study design)
Forest Watershed and Riparian Disturbance (FORWARD), Swan Hills, AB	54°12'N 115°46'W	131 – 248 km ²	1998-2004	1983-1997 (WSC)	1998	1999-2000	-89% burned 1998, 10% logged	-3 WSC gauges (day) -4 met. stations (P, T, rel H, wind speed & direction., incoming solar, plus 8 additional P gauges) -stream chemistry – ISCO's (1-2 week grab samples) -1 EC met. station (Whitecourt)	-baseflow, summer stormflow -TP export	Smith et al. 2003 Prepas et al. 2003
Duck Mountain , MB	51°01'W 100°39'N	103-4230 km ²	-earliest WSC site is 1954; range is 10 – 40 years -3 active gauges, 11 discontinued				-Louisiana Pacific is license holder	-14 WSC gauges drain Duck Mountain, seasonal (day) -9 MSC met. stations in the region	-water yield -seasonal runoff -peak flows -daily Q	unpublished draft report by Watertight Solutions Ltd. 2005
BOREAS , Thompson MN, & Prince Albert SK	55.91°N 98.49°W 53.86°N 104.62°W	NSA 8000 km ² SSA 11170 km ²	1993-1997				undisturbed system	- discharge - air temp - precipitation - other met variables - numerous stream gauges - numerous met. measurements - remote sensing - soil moisture - snow measures - flux towers	-water balance -ksat -SWE -daily hydrograph separations	Sellers et al. 1995 Metcalf & Buttle 1999, 2001
Coldwater Lakes , ON	49°10'N 92°10'W	70-194 ha	1990-1999	1990-1995	1996 & 1998	1999	L26 – 45% clear cut no shoreline disturbance L39 – 77% clear cut with shoreline disturbance L42 – 74% clear cut with shoreline disturbance -2 phase logging 1996 & 1998 - feller buncher & chain saw -aerially seeded in 1998	-stream gauges at inflow and outflow of 3 headwater lakes -1 land met. Station (1993), additional floating met stations on lakes -water quality & chemistry measurements (bi-weekly) -aquatic ecology (monthly May to Sep) -lakes not hydrologically independent, single lake approach	-limnological study of water quality	Steedman 2000 Nicholls et al. 2003
Pukaskwa River (aka White River), ON	48°03'N 85°48'W	537 km ² (total area)	2002-present WSC: 1940 - present	2002-2005	2006	2007	-logging within the watershed starting 2000 -logging within experimental catchments in 2006	-1 WSC & 3 OMNR seasonal stream gauges -1 OMNR & 1 GLFC seasonal met. stations	NA	-no publications to date

Site	Location	Area	Project Duration	Pre-harvest years	Year of Harvest	Post-harvest	Forest Operations	Research Infrastructure	Hydrologic and water quality metrics	Primary Reference (study design)
Current River, ON	48°33'N 89°14'W	403 km ² (total area)	2002-present WSC: 1968 - present	2002-2003	2004	2005-present	-extensive logging began 1995 in main watershed -logging in experimental watershed 2004	-1 WSC (1968) & 2 OMNR (2002) stream gauges -1 met. station at Lakehead University	NA	-no publications to date
Little Abitibi, ON	49.07°N 81.03°W	1408 km ² (total area)	2003-present WSC: 1955 - present			2003 - present	-extensive logging using various methods throughout	-20 Trent stream gauges (2003) -3 met. stations -10 standard rain gauges	NA	-no publications to date
Esler lakes Research Area, ON	49.38°N 81.01°W	200 km ² (total area)	2001-present	2001-2004	2005 (winter)	2005 - present	-extensive logging using clear cut (dominant) and variable retention methods	- 14 lakes with piezometers 10 lakes each with 3 zero-tension lysimeters and 6 tension (porous cup) lysimeters - Three rain gauges in study area - Water level indicators on 21 lakes	- Deep and shallow soil groundwater movement and chemistry - O ¹⁸ measurements on groundwater - Lake water chemistry (2001-present) - Seepage meter measurements (2006) - Mini piezometer measurements of lake sediments - Snow melt and runoff data for 3 intensive lakes (2004-present) - Throughfall (terrestrial and nearshore)	1) Hazlett et al 2005 2) Hazlett et al 2006 3) Starr et al 2006
Lac Laflamme, Montmorency Forest, QC	47°19'N 71°07'W	68 ha	1981-1997	NA	NA	NA	undisturbed system	-18 -27 snow sampling stations -9 stream gauging stations (hourly) -1 met. station: air T & P -snow & soil moisture collected 2-7 days Mar-May -lake levels -modelling: VSAS2	-daily Q and snow water content -peak flow	Prévost et al. 1990

Site	Location	Area	Project Duration	Pre-harvest years	Year of Harvest	Post-harvest	Forest Operations	Research Infrastructure	Hydrologic and water quality metrics	Primary Reference (study design)
Ruisseau des Eaux-Vollées Experimental Watershed (REVIEW) , Montmorency Forest, QC	47°16'N 71°09'W	122 - 917 ha	1967-1998	1) 1967-1974 2) 1985-1992	1) 1974-1975 2) 1993	1) 1976-1978 2) 1994-1998	1) 1974-1976 single patch cutting 31% of 6 2) 1993 85% basal area removal thru clearcut 7A -harvester & forwarder -5-20 m buffers	-3 stream gauges (nested): Main (5D), 6, 7A ; V-notch weirs (hr) -paired-watershed & single basin approaches 1) 6 was treatment & 7 was control 2) Second experiment was reverse	-peak flows -quick flow volume -baseflow volume -total storm flow volume -lag time -concentration time -rise time -falling time -base time	Guillemette et al. 2005 Plamondon & Ouellet 1980
Cote Nord & Haute – Mauricie , QC	50°43'N 67°30'W 48°31'N 73°25'W	0.83 – 4.09 km ²	1974-1977	1974	1975 1977	NA	Cote Nord (1975) - 7A, 7B, 7C, 5A & 5C controls - 5G 62% clearcut with 10 m buffer - 5H 26% clearcut no buffer -chain saws & rubber tire skidders - 5G & 5H are nested in 5C Haute Mauricie (1977) - Gilbert 100% clearcut no buffer - Huguette 40% clearcut 30 m buffer - Wajusk control -chain saw & skidders and feller buncher	-suspended sediments -water T -water quality and chemistry -paired basin	-concentration of inorganic suspended sediments - CA, K, DO, FE -pH, colour and conductivity	Plamondon et al. 1982
Copper Lake , NF	48°50'N 57°50'W	114 – 124 ha Copper Lake – 13.5 km ²	1993-1997	1993	1994-1996	1996-1997	1994- road construction through both wsheds T1-1 – 18% clearcut no buffer T1-2 26% clearcut w. 20m buffer T1-3 & T1-5 , controls -2 phases winter 1994-1995 & summer 1996	-stream surveys -lake bathymetry -stream gauge at outflow from Copper Lake (WSC) -water chemistry (monthly) -fish studies -stream T (hourly), 10 recorders -suspended sediment sampling (~2x/year) -benthic invert surveys (annual)	-average monthly discharge -seasonal T	Curry et al. 2002 Clarke et al. 1997

Site	Location	Area	Project Duration	Pre-harvest years	Year of Harvest	Post-harvest	Forest Operations	Research Infrastructure	Hydrologic and water quality metrics	Primary Reference (study design)
Triton Brook, NL & LB	48°36'N 54°35'W	23,297 ha	July 2001 – present	2001	2002	2002-present	-mechanical harvesting, fire, infestation over last 100 years -no experimental harvest -reforestation efforts ongoing (scarification & replanting)	-water quality at 3 sites, manual, (bi-weekly Jun to Oct) -WSC stream gauge (daily) -modelling; single basin	-daily Q	NA

Table 5: Impacts of forest management practices on water quality.

Parameter	Disturbance	Location	Length of study	Impact	Author	
Temperature of water	Wind; forestry	BC, Stuart-Takla	6 years	+ ; + +	Macdonald et al, 2003	
	Wind; forestry	Ontario, Coldwater lakes	8 years, 5 years pre-harvest, 3 years post-harvest	+ Thermocline depth	Steedman and Kushneriuk, 2000	
	Forestry	Ontario, White River	4 years, 2 pre-harvest 2 post-harvest	+ + without buffers	Kreutzweiser et al, 2009b	
	Forestry	Ontario, Coldwater lakes	4 years, 1 year pre-harvest, 3 years post-harvest	0	Steedman et al, 2001	
	Forestry	Quebec, Montmorency forest	5 year study, 2 years pre-harvest, 3 years post-harvest	0	Tremblay et al, 2009	
	pH	Fire; forestry	Quebec, Gouin Reservoir	3 years immediately post-fire; post-harvest	0; 0	Carignan et al, 2000a
		Fire; forestry	Quebec	3 years immediately post-fire; post-harvest	0; 0	Garcia et al, 2007
		Beaver dam; forestry	Alberta; Rocky Creek	3 years	+; 0	Hillman et al, 1997
		Forestry	Quebec, Laurentian Mountains	5 months, post-harvest	0	Hausmann and Pienitz, 2009
		Forestry	Alberta, TROLS study area	4 years, during harvest	0	Prepas et al, 2001a
	Forestry	Ontario, ELA	3 years, post-harvest	0	Steedman, 2000	
	Forestry	Quebec, Montmorency Forest	5 years, 2 years pre-harvest, 3 years post-harvest	-	Tremblay et al, 2009	
Nutrient (N,P)	Fire; forestry	Quebec, Gouin Reservoir	3 years, post-fire; post-harvest	+ P, + N; + P, + N	Carignan et al, 2000	
	Fire; forestry	Quebec, Gouin Reservoir	3 years, post-fire; post-harvest	+ P, + N; + P, + N	Garcia et al, 2007	
	Fire; forestry	Quebec, Gouin Reservoir	3 years, post-fire, post-harvest	+ P, + N; 0 P, 0 N	Garcia and Carignan, 2000	
	Fire; forestry	Quebec, Haute-Mauricie	3 years, post-fire; post-harvest	+ P, + N; + P, + N	Lamontagne et al, 2000	
	Fire; forestry	Alberta, Caribou Mountains	1 year, 2 years post-fire	+ P, + N; + P, + N	McEachern et al, 2000	
	Forestry	Quebec, Gouin Reservoir	3 years, some during harvest	0 P 0 N	Bertolo and Magnan, 2007	
	Forestry	Alberta, Moose Lake	1 year, post-harvest	0 P	Evans et al, 2000	
	Forestry	Ontario, Experimental Lakes	1 year	- - P	France et al, 1996	

Parameter	Disturbance	Location	Length of study	Impact	Author
	Forestry	Quebec	5 months, post-harvest	+ P + N	Hausmann and Pienitz, 2009
	Forestry; beaver dams; roads	Alberta	3 years, during harvest	0 N; + N; 0 N	Hillman et al, 1997
	Forestry	Ontario, White River	5 weeks, post-harvest	0 P 0 N	Kreutzweiser et al, 2008a
	Forestry	Alberta	4 years, 2 years pre-harvest, 2 years post-harvest	+ P 0 N	Prepas et al, 2001a
	Forestry	Ontario, Coldwater Lakes	8 years, 5 years pre-harvest, 3 years post-harvest	0 P + N	Steedman, 2000
	Forestry	Quebec, Montmorency Forest	5 years, 3 years pre-harvest, 2 years post-harvest	0 P + + N	Tremblay et al, 2009
	Forestry	Quebec	2 years, 1 year pre-harvest, 1 year post-harvest	+ P 0 N	Winkler et al, 2009
DOC	Fire; forestry	Quebec, Gouin Reservoir	3 years post-fire; post-harvest	0 ; +	Carignan et al, 2000
	Fire; forestry	Ontario, Quebec	13 year study, 4-13 years post- burn or harvest	+ ; +	France et al, 2000
	Fire; forestry	Quebec	3 years, post-fire and post- harvest	0 ; +	Garcia et al, 2007
	Fire; forestry	Quebec, Haute-Mauricie	3 years, post-fire and post- harvest	0 ; +	Lamontagne et al, 2000
	Forestry	Quebec	3 years, during harvest	+	Bertolo and Magnan, 2007
	Forestry	Ontario, Experimental Lakes	1 year	- -	France et al, 1996
	Forestry	Quebec, Laurentian Mountains	5 months, post-harvest	+ +	Hausmann and Pienitz, 2009
	Forestry; beaver dams; road	Alberta	3 years, 1 year pre-harvest, 2 years post-harvest	0 ; - ; 0	Hillman et al, 1997
	Forestry	Ontario, Cold-water lakes	4 years, post-harvest	0 (any increase seen could not be attributed to cut)	Knapp et al, 2003
	Forestry	Ontario, Turkey Lakes	5 weeks, post-harvest	0	Kreutzweiser et al, 2008
	Forestry	Ontario, Cold-water lakes	4 years, 2 years pre-harvest, 2 years post-harvest	+	Steedman, 2000
	Forestry	Quebec	2 years, 1 year pre-harvest, 1 year post-harvest	+	Winkler et al, 2009

Parameter	Disturbance	Location	Length of study	Impact	Author
	Forestry	Quebec	2 years, 1 year pre-harvest, 1 year post-harvest	0	Desrosiers et al, 2006a
Dissolved oxygen	Forestry	Ontario, Coldwater lakes	8 years, 5 years pre-harvest, 3 years post-harvest	0	Steedman and Kushneriuk, 2000
	Forestry	Quebec	2 years, 1 year pre-harvest, 1 year post-harvest	0	Winkler et al, 2009
Base cations	Fire; forestry	Quebec, Gouin Reservoir	3 years post-fire; post-harvest	+ ; 0	Carignan et al, 2000
	Fire; forestry	Quebec	3 years post-fire; 3 years post-harvest	+ ; 0	Garcia et al, 2007
	Fire; forestry	Quebec, Gouin Reservoir	3 years post-fire; post-harvest	+ ; +	Lamontange et al, 2000
	Fire; forestry	Quebec, Gouin Reservoir	3 years post-fire; post-harvest	0 ; 0	Garcia and Carignan, 2000
	Beaver dam; forestry; roads	Alberta	3 years	+ ; 0; 0	Hillman et al, 1997
	Forestry	Ontario	5 weeks post-harvest	0	Kreutzweiser et al, 2008a
	Forestry	Ontario, Coldwater Lakes	8 years, 5 years pre-harvest, 3 years post-harvest	-	Steedman, 2000
	Forestry	Quebec, Montmorency Forest	5 years, 3 years pre-harvest, 2 years post-harvest	+	Tremblay et al, 2009
Acid anions	Fire; forestry	Quebec, Gouin Reservoir	3 years post-fire; post-harvest	+ + ; 0	Carignan et al, 2000
	Fire; forestry	Quebec, Gouin Reservoir	3 years post-fire; post-harvest	+ ; 0	Lamontagne et al, 2000
	Fire; forestry	Canada, meta-analysis	Varied	+ ; -	Nitschke, 2005
	Forestry	Quebec, Laurentian Mountains	5 months post-harvest	0	Hausmann and Pienitz, 2009
	Forestry	Alberta, TROLS area	4 years, harvest at year 2	0	Prepas et al, 2001a
	Forestry	Ontario, ELA	8 years, 5 years pre-harvest, 3 years post-harvest	0	Steedman and Kushneriuk, 2000
	Forestry; fire	Quebec, Gouin Reservoir	3 years, post-harvest; 3 years, post-fire	0 ; +	Garcia and Carignan, 2000